



# Seismic wave Propagation and Imaging in Complex media: a European network

**Shane Murphy**  
Early Stage Researcher

**Host Institution:** University of Naples

**Place of Origin:** Dublin, Ireland

**Appointment Time:** January 2006

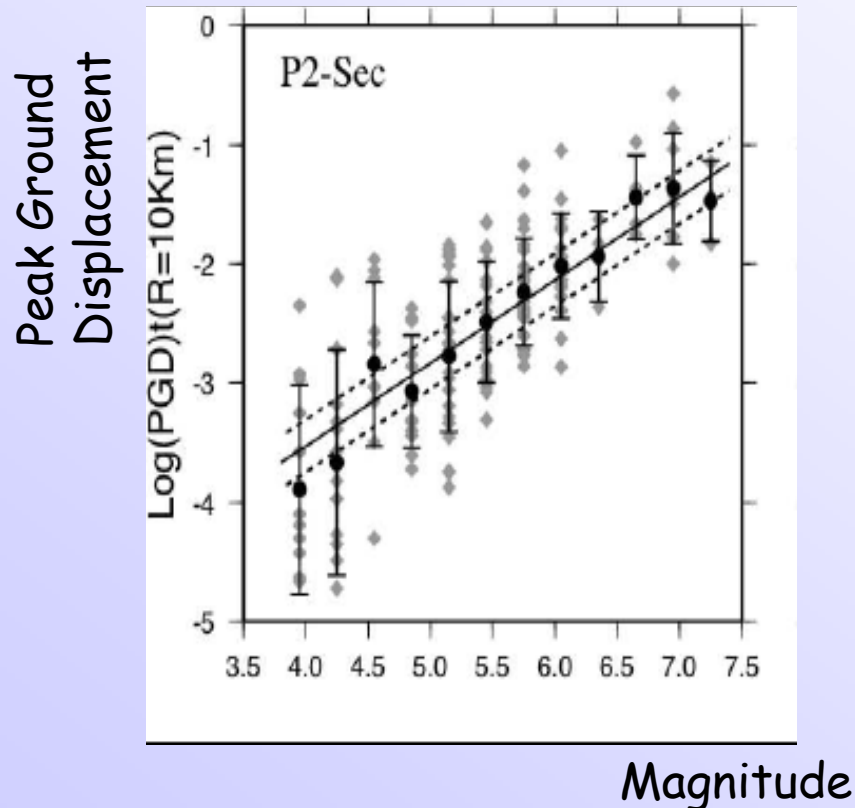
**Project:** How can initial ground motion scale with final rupture size?

**Task Groups:** Local Scale

**Cooperation:** INGV Roma

# Early Magnitude Estimation

*Scaling of Peak Displacement with Magnitude (Wu et al., GRL, 2006 and Zollo et al., GRL, 2006)*



After: Zollo, Lancieri & Nielsen, GRL, 2006

Source: European Strong Motion Dataset (Ambraseys et al., Geol. Teo. Appl., 2004)

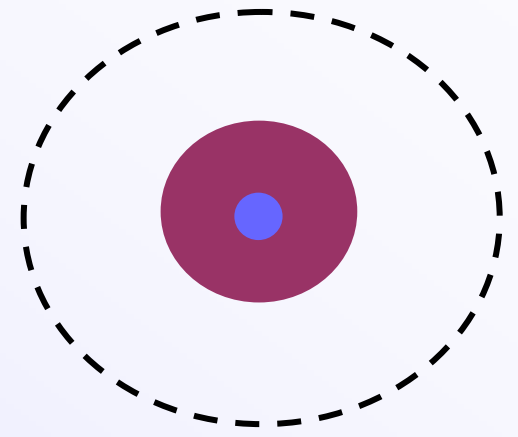
# Explanation

- What physical model could explain this?
  - Breakout phase dynamically determines final earthquake size
  - Kinematic effects due to rupture and wave propagation



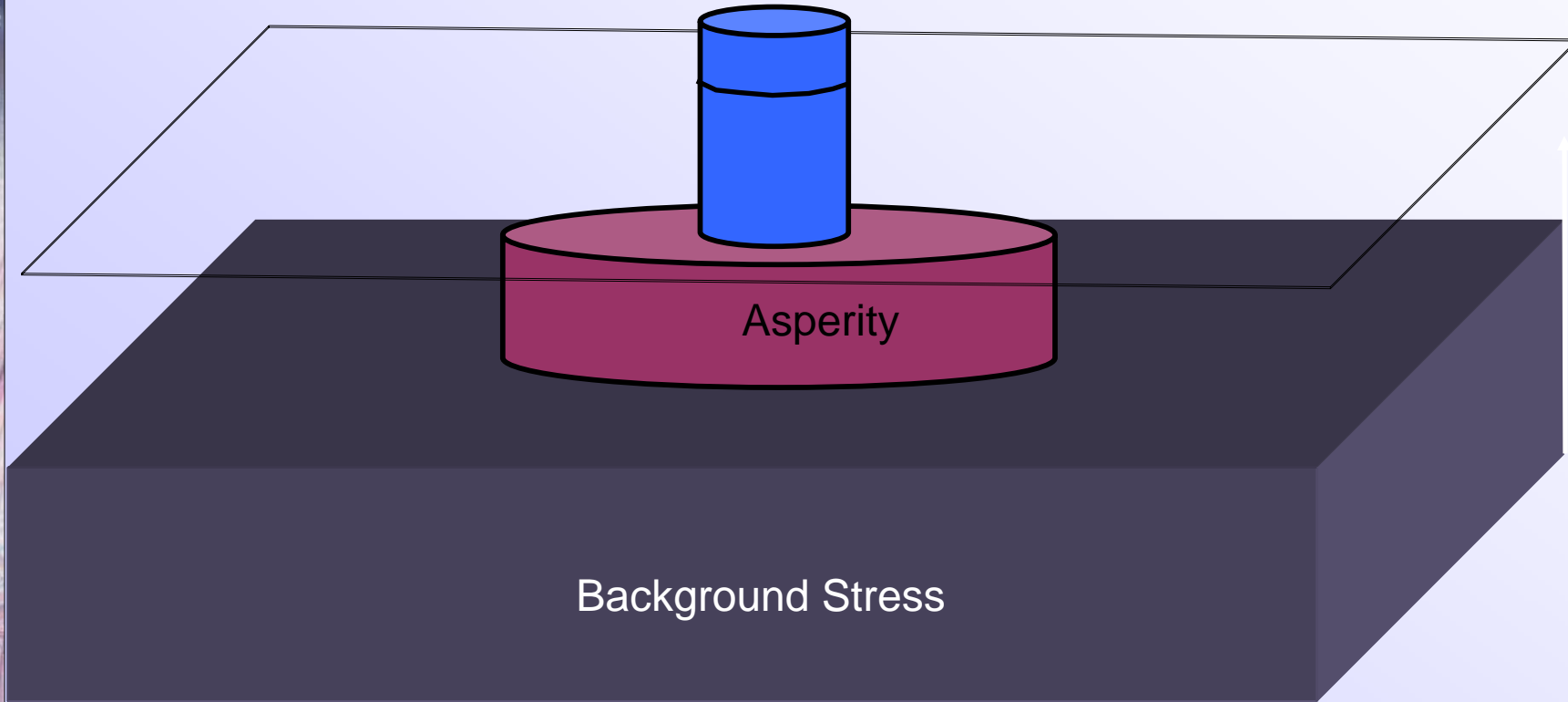
# Dynamic Modeling

# Model Setup



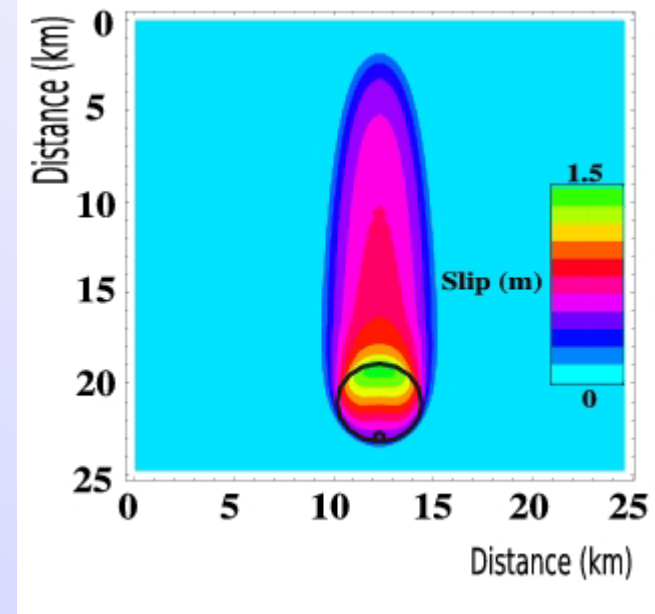
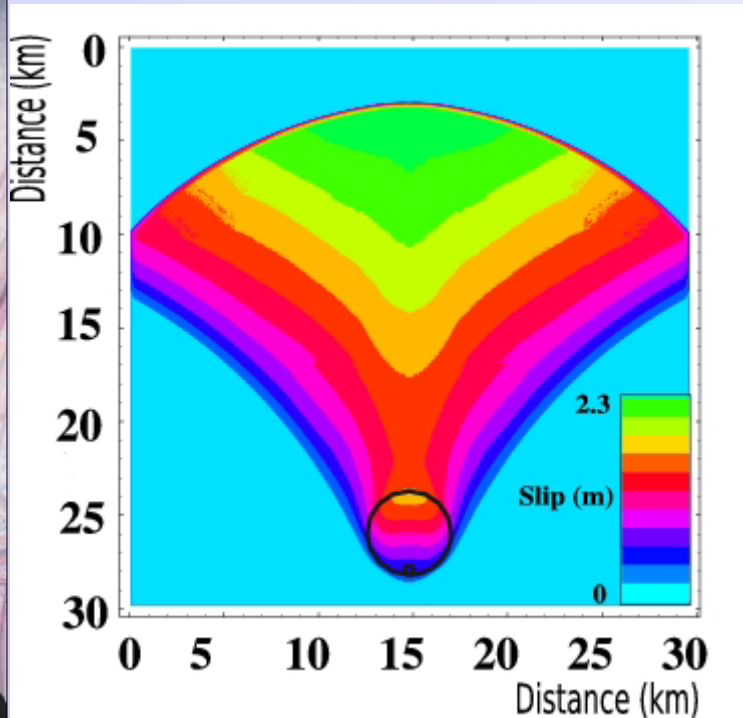
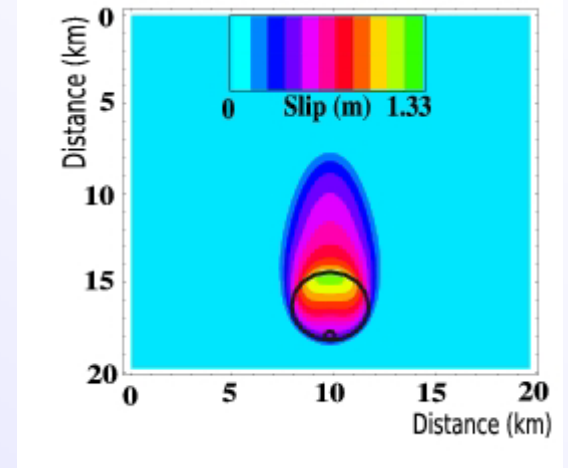
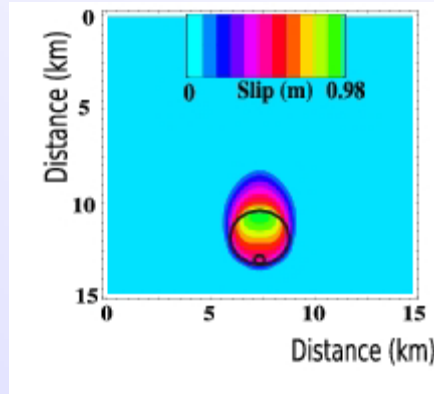
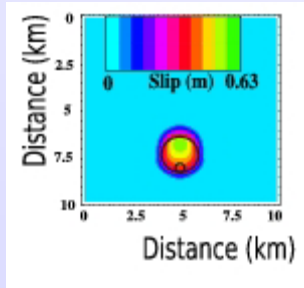
Yield Stress

Triggering patch



# Dynamic Modelling Results -I

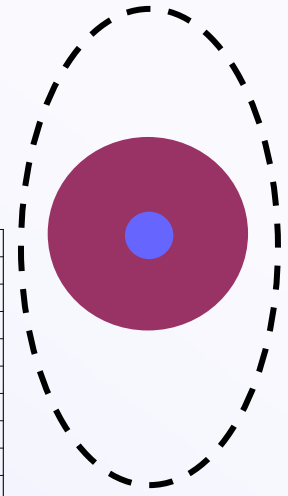
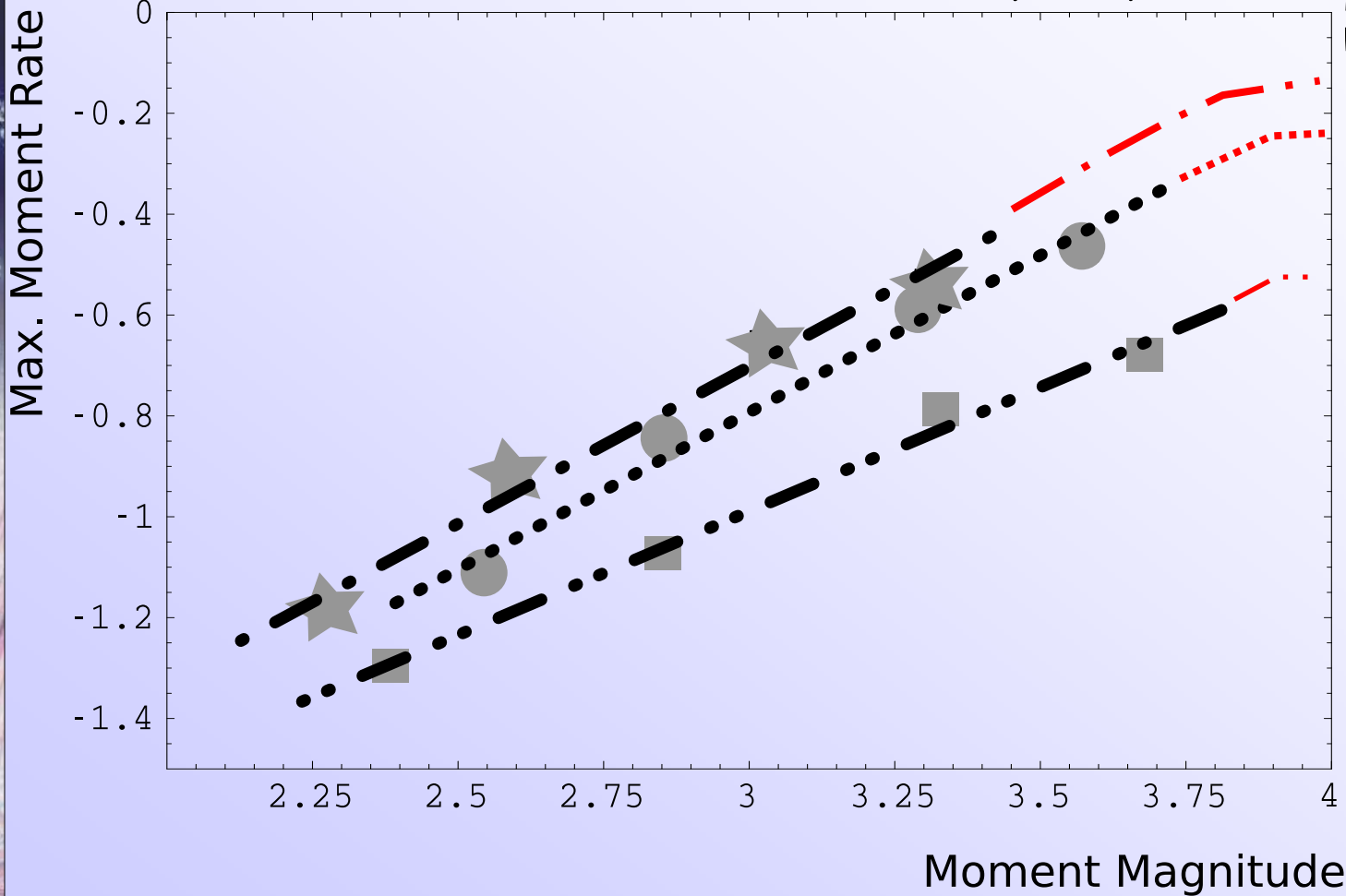
*Total slip on the fault*





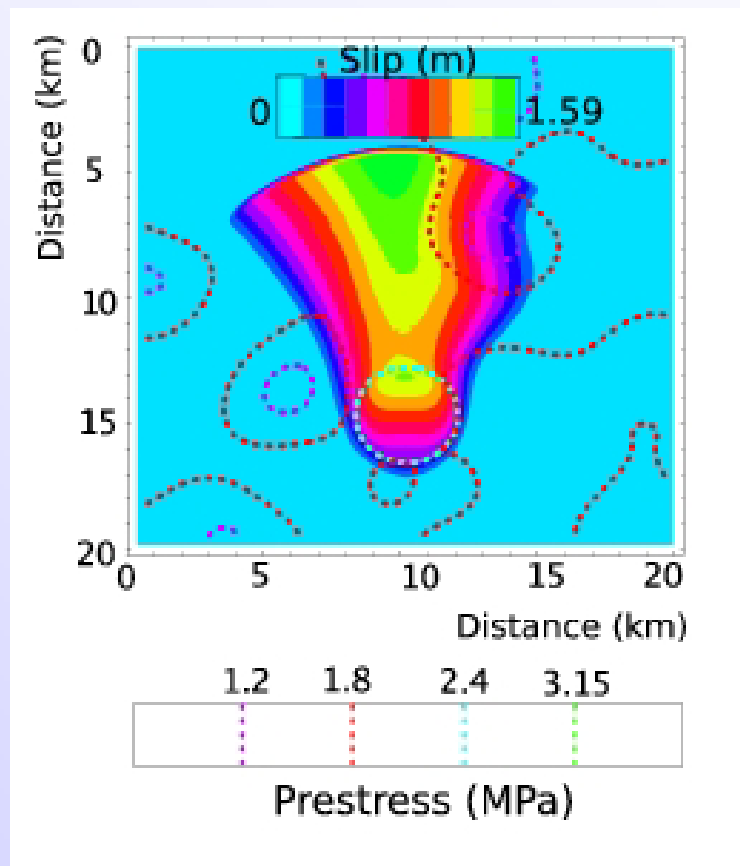
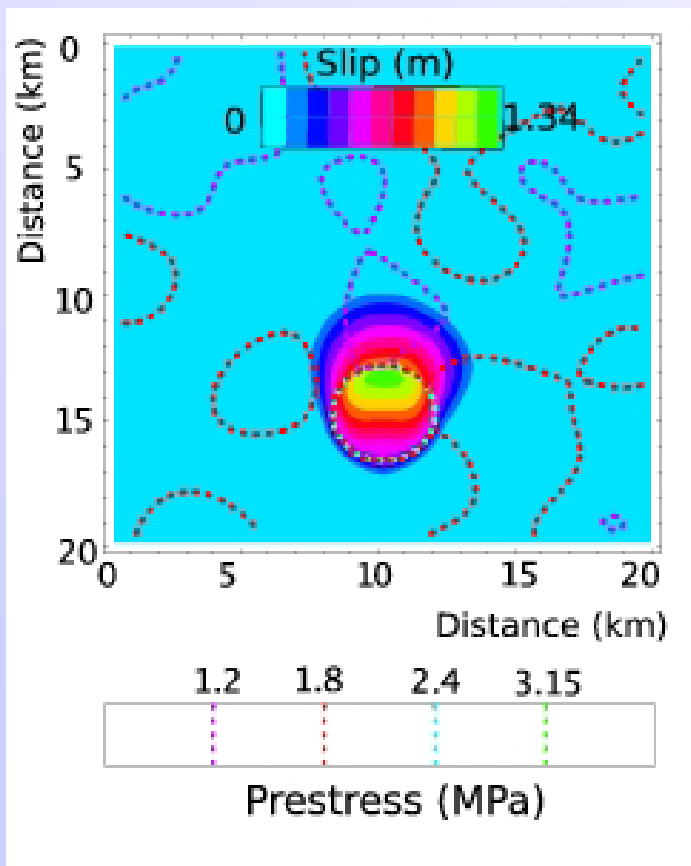
# Dynamic Results -III

*Max. Moment Rate Vs ML Outside of asperity*



# Cascade Feature of Rupture

*Application of heterogeneous pre-stress*





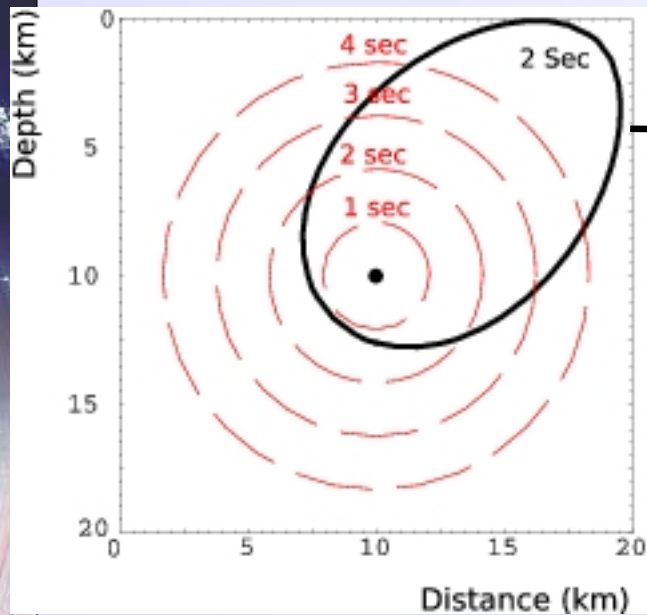
# Dynamic Conclusions

- *Dynamic Rupture is Deterministic over a very short scale*
- *The stopping of earthquakes is predominantly a cascade in nature*
- *Scaling only works for a particular class of rupture models (self healing pulses)*

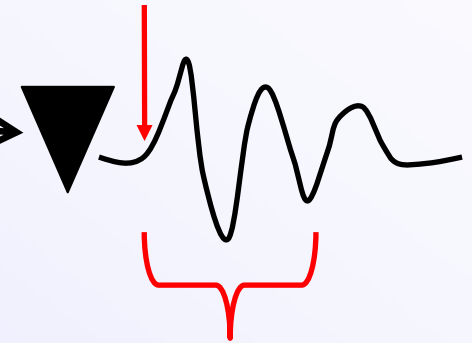


# Kinematic Modeling

# Isochrone Back Projection



S-Wave Arrival



2 seconds

$$u(x,y,z,t) = \Delta u \frac{\mu R^c}{4\pi \rho \alpha^3 r} \int_S dS$$

Slip  $\propto$  Length of Fault

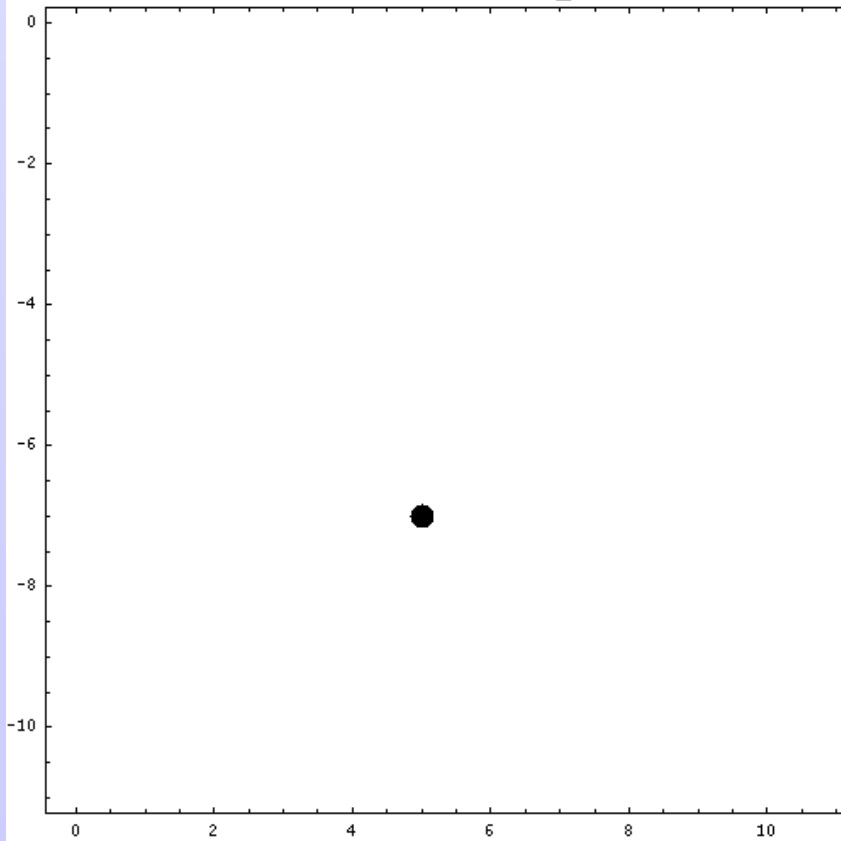
After: Bernard & Madariaga, 1984

# Isochrone back-projection

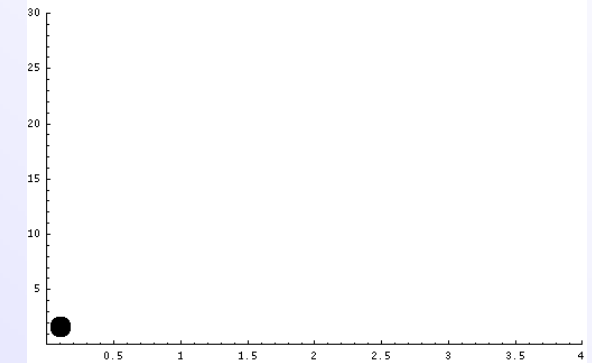
*Station Distance: 30km*  
*Rupture Velocity: 2.5km/s*  
*Hypocentre Depth: 7km*

*P-wave Velocity: 4km/s*  
*Azimuthal Angle: 45 degrees*  
*Fault Size: 10km x 10km => 5-6.5M<sub>L</sub>*

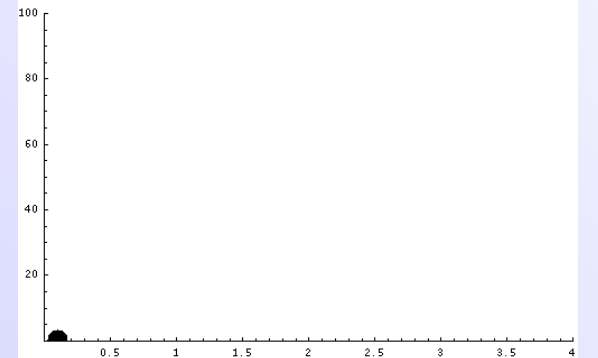
## Isochrone Shape



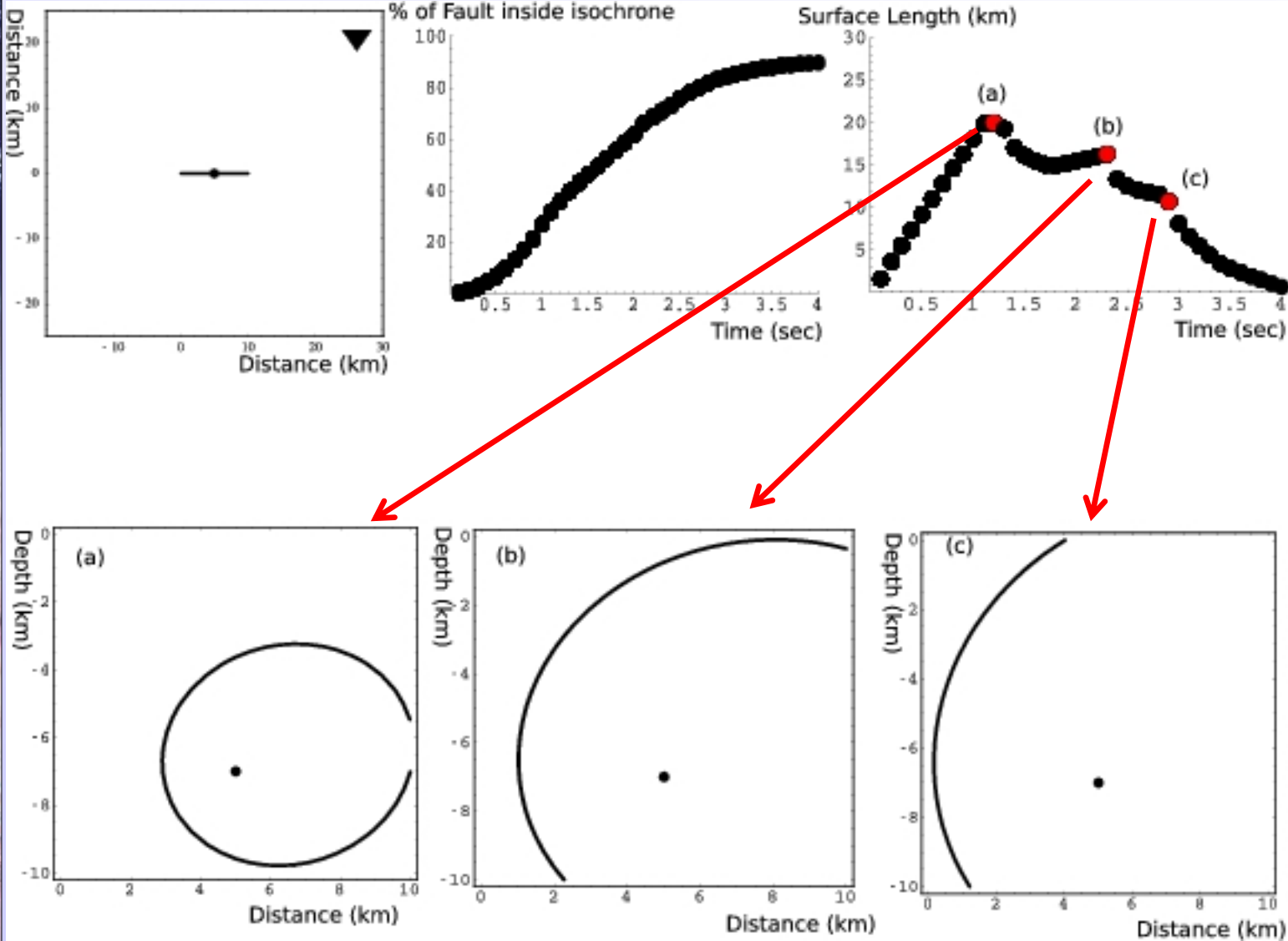
## Isochrone Length



## % of Total Fault Area

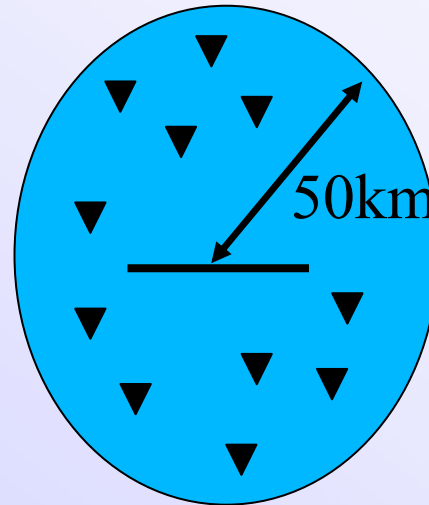


# Isochrone evolution over time



# Model Set up

*Station Distribution: Random Distribution*



*Two velocity Profiles: (a)  $v_p=4.5\text{km/s}$       (b)  $v_s= 2.6\text{km/s}$*

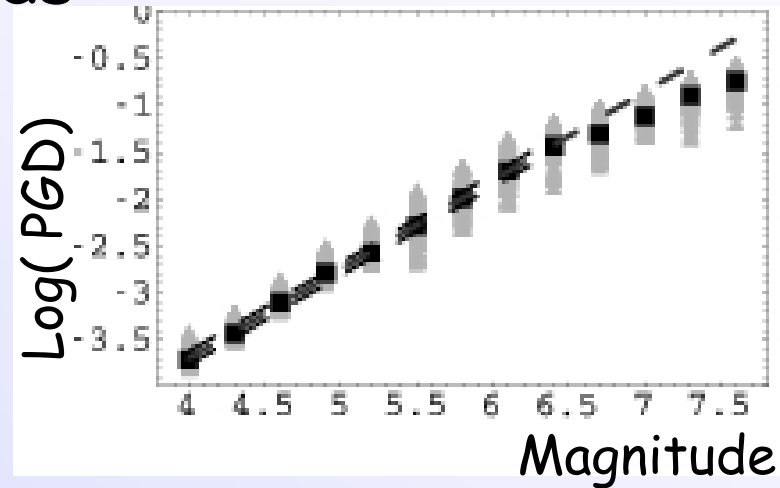
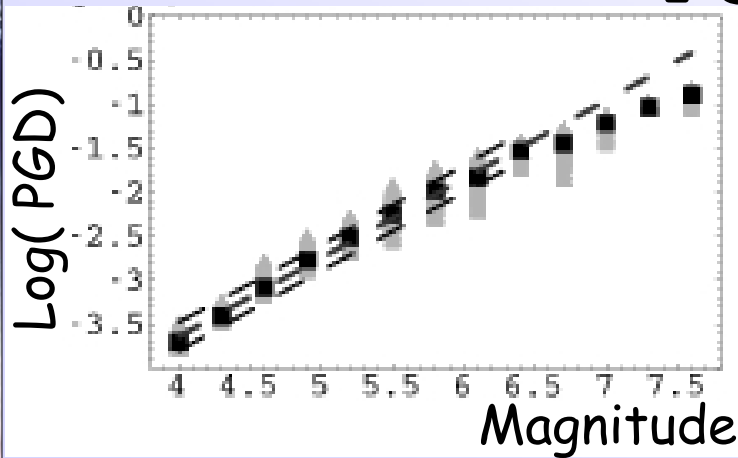
*Hypocentre Location: Randomly distributed below 10km*

*Fault Length for Square Fault:    0.94km - 59.5km*

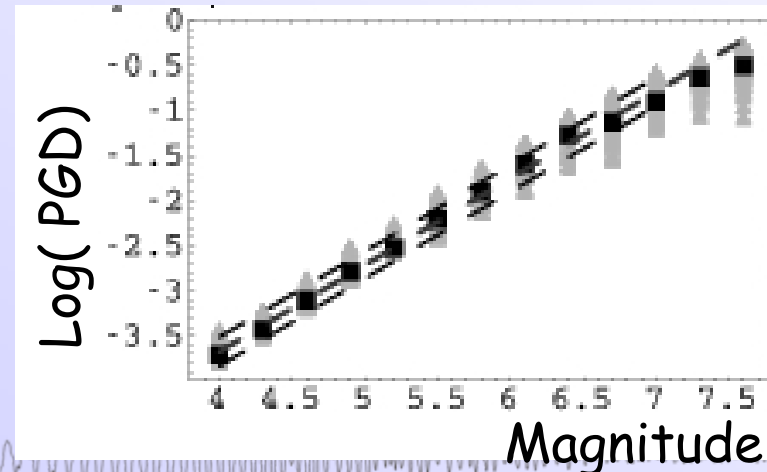
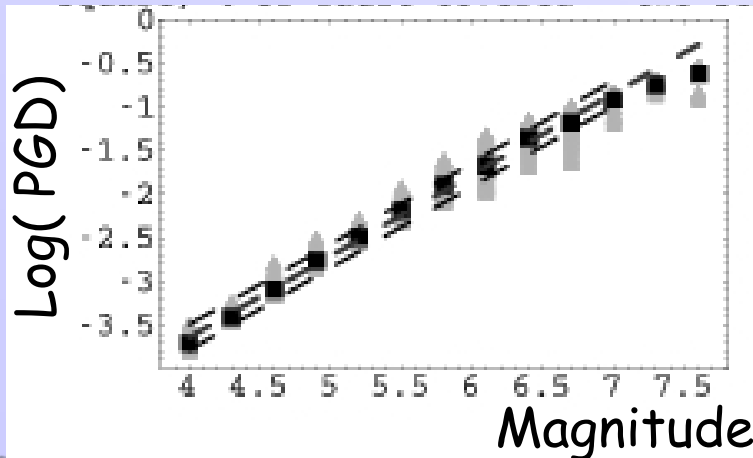
# Results for Simplistic Peak Ground Motion

P-wave: 4.5 km/s      S-wave: 2.6 km/s

1 Seconds



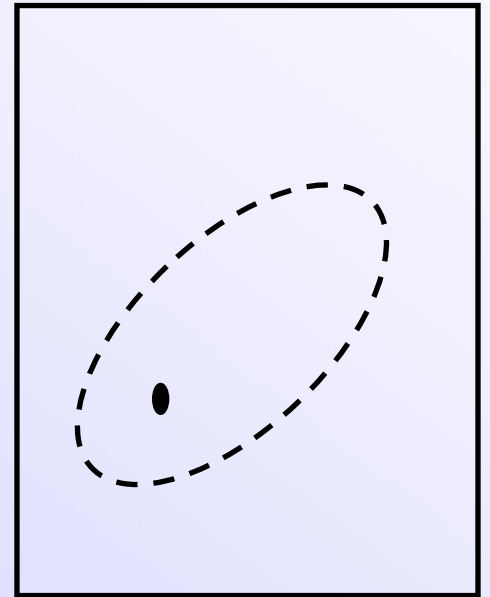
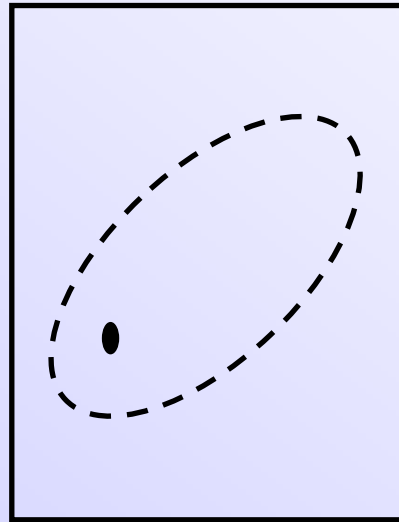
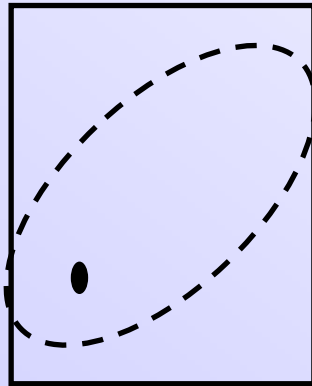
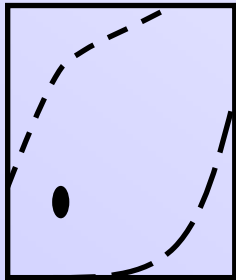
2 Seconds



# Why does the isochrone length become constant?

- *Isochrone = 2 secs in all cases*

*Increasing Magnitude*



*Maximum Isochrone Length  
Controlled by interaction with  
Fault sides*

*Isochrones are no longer  
touching fault sides at 2 sec  
period*



# Kinematic Conclusion

- Earthquake magnitude estimation is highly dependant on fault kinematics
- Measuring Isochrone length produces similar features seen in peak displacement for early magnitude estimations



# Final Conclusion

- Kinematic rather than dynamic would appear to influence early magnitude estimation



# The End