

Seismic wave Propagation and Imaging in Complex media: a European network

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Early Stage Researcher



Host Institution: Utrecht University
Place of Origin: Istanbul, Turkey
Appointment Time: February 2005

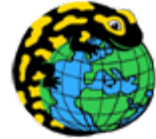
Project: Testing and Improving Tomographic Models Using
Numerical 3D Wave Propagation

Task Groups: TG Planetary Scale

Cooperation: Oxford University



Universiteit Utrecht



SPICE

Assessment of the tomographic mantle models using SEM seismograms

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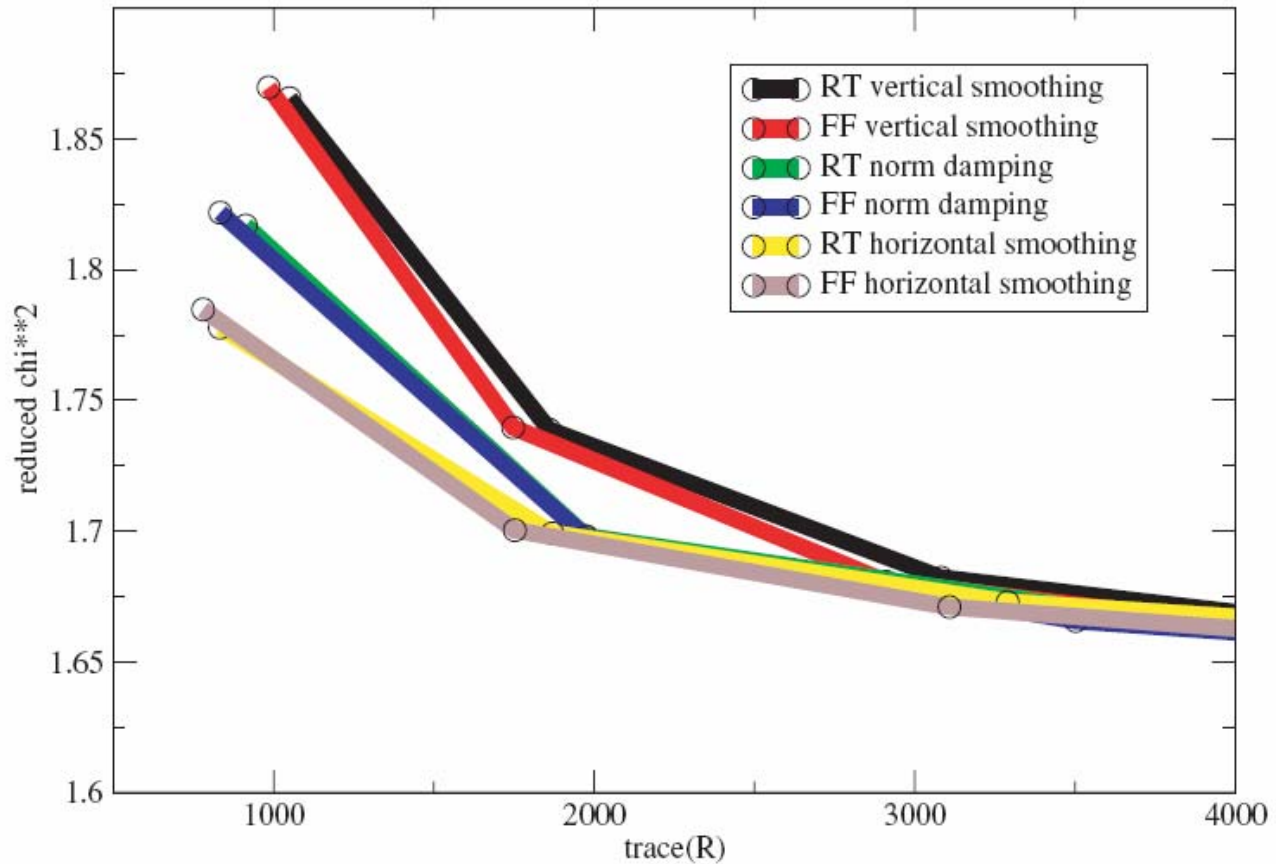
Utrecht University, Utrecht, the Netherlands

Importance of damping in inversion

$$\mathbf{m} = (\mathbf{A}^T \mathbf{C}_d^{-1} \mathbf{A} + \gamma \mathbf{D})^{-1} \mathbf{A}^T \mathbf{C}_d^{-1} \mathbf{d}$$

- Regularization is needed to make the inversion stable.
- Damping factor (γ) controls the smoothness of the models.
- Choice of γ depends on subjective observations.

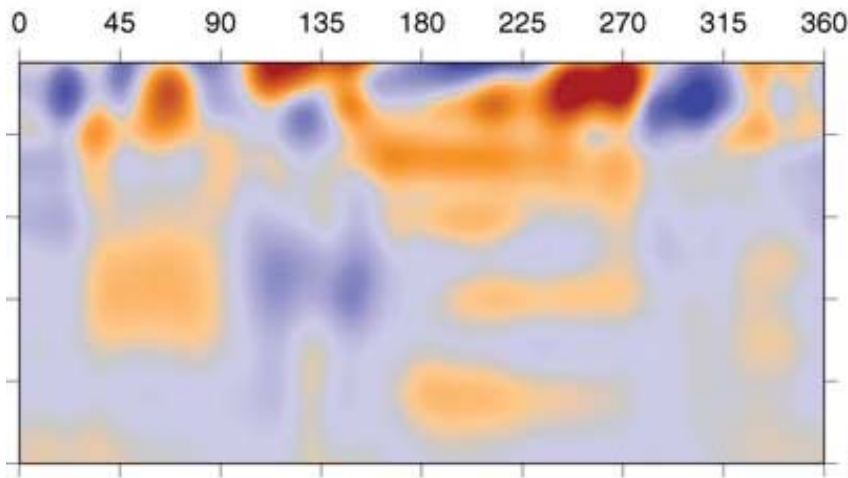
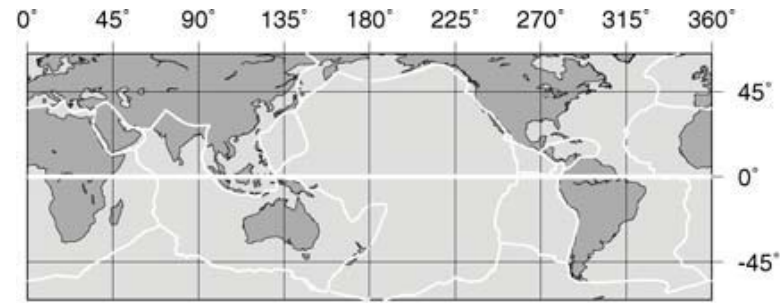
Sample trade-off curves



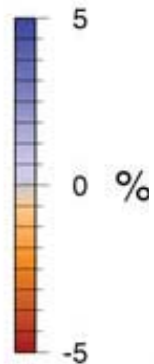
Trampert & Spetzler (2006)



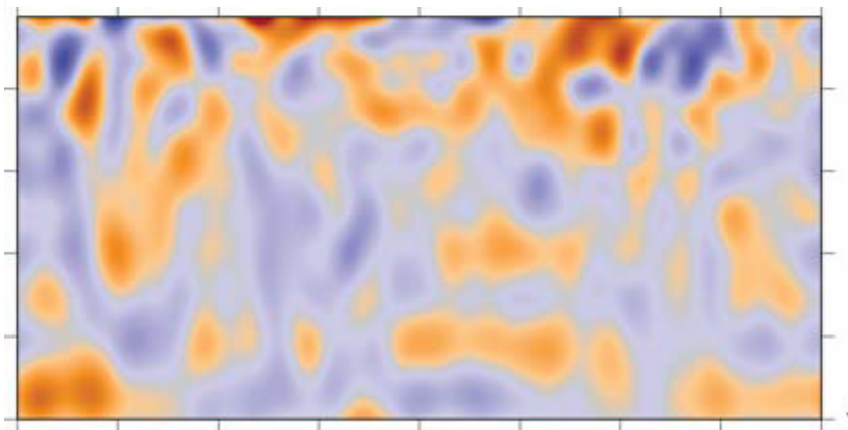
Cross-sections along the equator of the mantle models from Trampert & Spetzler (2006)



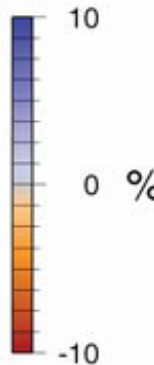
horizontal smoothing



$$\gamma = 0.01$$



horizontal smoothing



$$\gamma = 0.0001$$



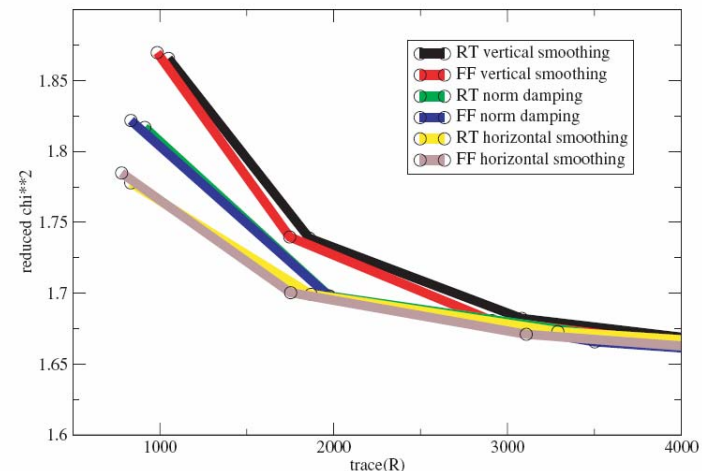
The objective

- Investigate the effect of **varying levels of damping** on tomographic mantle models.
- Check the agreement between real and **SEM (Komatitsch & Tromp 2002)** seismograms.

Synthetic seismograms using SEM

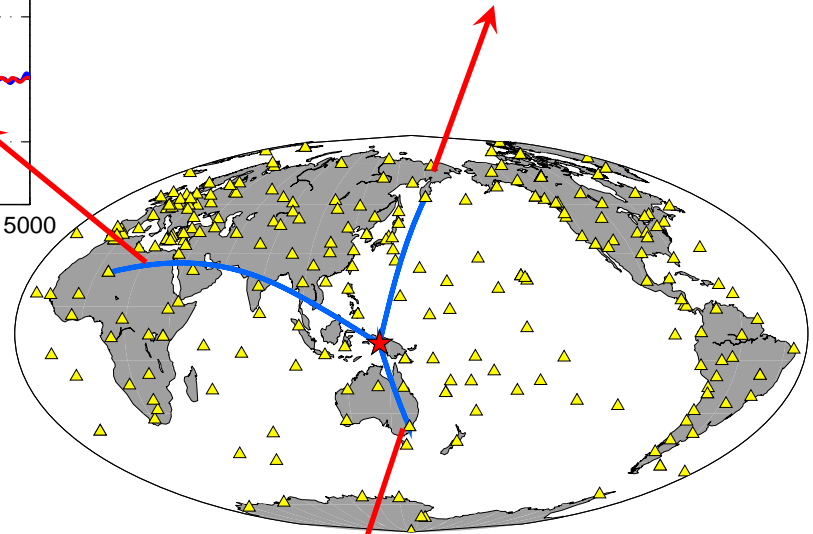
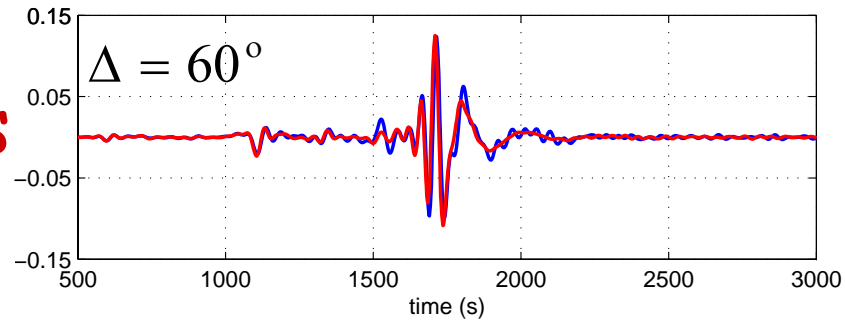
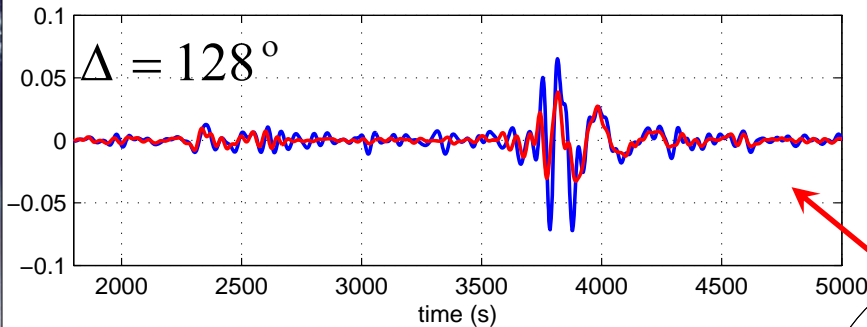
- Crust2.0 (Bassin et al. 2000)
- Four mantle models (Trampert & Spetzler 2006)

- $\gamma = 0.01$
- $\gamma = 0.001$
- $\gamma = 0.0001$
- $\gamma = 0.00001$



- PREM (Dziewonski & Anderson 1981)

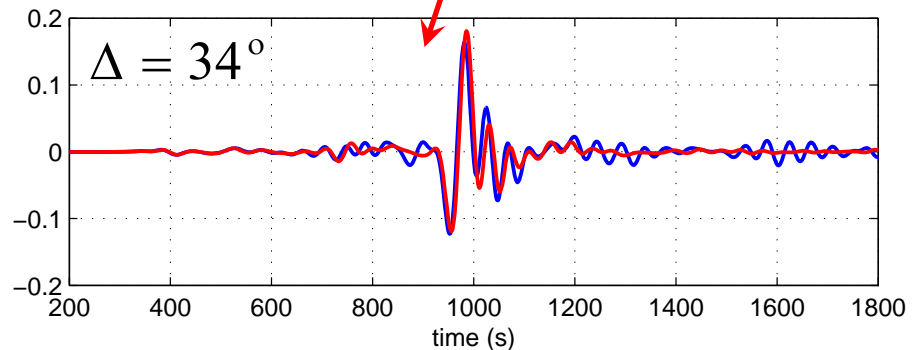
Comparison of real - SEM seismograms



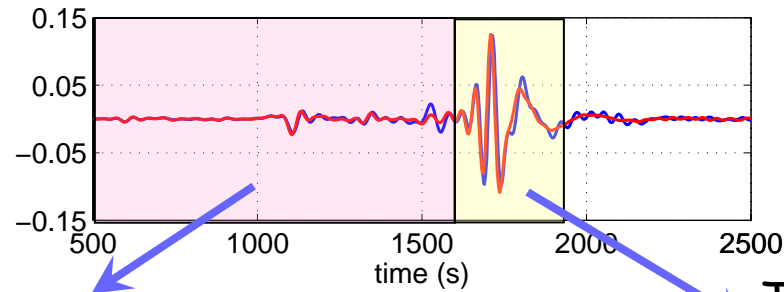
★ Irian-Jaya (2004) $M_w=7$

Blue: real seismograms

Red: SEM seismograms



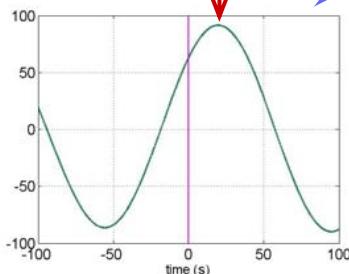
Measuring time shifts as a function of frequency



Windowing the body waves

Cross-correlation
of the seismograms

Narrowband filter
at desired periods
250 s - 40 s



$$dt_{body}(\omega)$$

$$dt_{Rayleigh}(\omega) = \frac{d\phi(\omega)}{\omega}$$

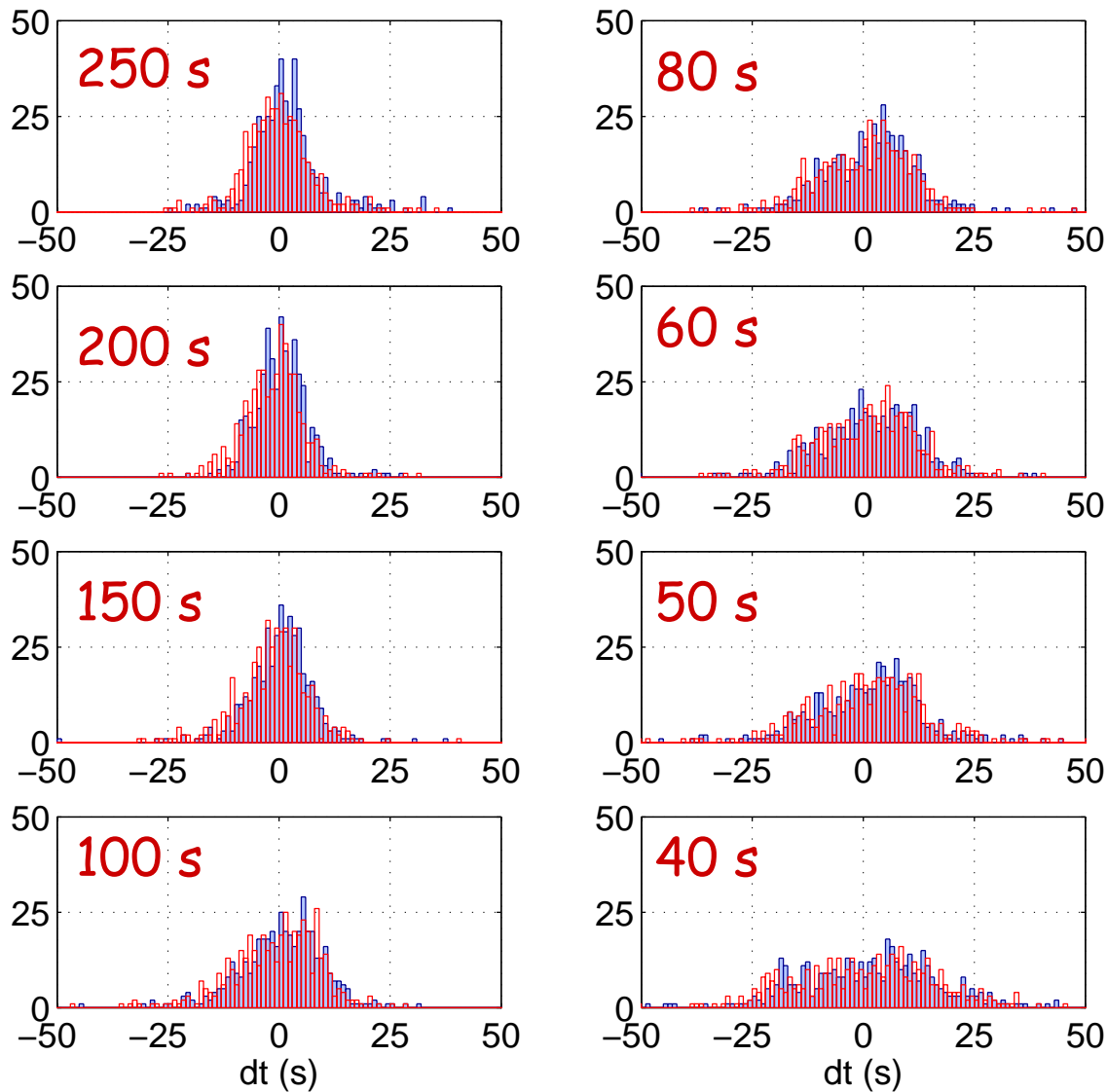
Time-variable filter
to extract the
0th mode Rayleigh waves

Cross-correlation
of the seismograms

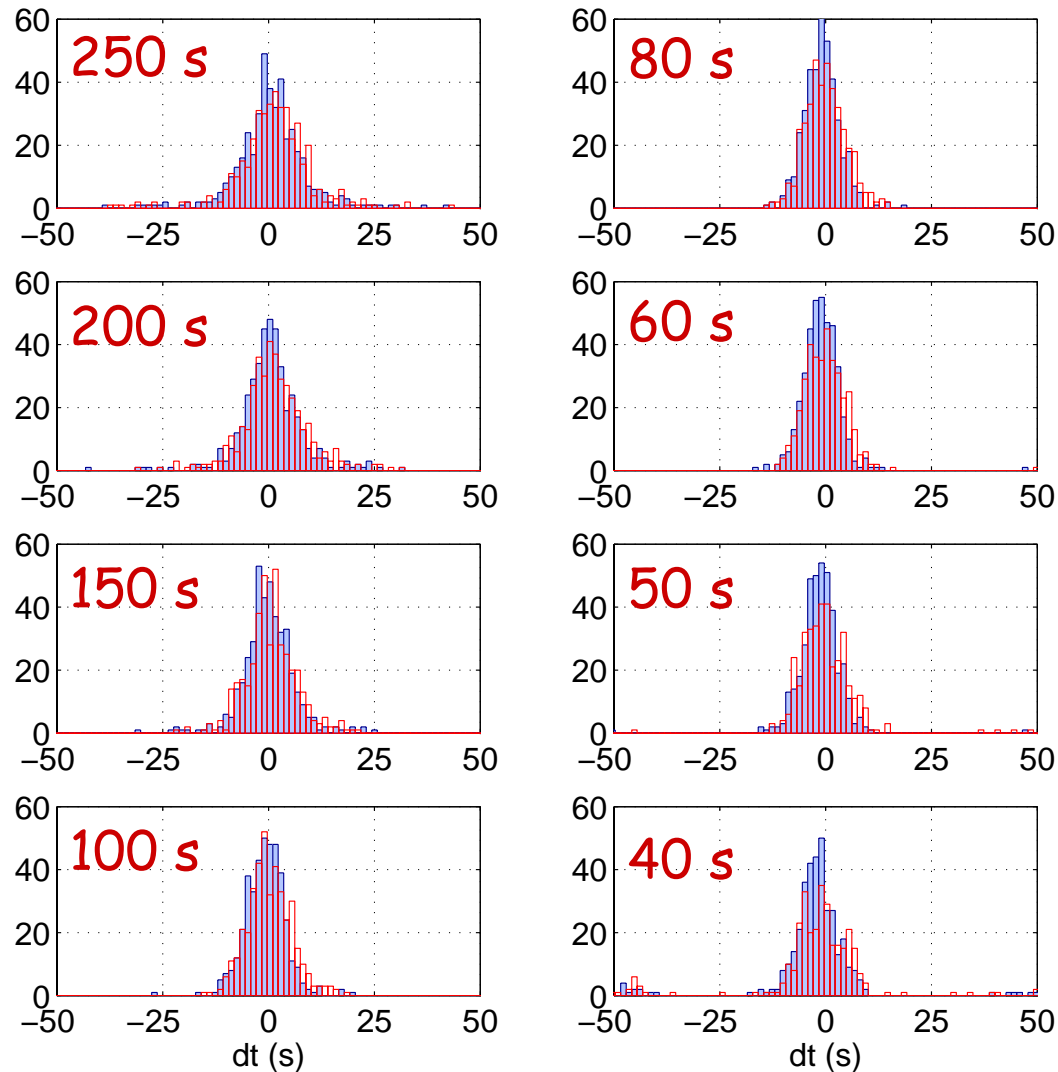
Measure the phase $d\phi$
of the cross-correlation

Unwrap the phase

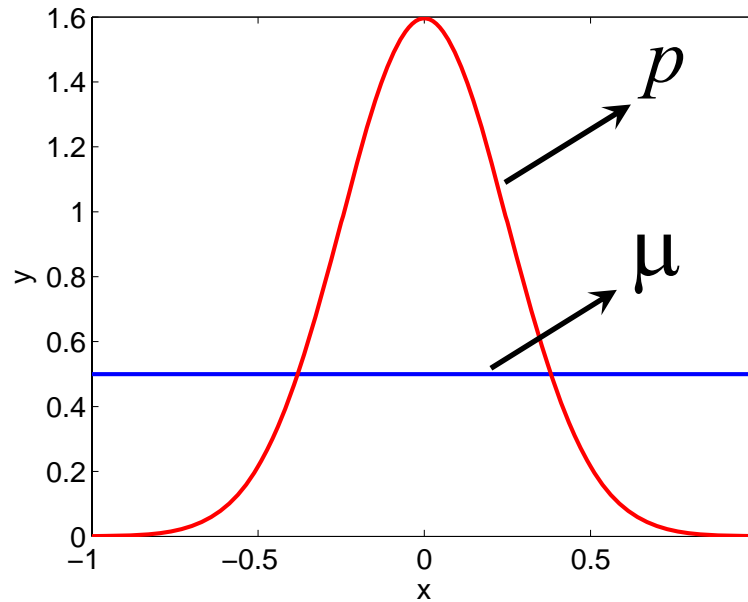
Time shift histograms from Rayleigh waves for the mantle models $\gamma = 0.001$ (blue bars) and $\gamma = 0.00001$ (red bars).



Time shift histograms from **body waves** for the mantle models $\gamma = 0.01$ (blue bars) and $\gamma = 0.00001$ (red bars).



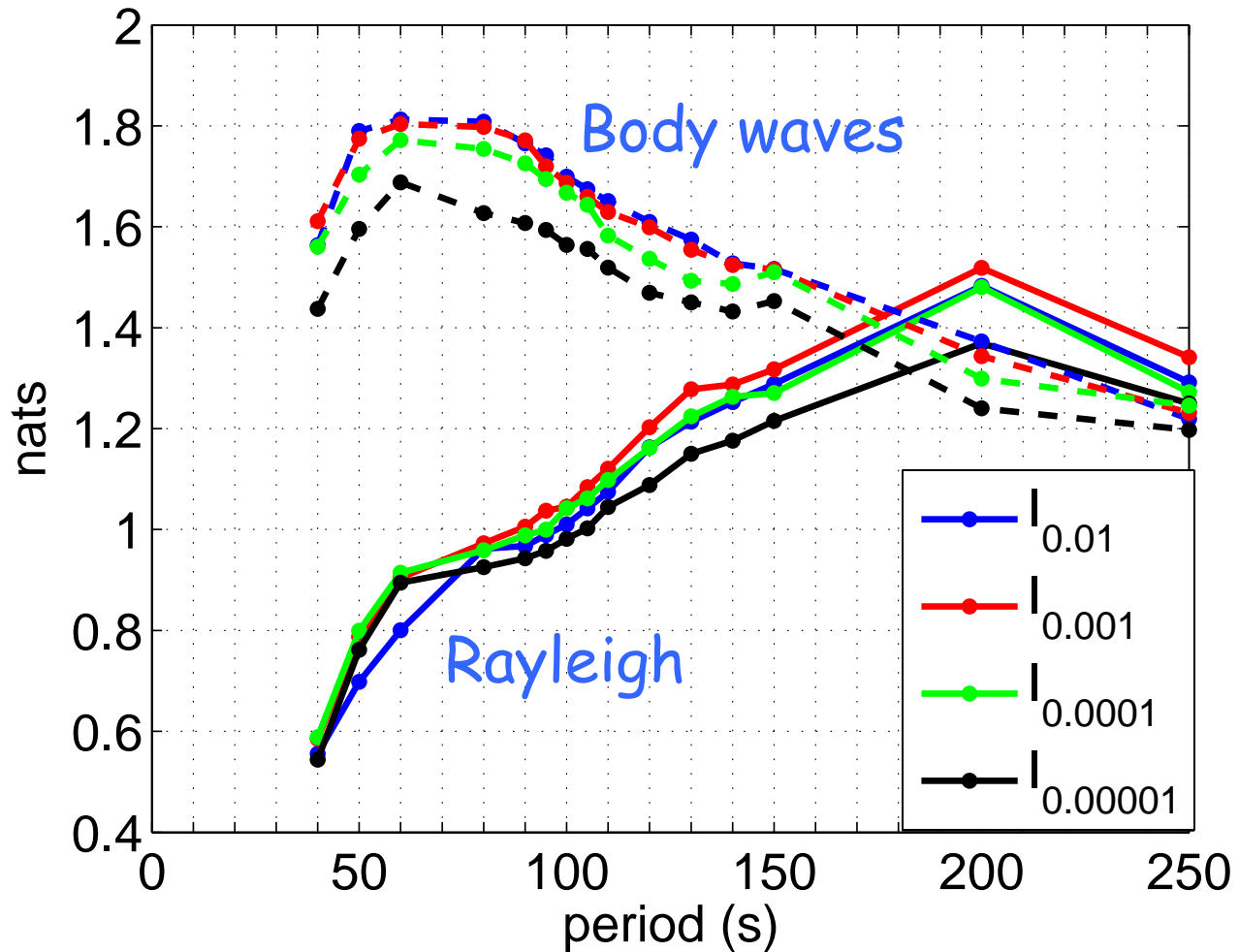
Information content in a probability density function



$$I_i = \int_{-\infty}^{\infty} p_i(m) \log \left(\frac{p_i(m)}{\mu(m)} \right) dm \quad (\text{Shannon 1948})$$

$$I_{\text{gain}} = I_i - I_j$$

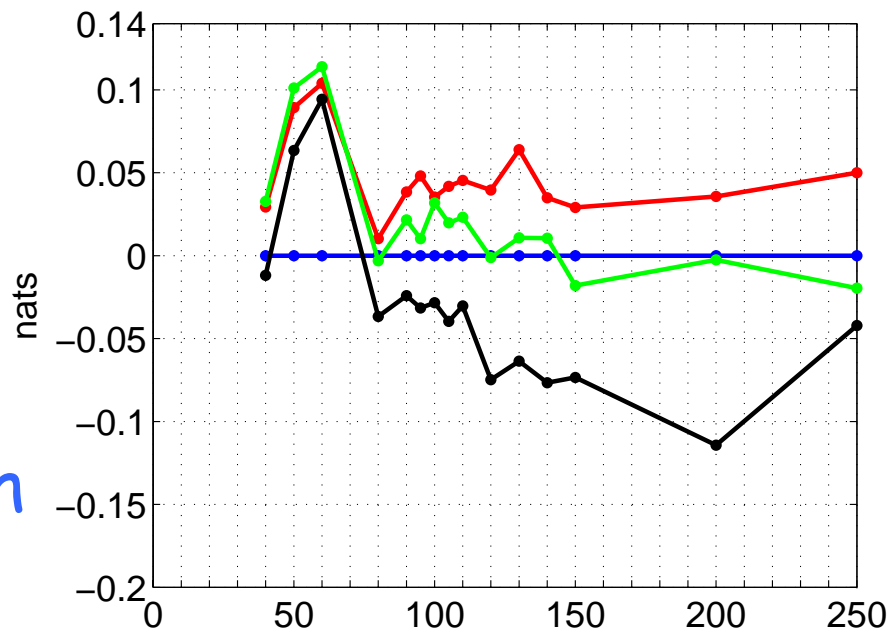
Information content



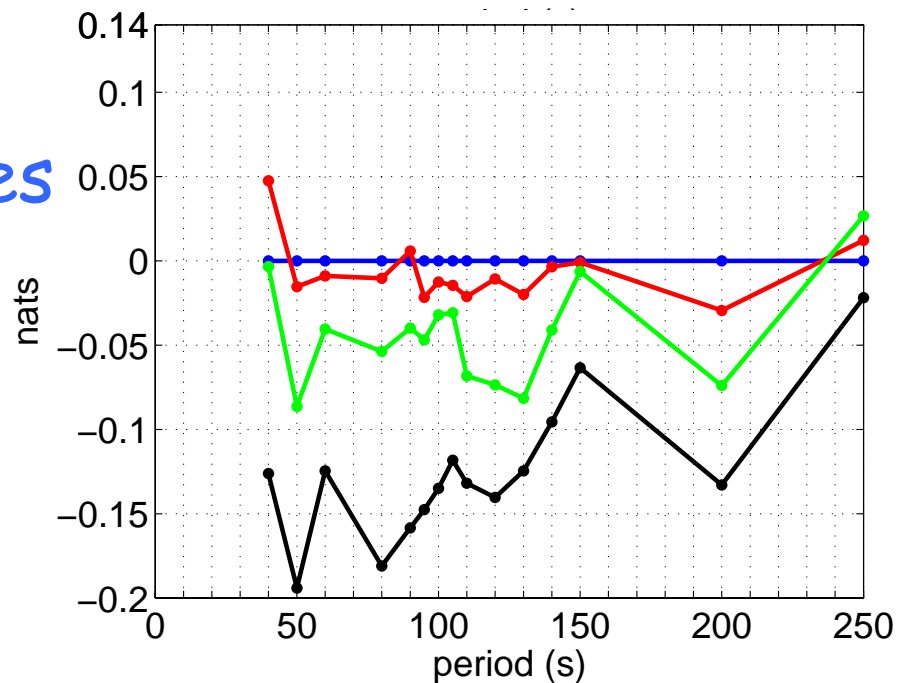






Information gain with respect to smoothest model

Rayleigh



Body waves



-  $I_{0.01} - I_{0.01}$
-  $I_{0.001} - I_{0.01}$
-  $I_{0.0001} - I_{0.01}$
-  $I_{0.00001} - I_{0.01}$

Conclusions

- For Rayleigh waves, the match between real and SEM seismograms is much better at long periods.
- The difference between the models is not large.
- For Rayleigh waves, problems at short periods might be due to the crustal model.
- Increase the number of paths!