



TIME REVERSAL and SEISMIC SOURCE IMAGING

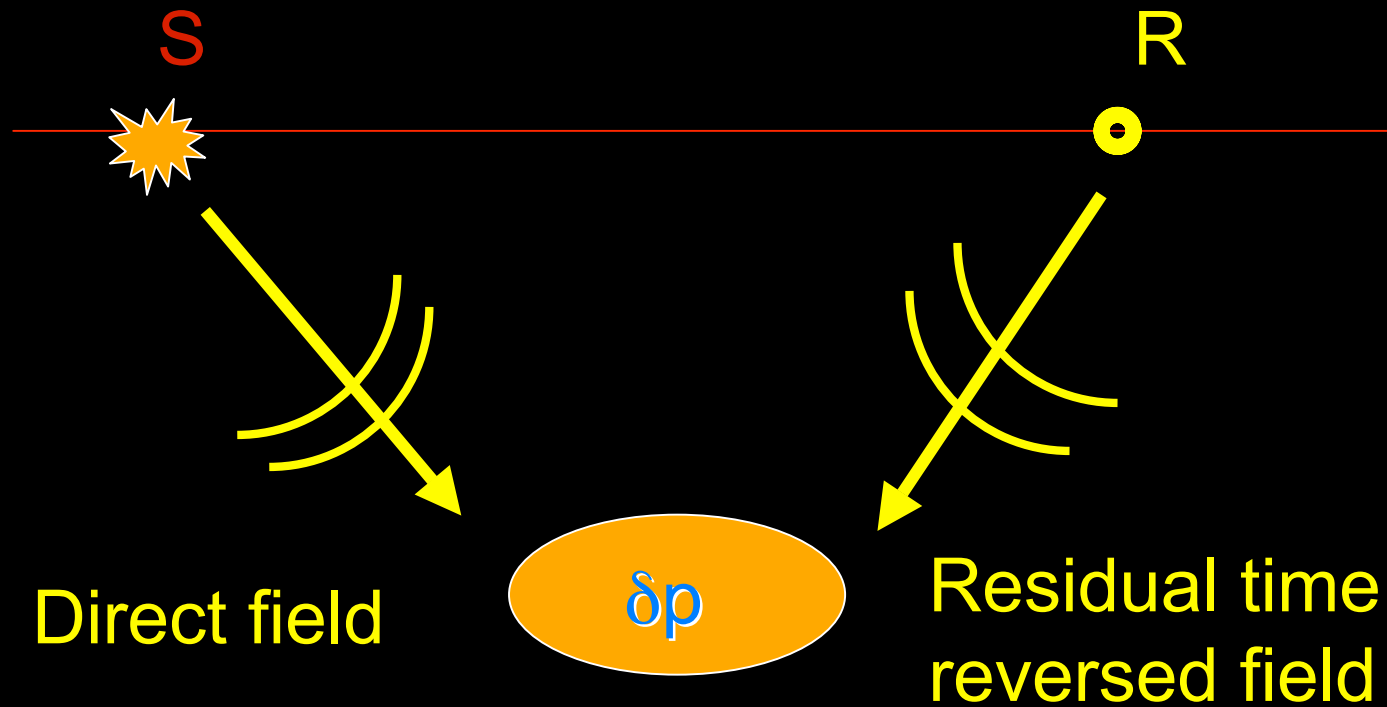
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Yann Capdeville, Eric Clévéde, Huong Nguyen,....
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LOA, ESPCI, Paris

Time Reversal- Adjoint Tomography:

-Source focusing (Green function known)

-Adjoint Tomography \Rightarrow Structure
(source known)



Background

- Time reversal experiments (Fink's group) => focusing

- Problem of Green function?

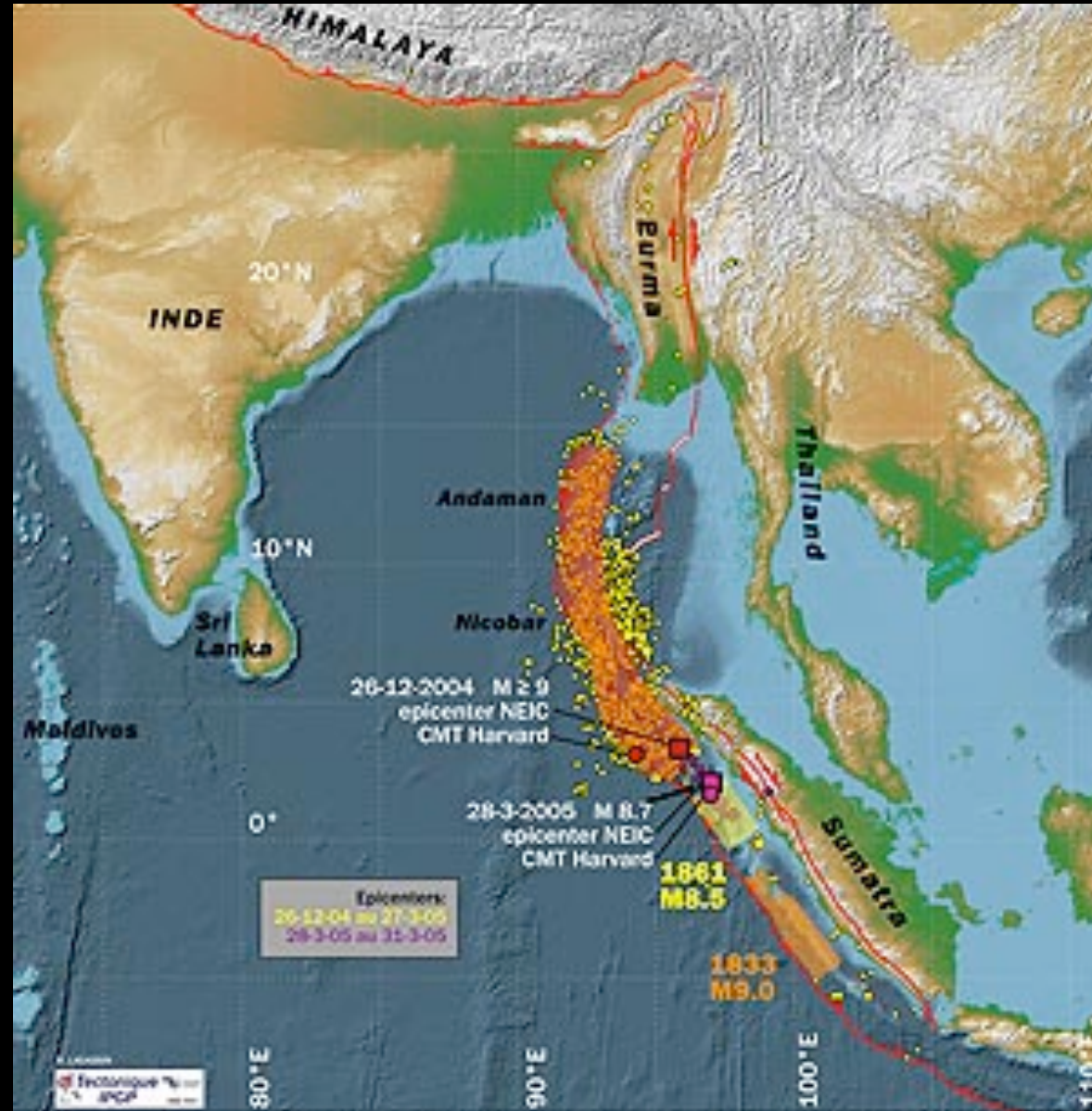
- Seismology: development of numerical techniques

SEM-CSEM (Komatitsch, Tromp, Vilotte, Capdeville,...)

Accurate Green functions?

- Motivation: Big Sumatra-Andaman earthquake

Rupture length 1200km



Time reversal Concept

Elasto-dynamics equation, for seismic displacement field $u(r,t)$

$$\partial^2 u / \partial t^2 = H \cdot u$$

In the absence of attenuation, rotation, time invariance and spatial reciprocity

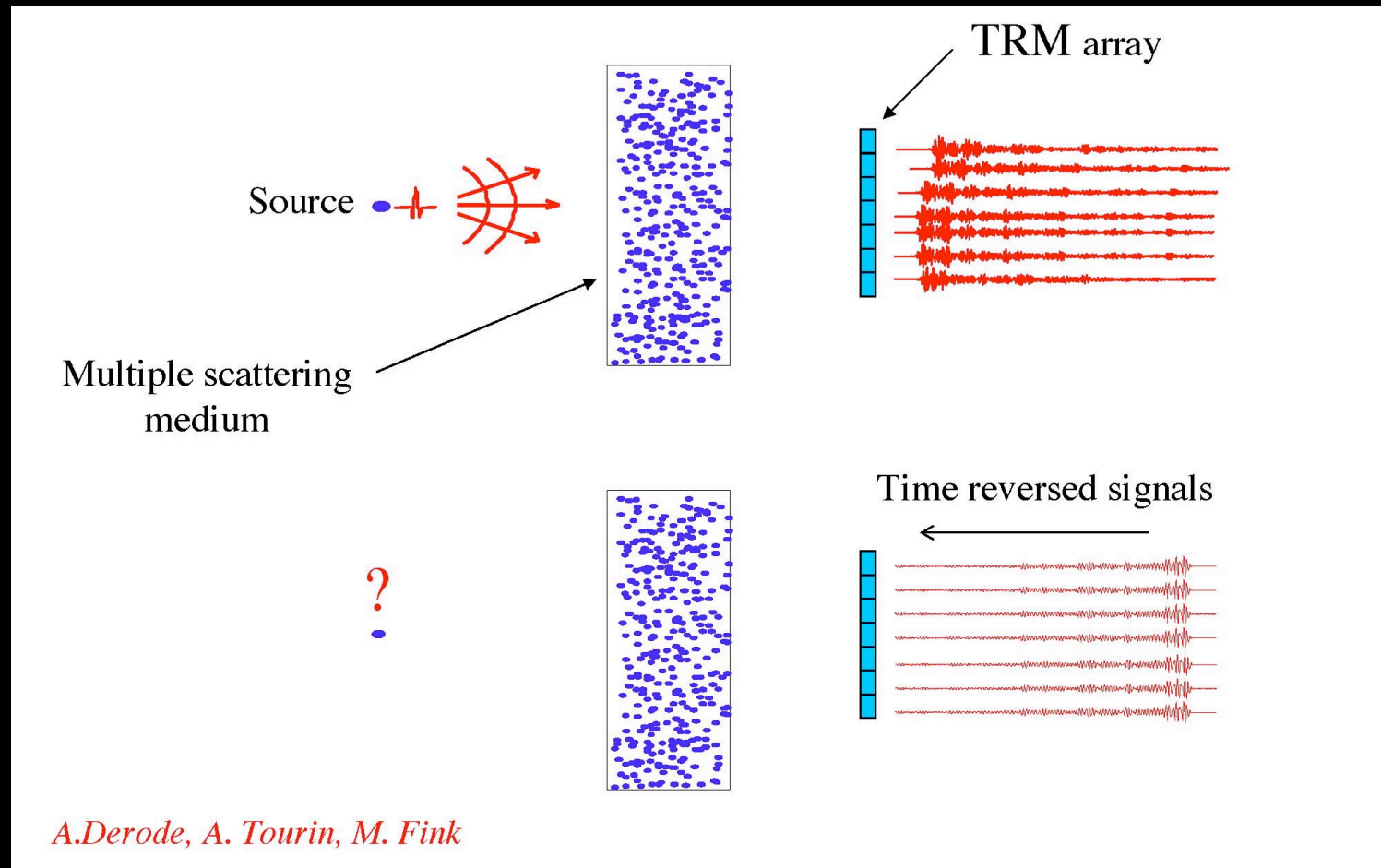
if $u(t)$ is a solution, $u(-t)$ is also a solution.

**We can send back waves with reversed time:
how to get a good focusing?**

Basic Principle for acoustics:

■ Acoustic Source -> receivers

■ Existence of transducers being at the same time recorders and emitters



Refocusing at the source location by sending back signal ($-t$) through the SAME medium from a small number of emitters



Seismic Source Imaging by time reversal

Method Principle:

- Acoustic Source -> receivers
- Existence of transducers at the same time recorders and emitters sending back signal in the same medium

How to apply this concept to seismic waves within the Earth?

1C (scalar) ->3C (elastic case)?

Limited number of receivers?

Realistic Propagating Medium? 1D-3D Earth models
(Green functions)

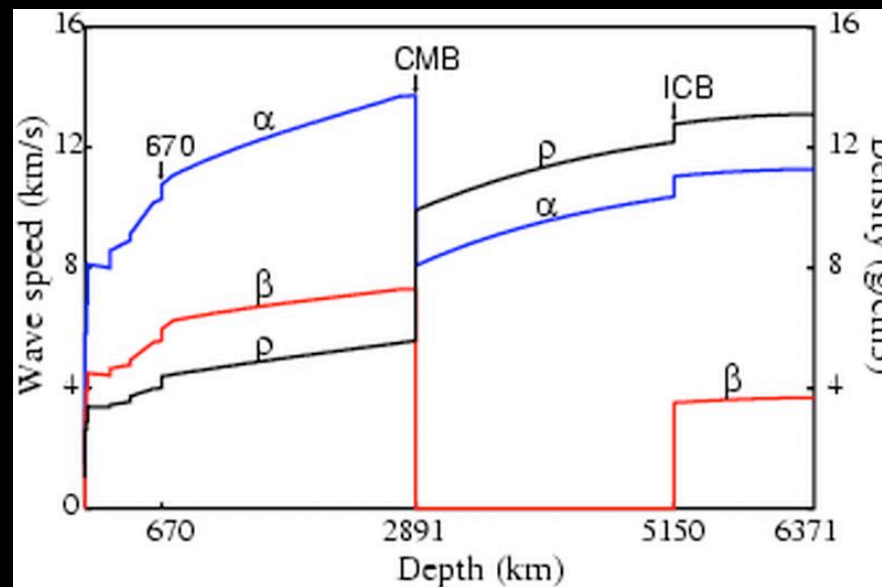
Time reversal

Seismic displacement field $u(r,t)$ can be calculated everywhere by the Normal modes or CSEM- method

It is possible to backpropagate $u(-t)$

- Vertical component
- Very long periods $T > 150s$

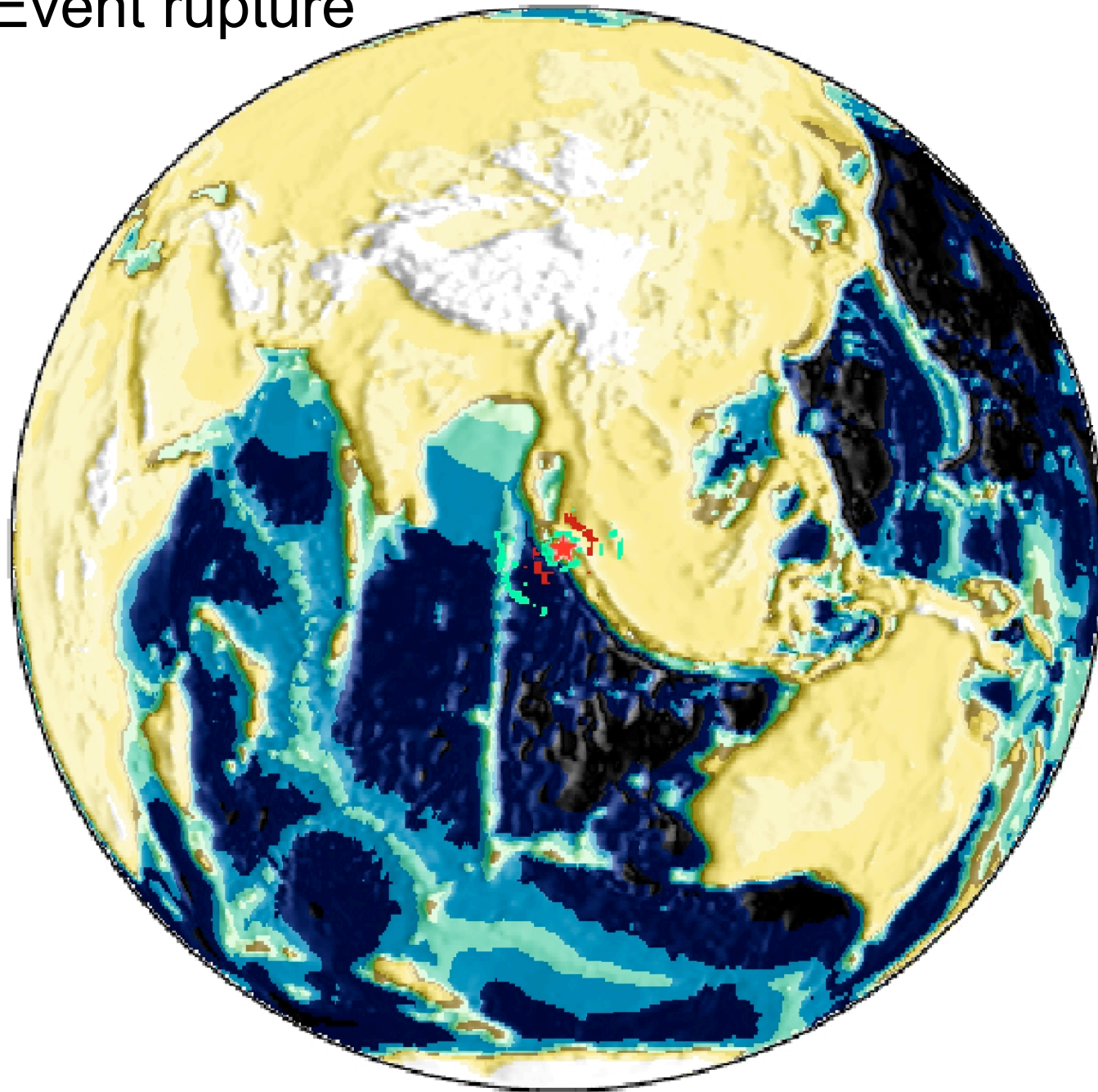
• 1D PREM



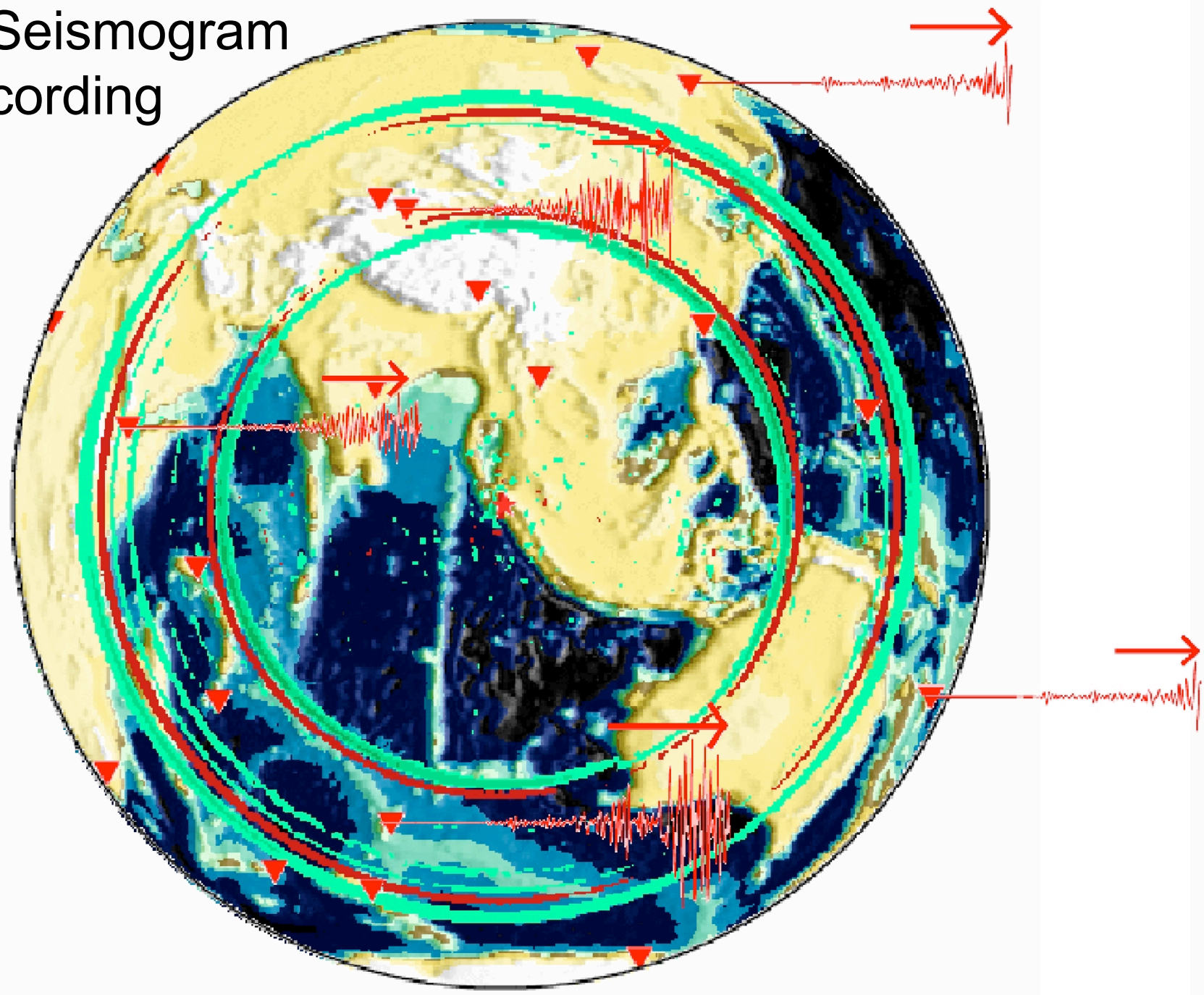
3D Model

Larmat et al., 2006

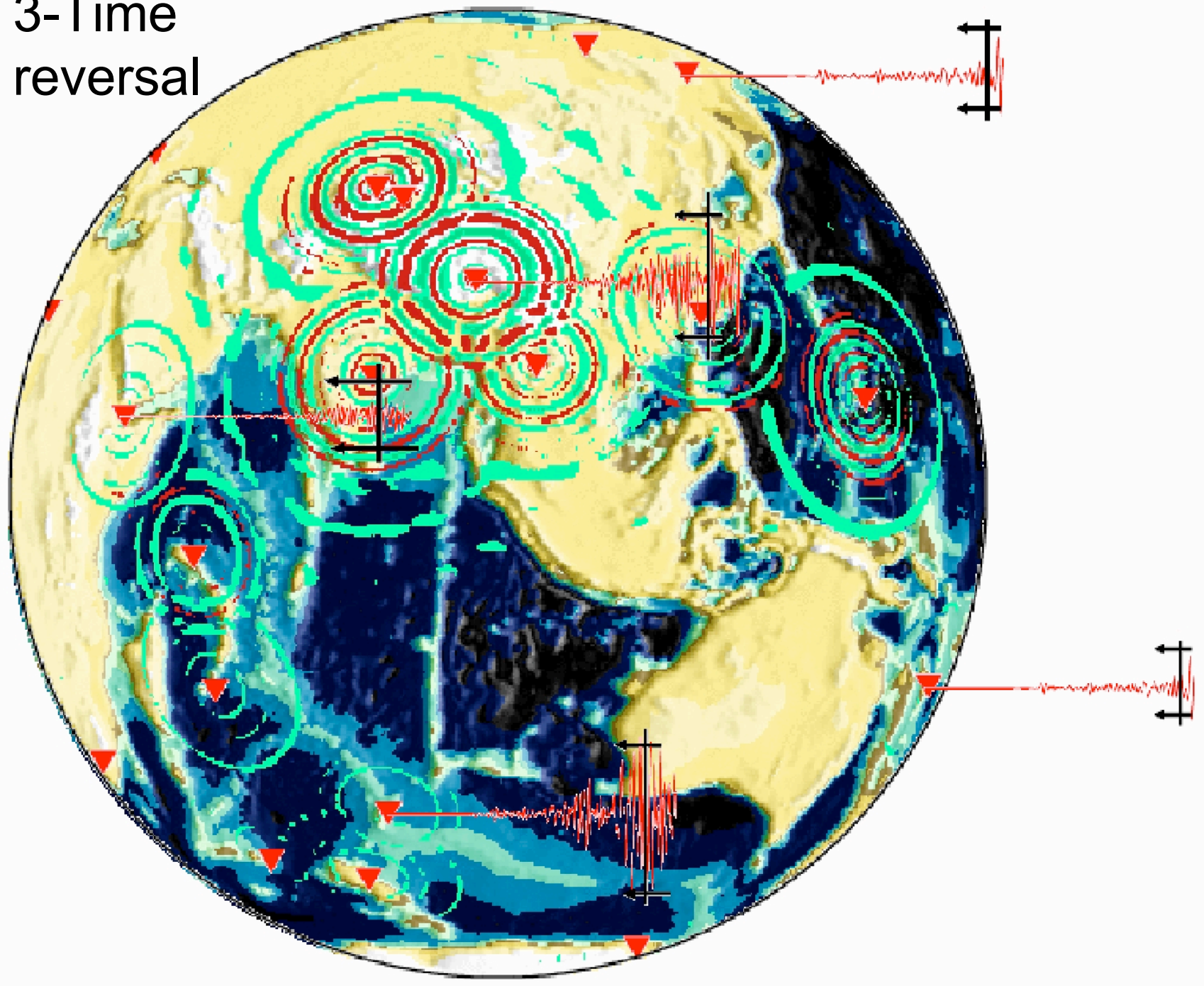
1-Event rupture



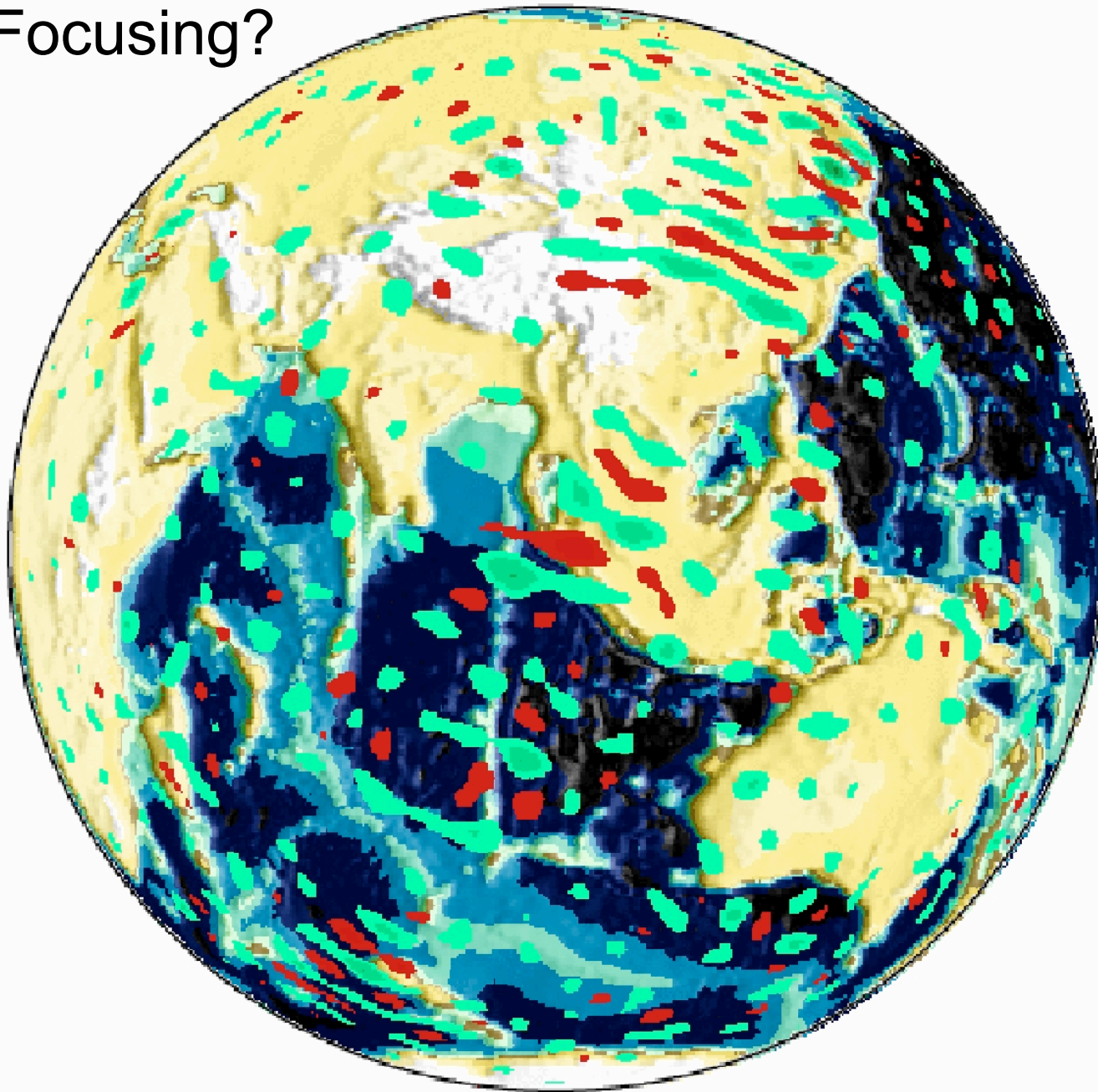
2-Seismogram recording



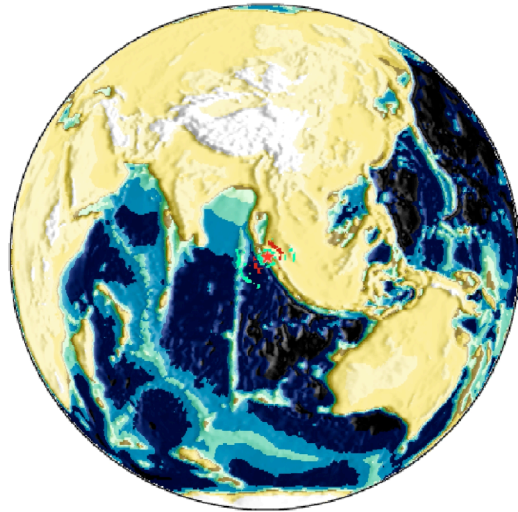
3-Time reversal



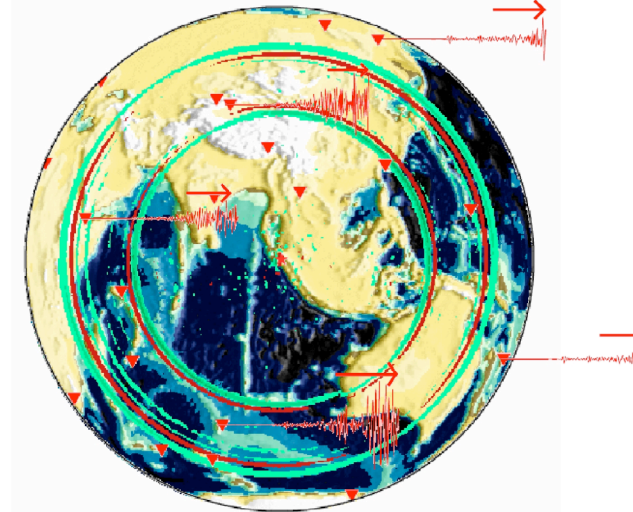
4- Focusing?



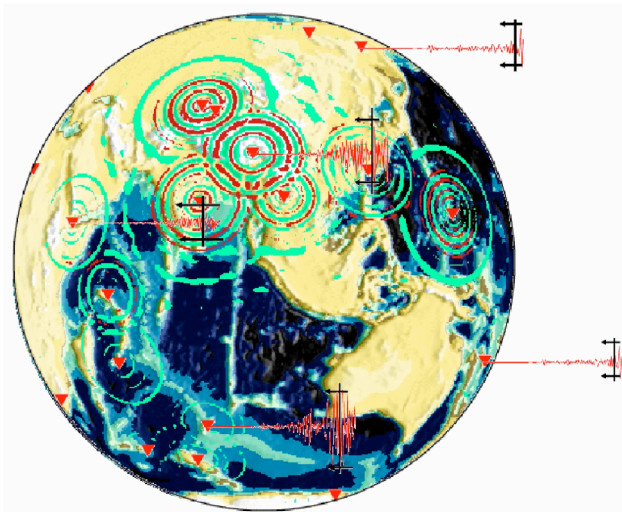
1- Event rupture



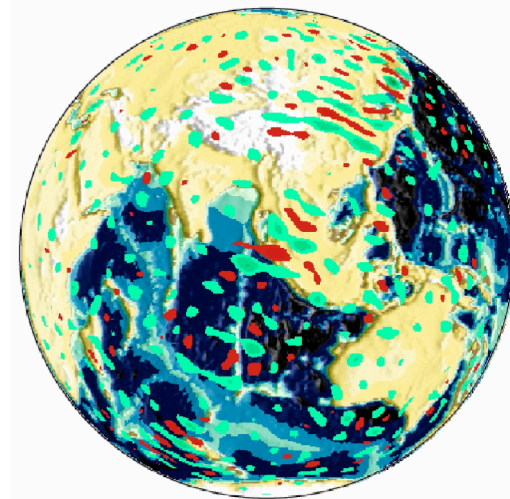
2- Seismogram recording



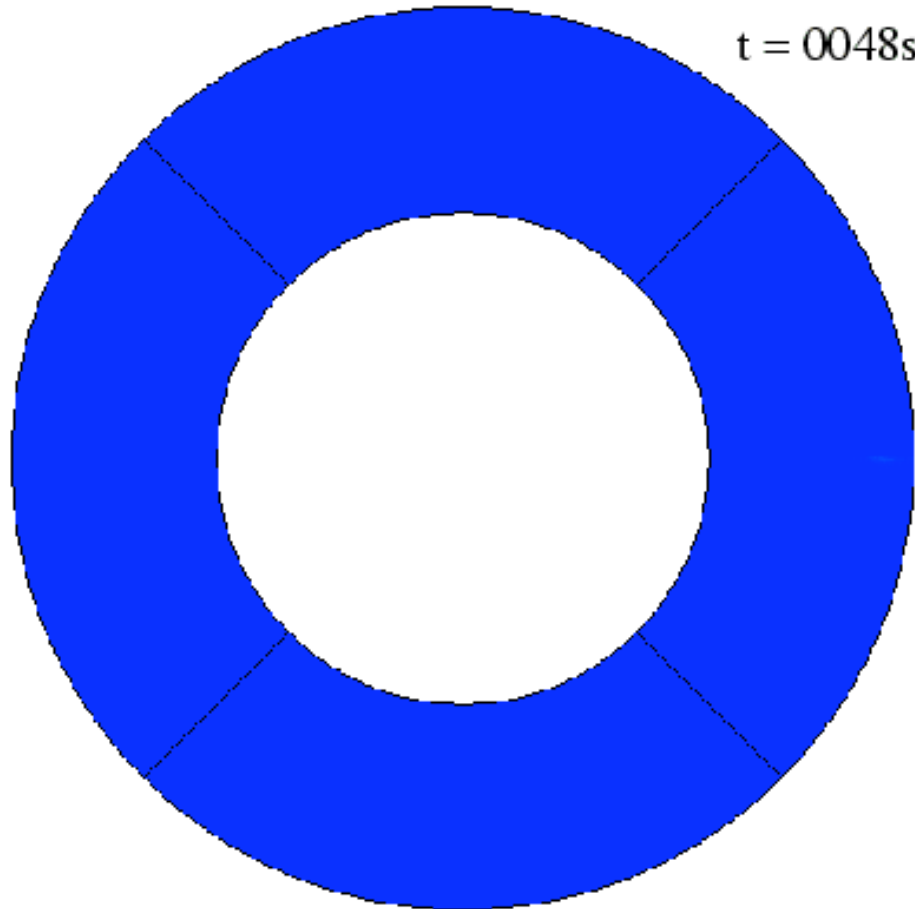
3- Time reversal experiment



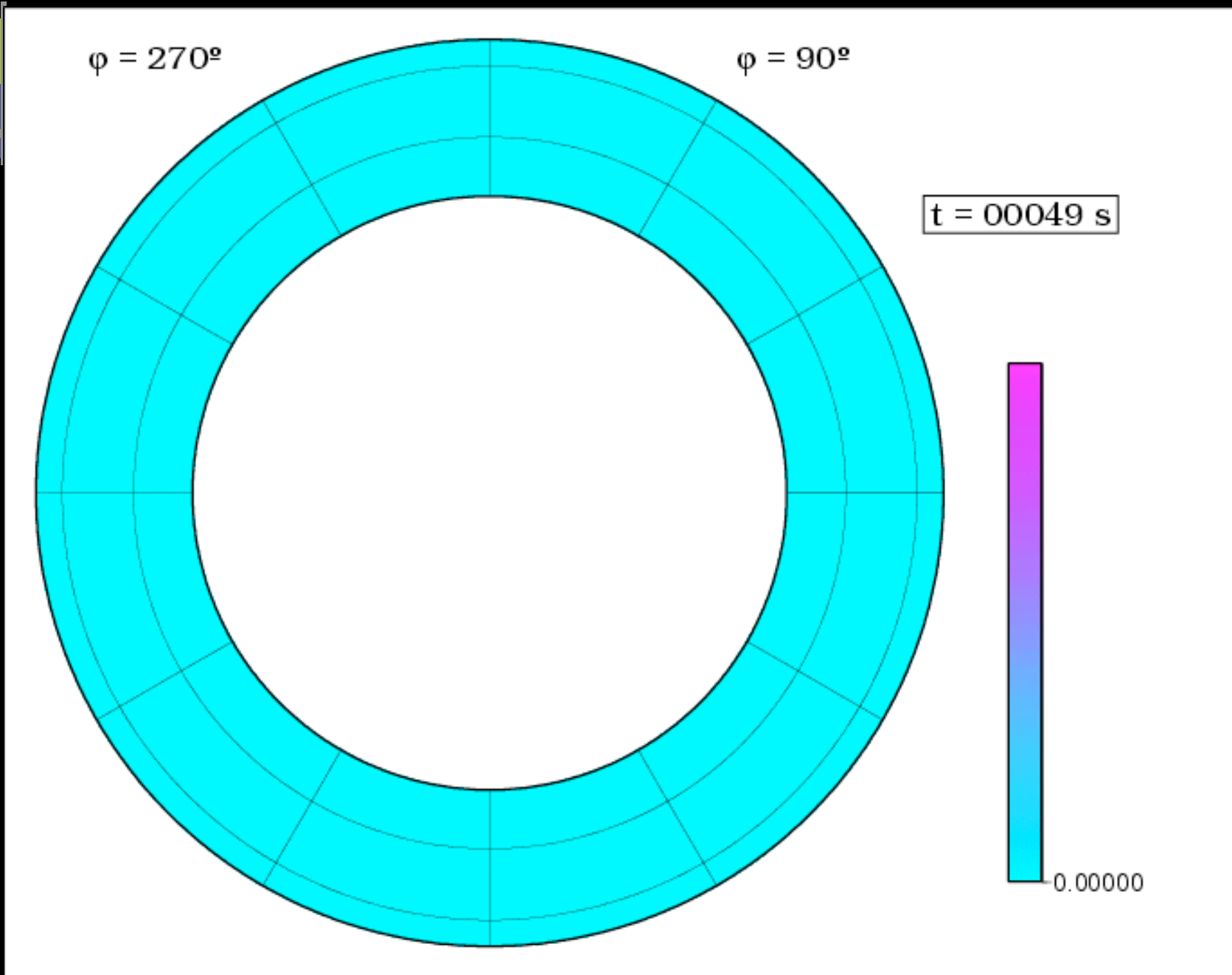
4- Focusing



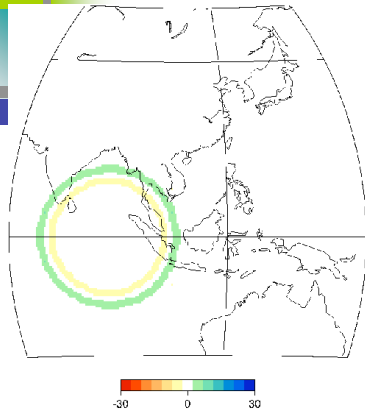
Synthetic Test: Point Source: Forward problem



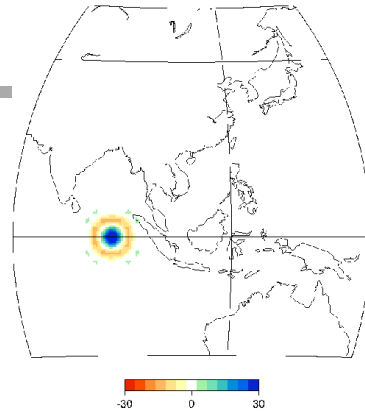
Time reversal



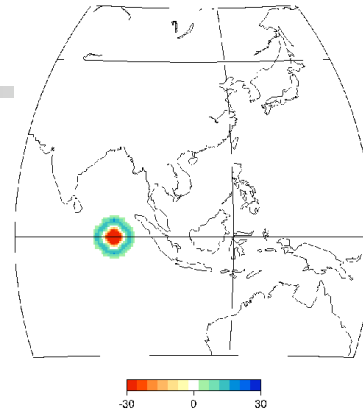
Homogeneous azimuthal distribution



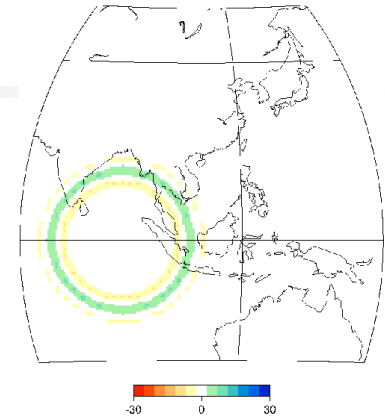
$t = -500s$



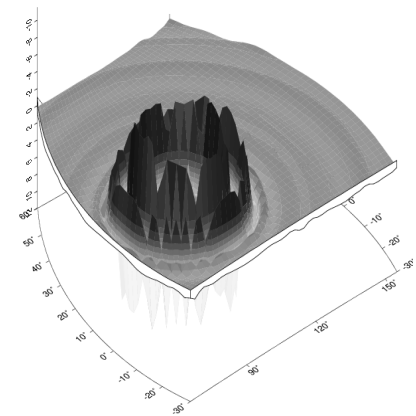
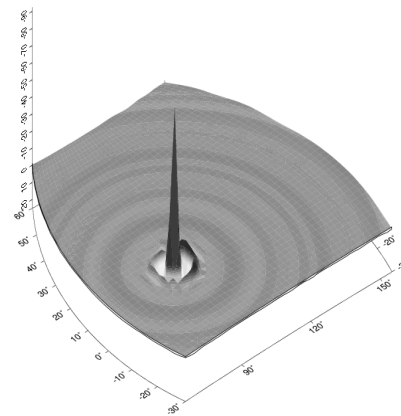
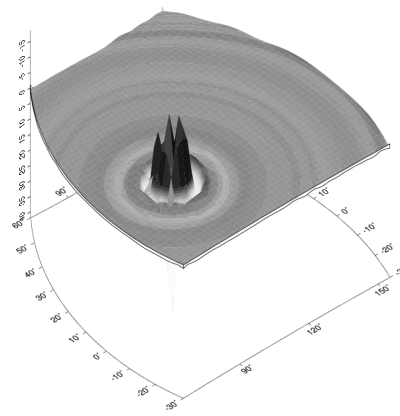
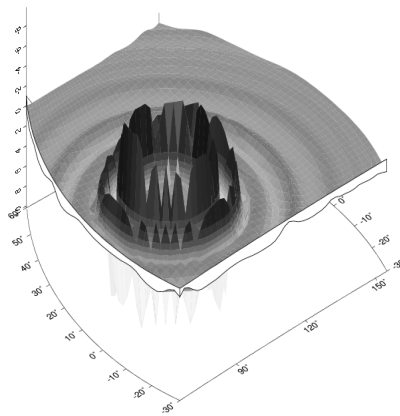
$t = -100s$



$t = 0s$

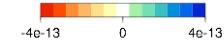
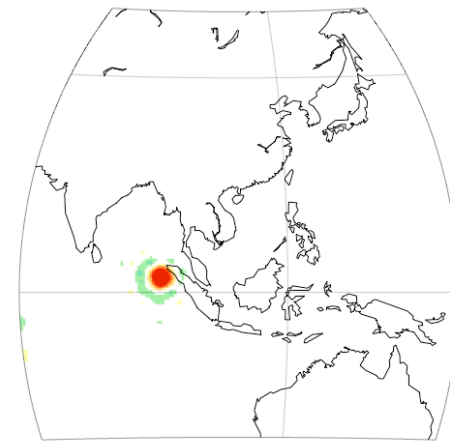
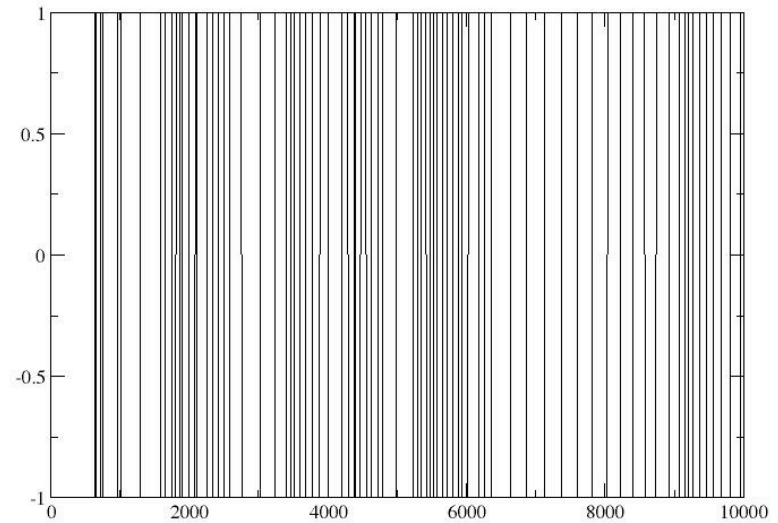
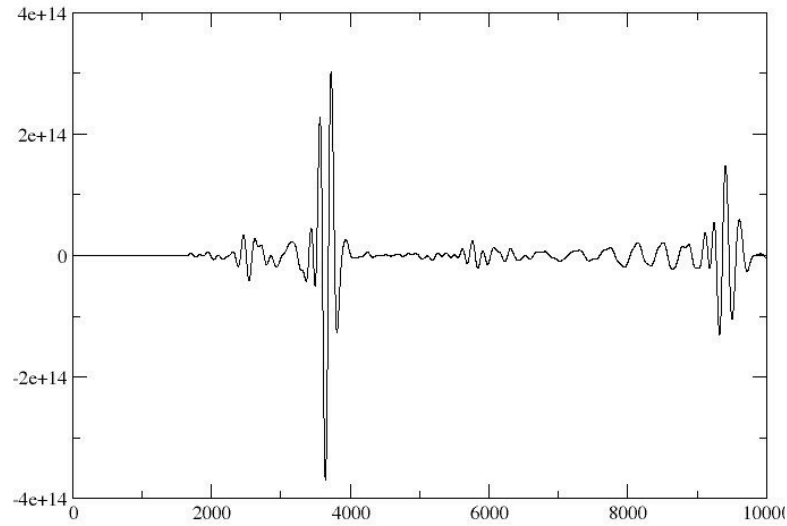


$t = 500s$

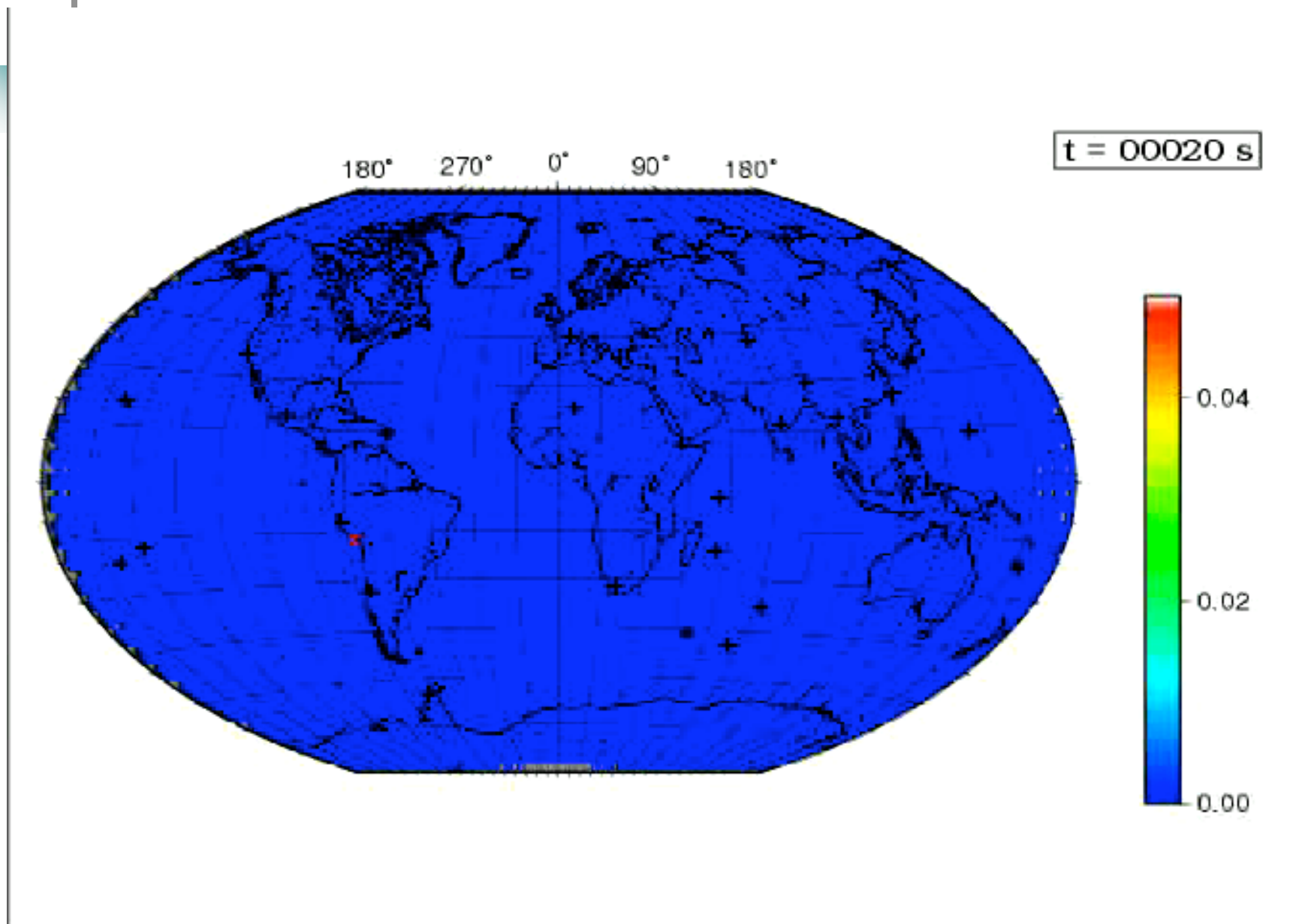


Complete seismogram

normalized

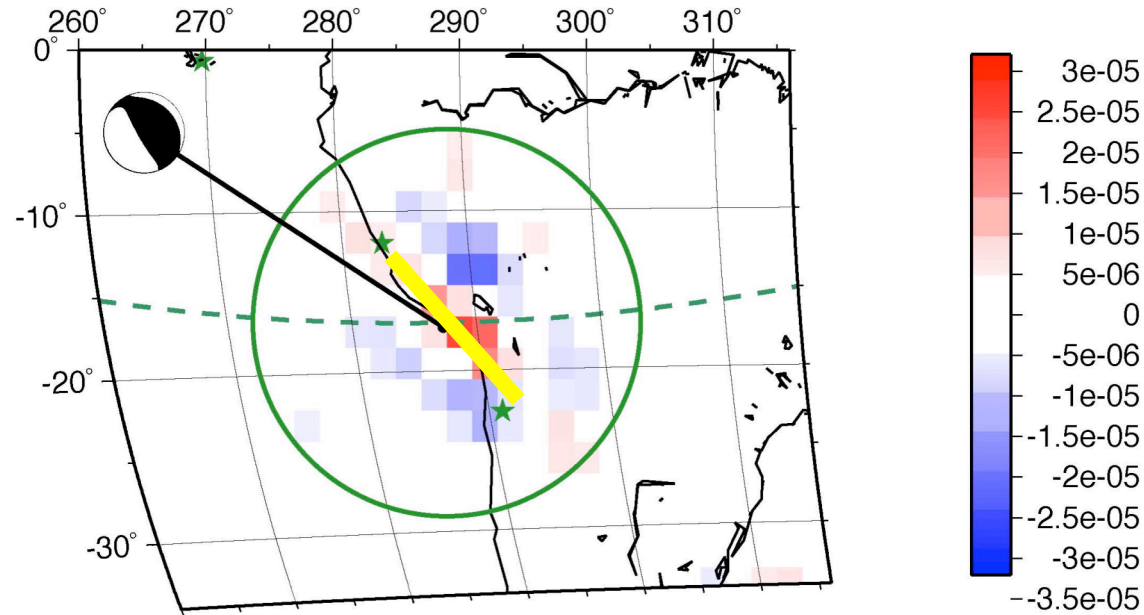


DATA: Peru Earthquake (23-06-2001) Mw= 8.4

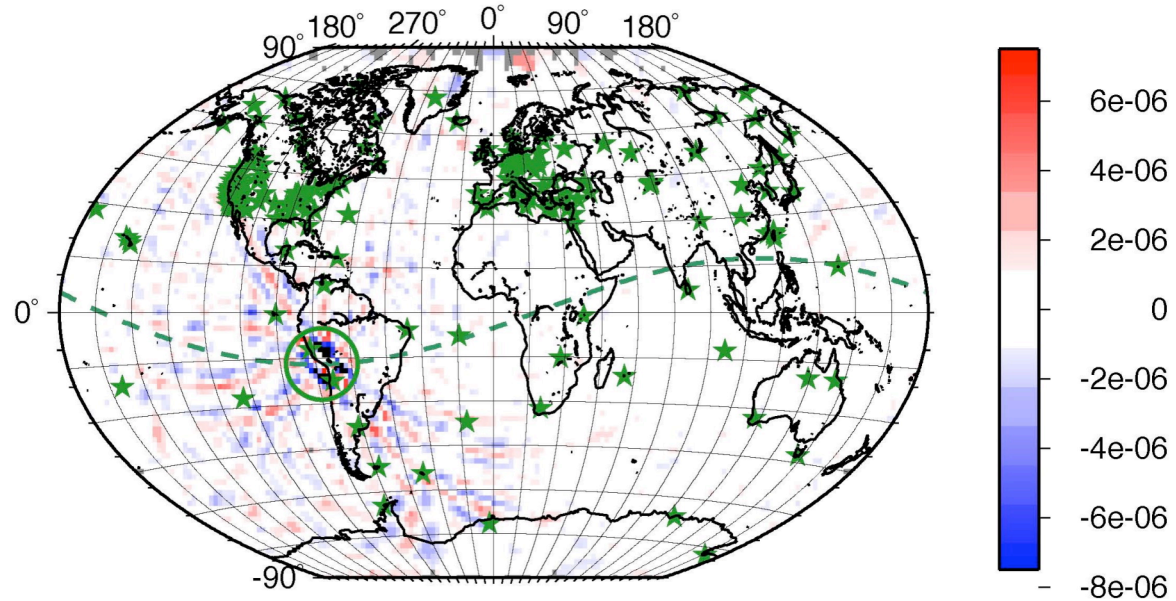


PERU 23 June 2001 - 8.4

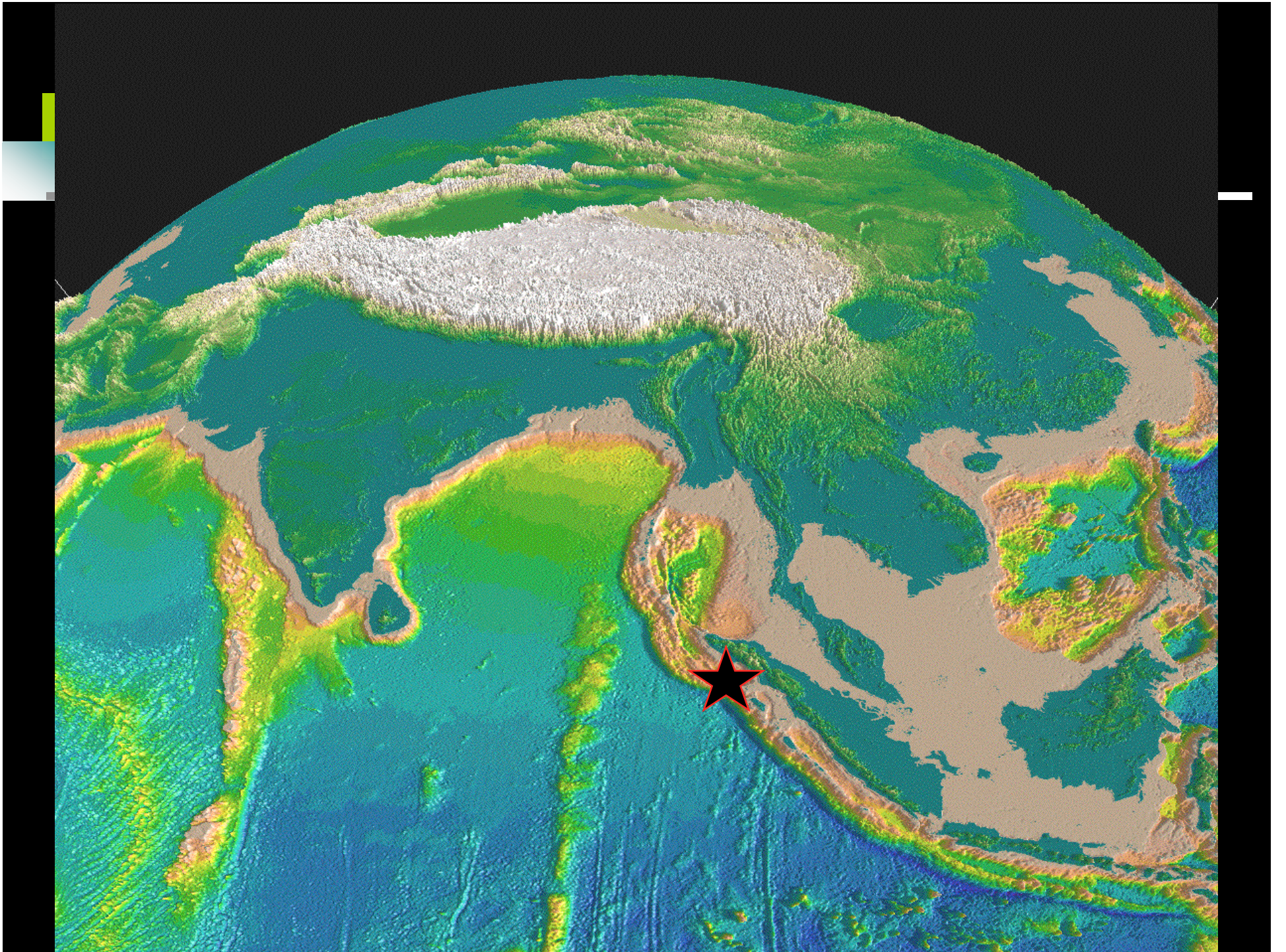
Fault Plane



$t = 0$ s

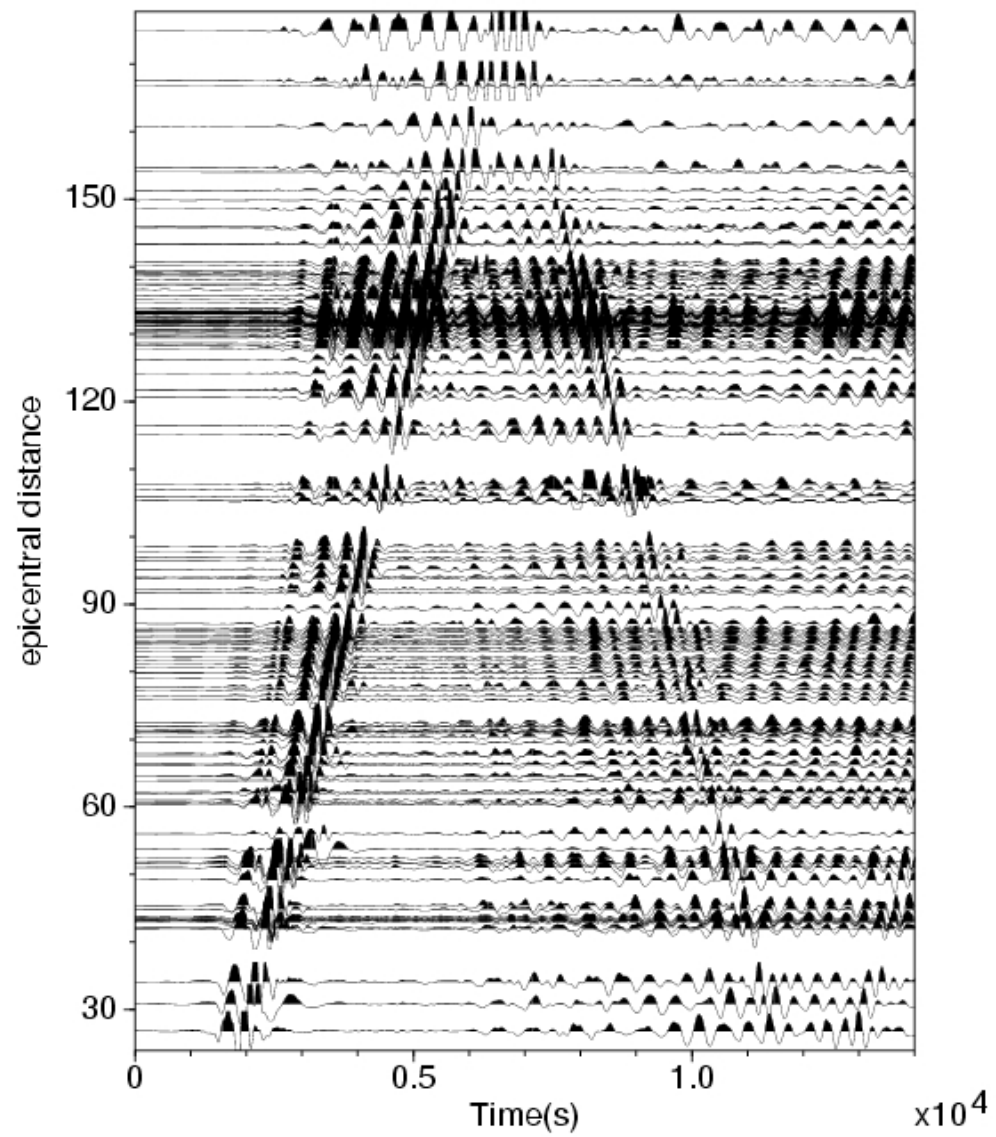


C. Larmat



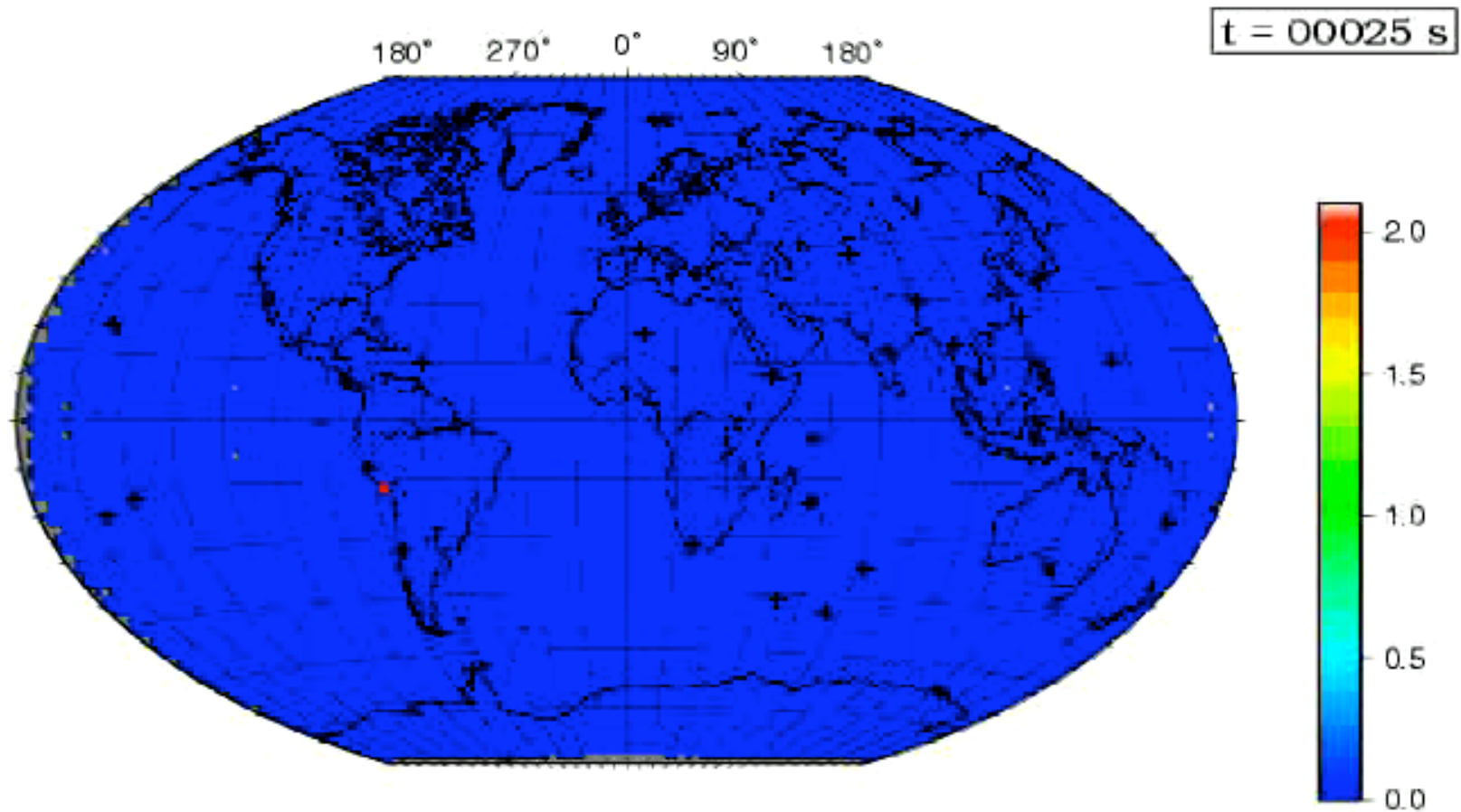


12/26/2004 data (1.2-9mHz)



The 121 real records we work with in this experiment (#11).

Sumatra Earthquake 26/12/04 -NM-SEM

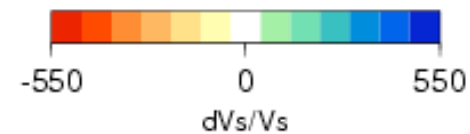
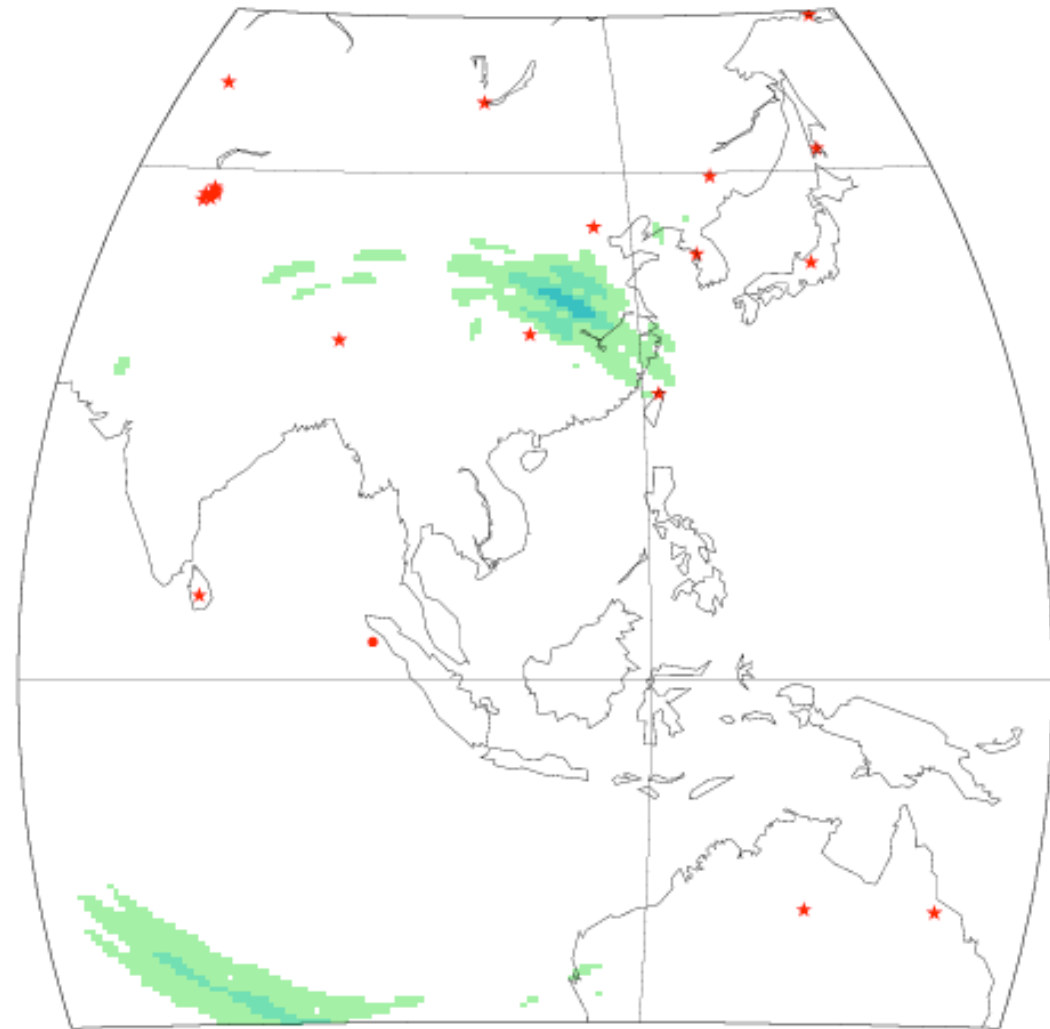


Sumatra

1D earth model

Synthetic Test
Normal modes

Point Source



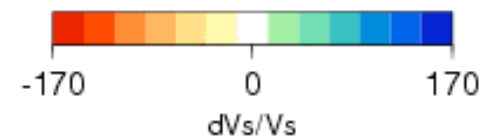
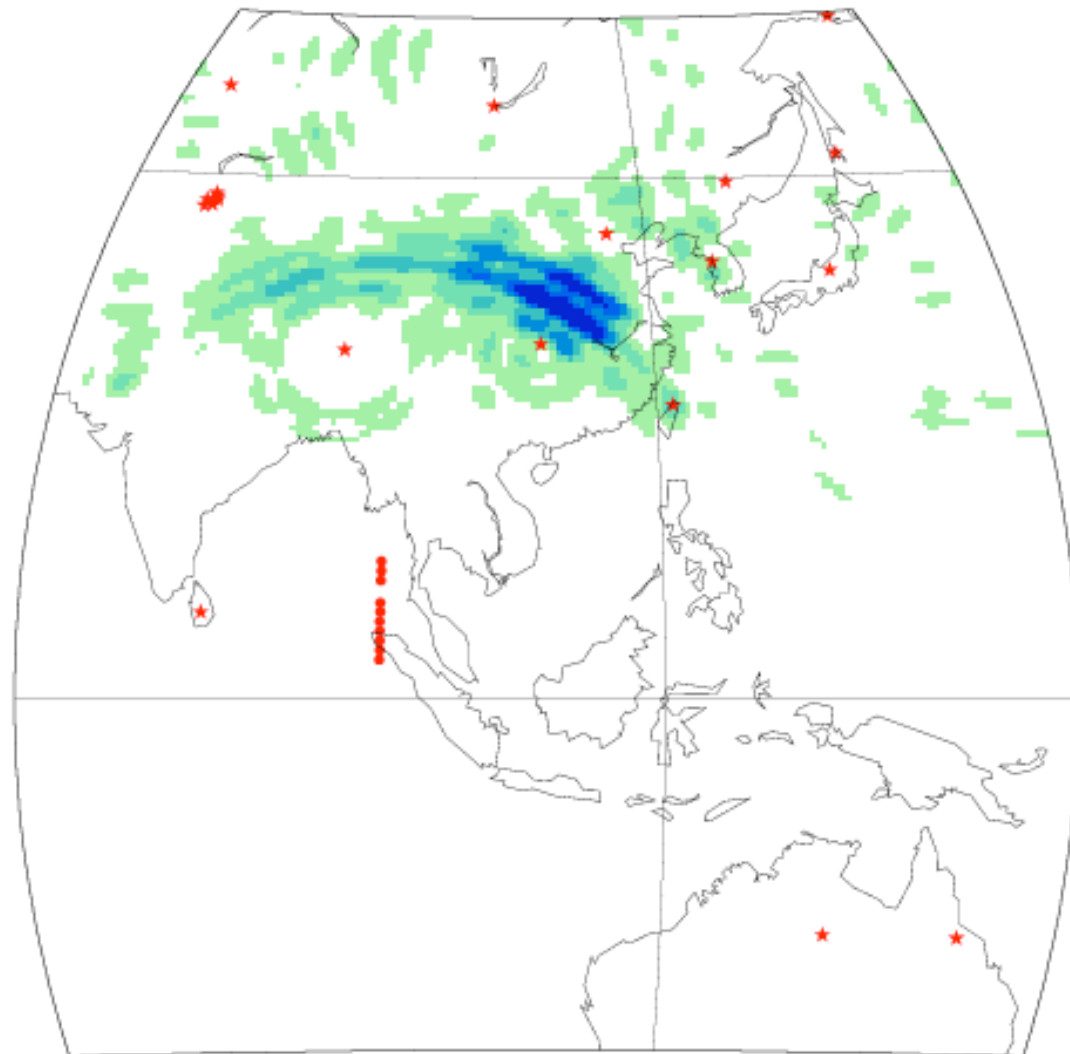
Y. Capdeville

Sumatra

Synthetic test

Normal modes:

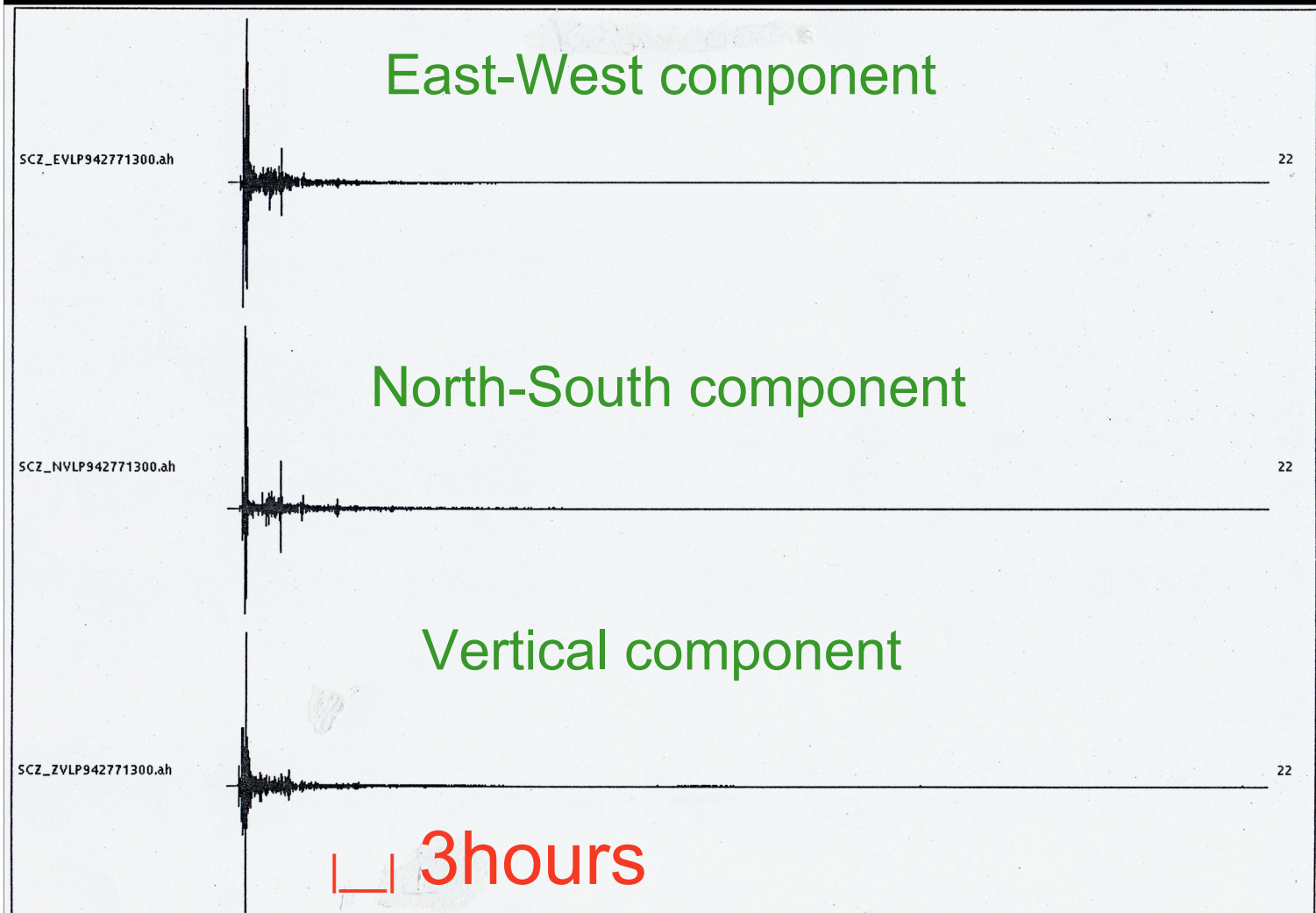
Extended
source



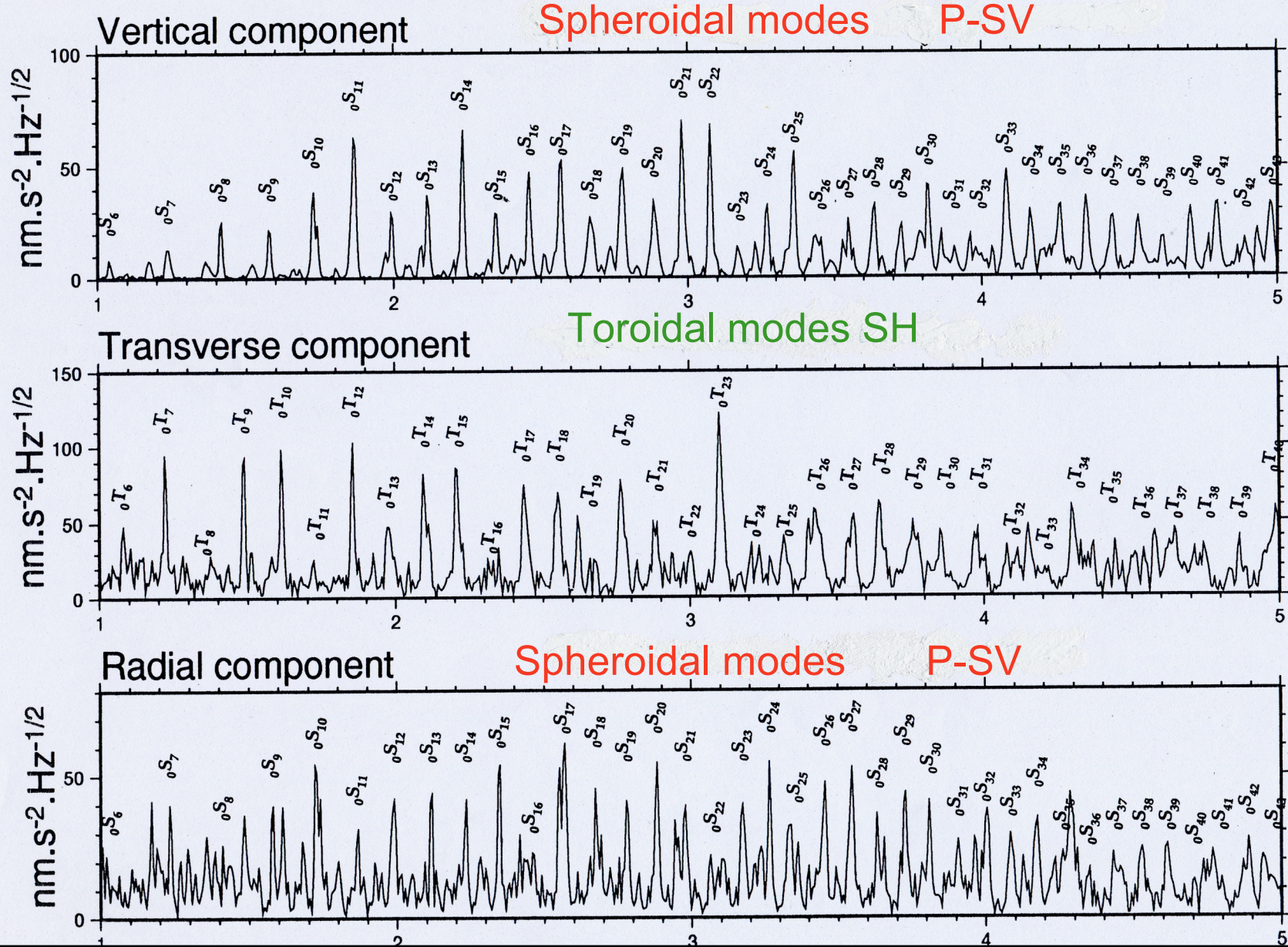
WHY does Time Reversal work when applied to seismic waves ?

Normal Mode Approach of Draeger & Fink (1999)
Cavity, Scalar case

Kurils islands 1994-277 Ms=8.3



KURIL 94 277 - SCZ VLP - 36h.

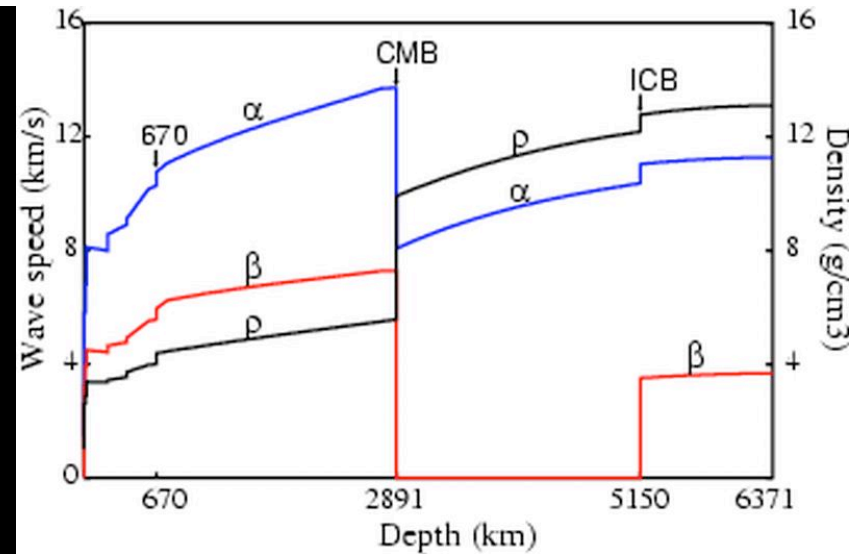


1D-Reference Earth Model:

$M_0(r)$, $\rho(r)$, $V_P(r)$, $V_S(r)$

(PREM, Dziewonski and Anderson, 1981)

$$\rho \partial_{tt} \mathbf{u}_0 + \mathbf{H}_0 \mathbf{u}_0 = \mathbf{0}$$



Eigenfrequencies: ${}_n\omega_l$

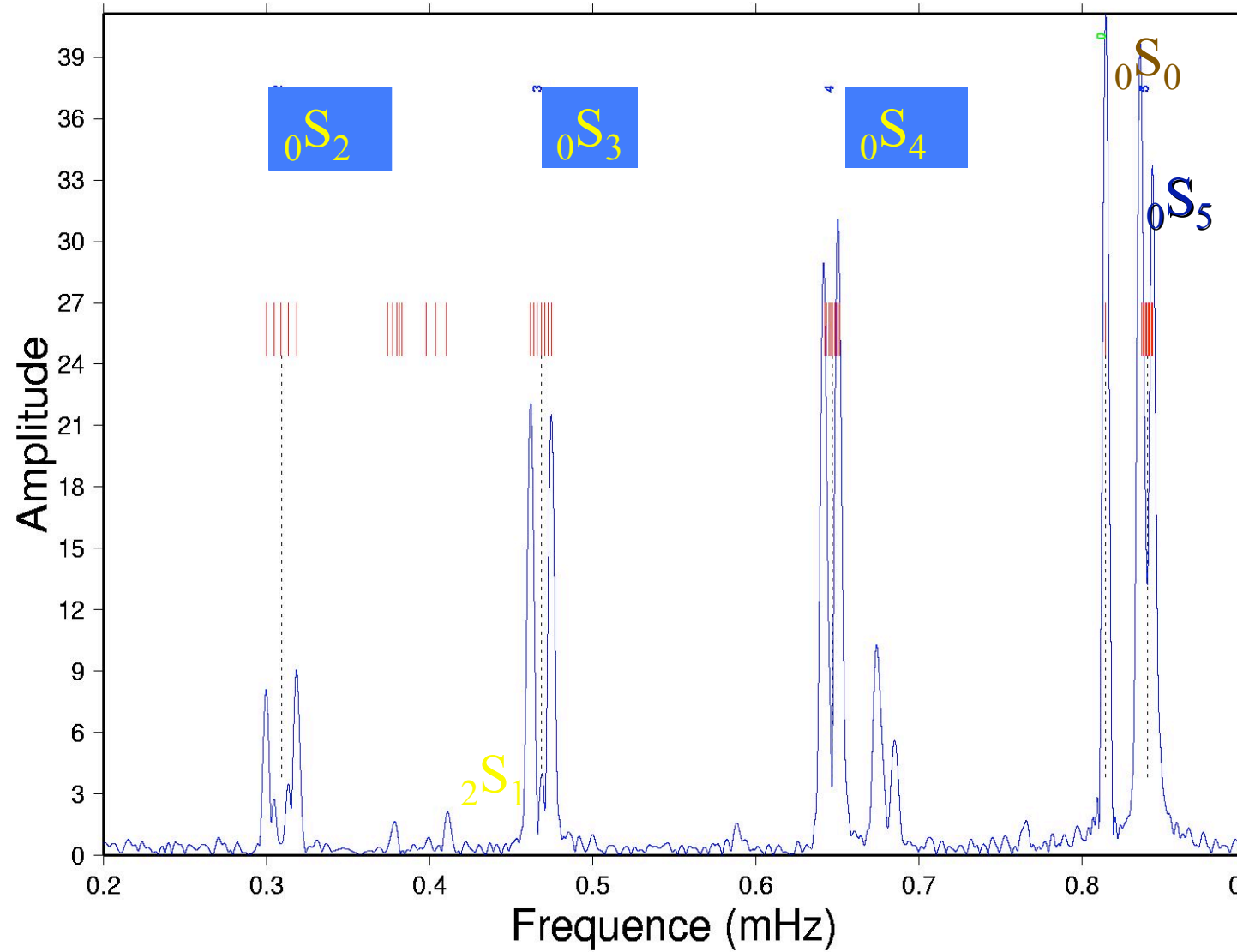
Eigenfunctions: ${}_n u_l^m(r,t) = |n,l,m\rangle$ (complete basis)

2 kinds of modes: Toroidal ${}_n T_l$, Spheroidal ${}_n S_l$

Degeneracy of eigenfrequencies ${}_n\omega_l$: $2l+1$
for radially symmetric models

Sumatra

can360.11.vhz, nb d'heures:118 h



Roult et
al, 2005

1D- Reference Earth Model

Seismic Source

$$\rho \partial_{tt} \mathbf{u} + \mathbf{H}_0 \mathbf{u} = \mathbf{F}_s$$

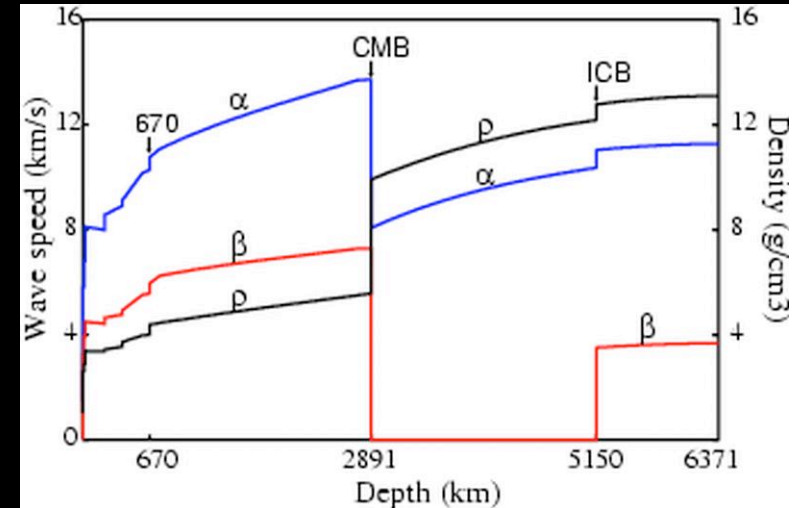
■ Synthetic Seismograms by normal mode summation ($\mathbf{k}=\{n,l,m\}$).

Displacement at point \mathbf{r} at time t due to a force system \mathbf{F} at point source \mathbf{r}_s

$$\mathbf{u}(\mathbf{r},t) = \sum_{\mathbf{k}} -(\mathbf{u}_{\mathbf{k}} \cdot \mathbf{F})_{\mathbf{s}} u_{\mathbf{k}}(\mathbf{r}) \cos \omega_{\mathbf{k}} t / \omega_{\mathbf{k}}^2 \exp(-\omega_{\mathbf{k}} t / 2Q_{\mathbf{k}})$$

Source Term $(\mathbf{u}_{\mathbf{k}} \cdot \mathbf{F})_{\mathbf{s}} = (\mathbf{M} : \boldsymbol{\varepsilon})_{\mathbf{s}}$

\mathbf{M} Seismic moment tensor, $\boldsymbol{\varepsilon}$ deformation tensor



Why does time reversal works when applied to seismic waves?

\mathbf{r}_R receiver, \mathbf{r}_S source location

$$\mathbf{u}(\mathbf{r}_R, t) = \sum_{\mathbf{k}} -(\mathbf{F} \cdot \mathbf{u}_{\mathbf{k}})_S \cos \omega_{\mathbf{k}} t / \omega_{\mathbf{k}}^2 \mathbf{u}_{\mathbf{k}}(\mathbf{r}_R)$$

$$\mathbf{u}(\mathbf{r}_R, t) = \mathbf{F}_S(t) * \mathbf{G}_{SR}(t)$$

Time reversed seismogram in \mathbf{r}_R : $\mathbf{F}_S(-t) * \mathbf{G}_{SR}(-t)$

in \mathbf{r}_M : $\mathbf{v}(\mathbf{r}_M, t) = \mathbf{F}_S(-t) * \mathbf{G}_{SR}(-t) * \mathbf{G}_{RM}(t)$

Why does time reversal works when applied to seismic waves?

for a point source,

If M in S, autocorrelation:

$$\mathbf{v}(\mathbf{r}_S, t) = \mathbf{G}_{SR}(-t) * \mathbf{G}_{RS}(t) = \int \mathbf{G}_{SR}(t+\tau) \mathbf{G}_{RS}(\tau) d\tau$$

If M not in S, cross-correlation:

$$\mathbf{v}(\mathbf{r}_M, t) = \mathbf{G}_{SR}(-t) * \mathbf{G}_{RM}(t) = \int \mathbf{G}_{SR}(t+\tau) \mathbf{G}_{RM}(\tau) d\tau$$

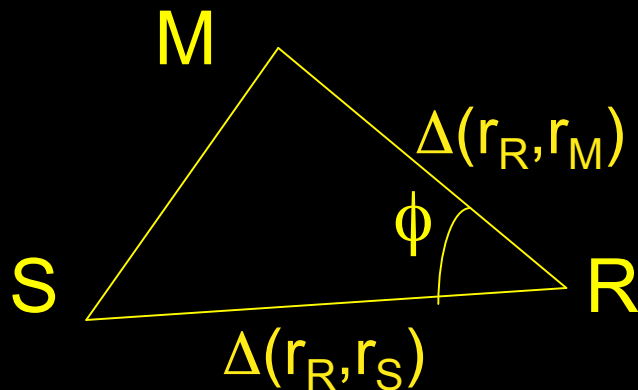
Why does time reversal works when applied to seismic waves?

$$\mathbf{v}(\mathbf{r}_M, t) = \sum_k \sum_{k'} \mathbf{u}_{k'}(\mathbf{r}_R) u_k(\mathbf{r}_S) F_S u_k(\mathbf{r}_R) \mathbf{u}_k(\mathbf{r}_M) \int \mathbf{g}(t, \tau) d\tau$$

k multiplet: $\{n, l, m\}$ $u_k(\mathbf{r}_R) = {}_n D_l(r_R) Y_l^m(\theta, \phi)$


Addition theorem: $\sum_k Y_l^m(\theta_1, \phi_1) Y_l^m(\theta_2, \phi_2) = P_l^0(\cos \Delta(r_1, r_2))$

$$\mathbf{v}(\mathbf{r}_M, t) = \sum_{n,l} \sum_{n',l'} {}_n D_l P_l^0(\cos \Delta(r_R, r_M)) F_S {}_n D_l P_l^0(\cos \Delta(r_R, r_S)) \int \mathbf{g}(t, \tau) d\tau$$

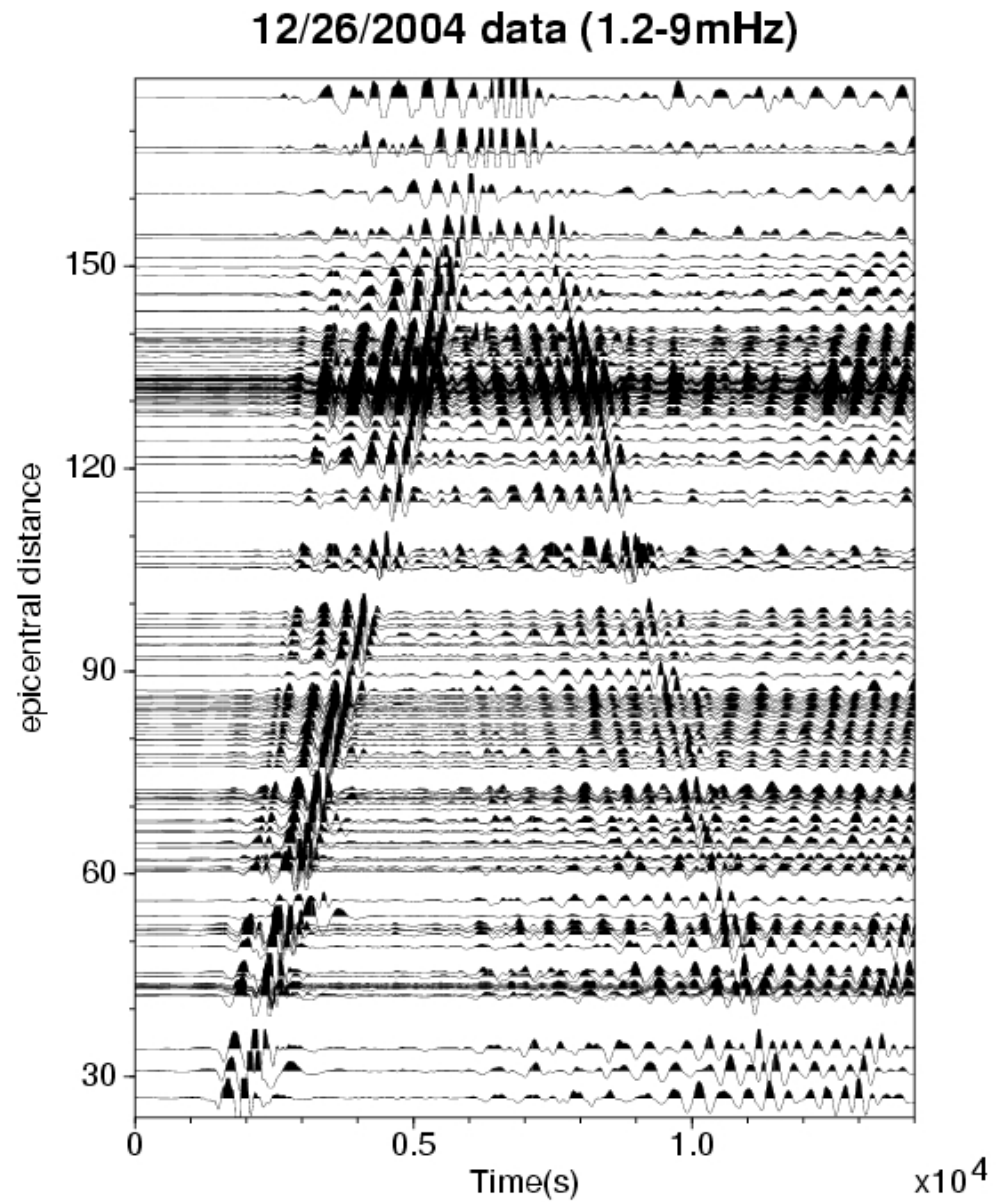


=> Max if $\phi = 0$ or π

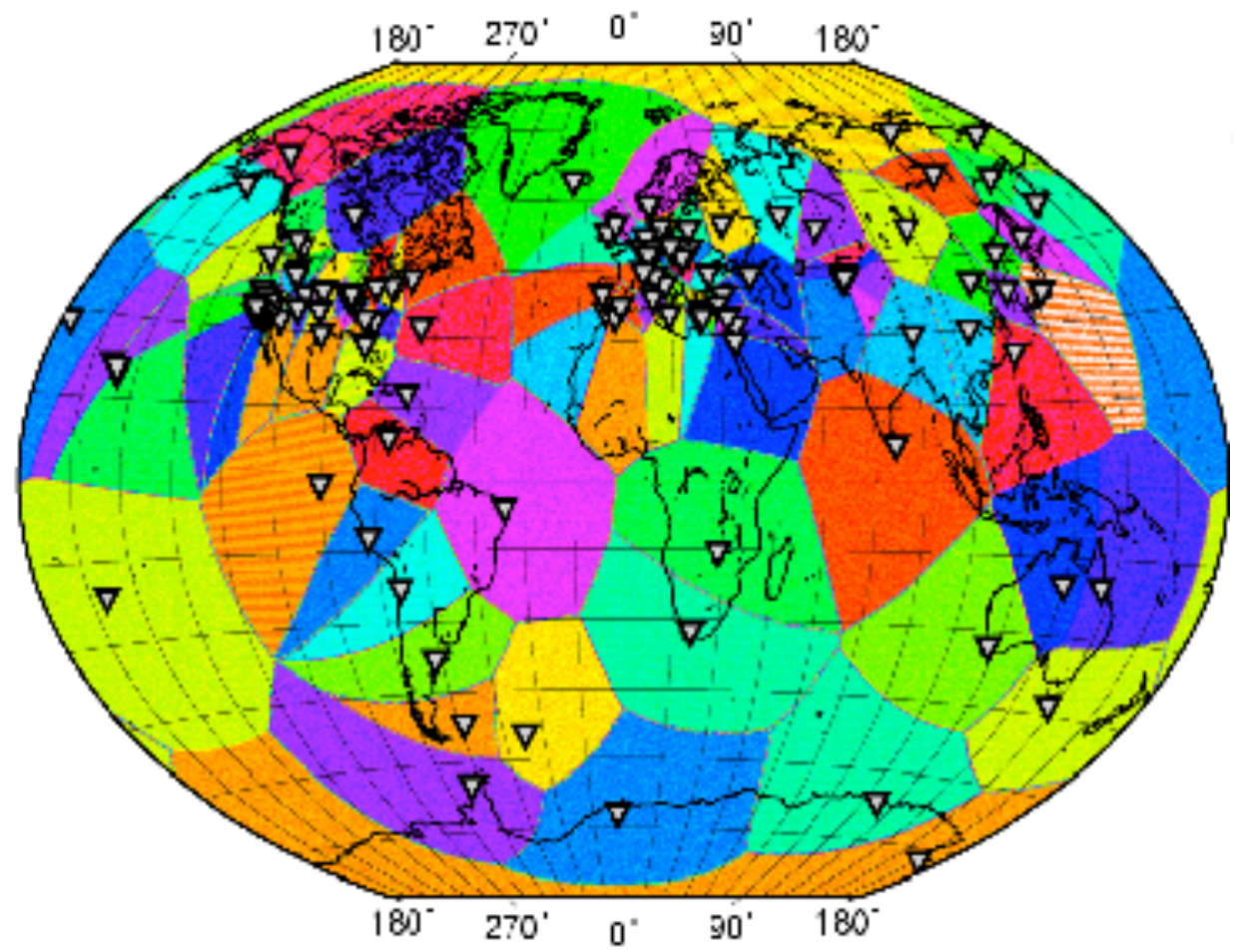
(Romanowicz, Snieder, ...)



Sumatra- Andaman Earthquake



The 121 real records we work with in this experiment (#11).



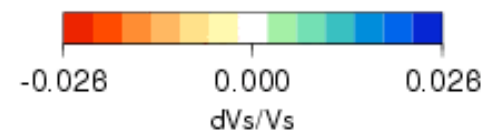
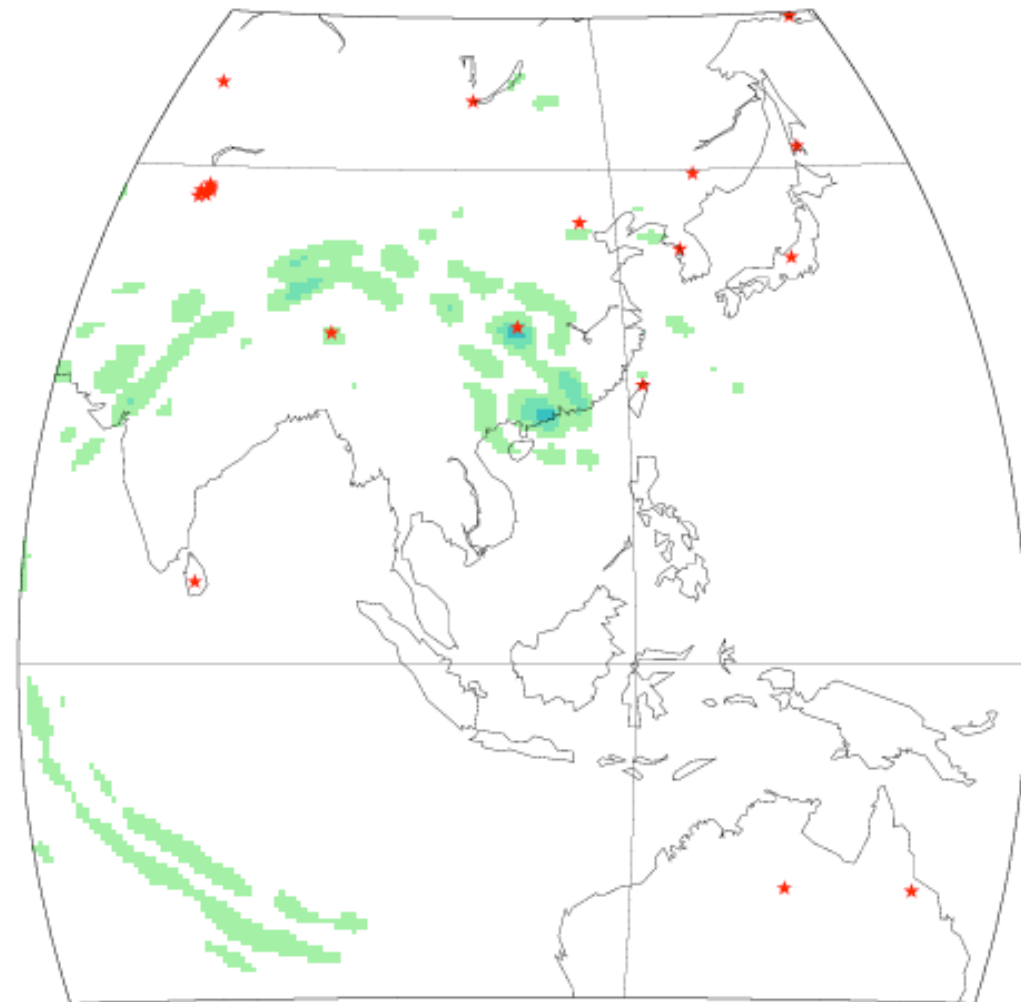
b) Stations network in a seismic TR experiment



Sumatra

Normal mode
Time reversal

Real Data



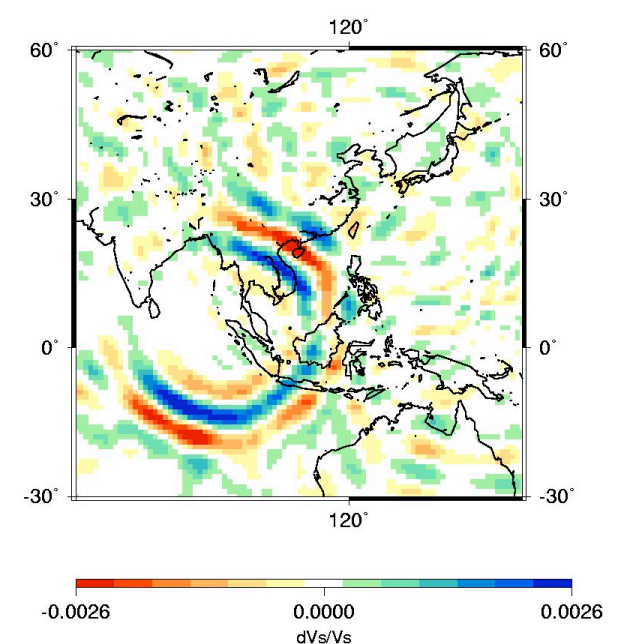
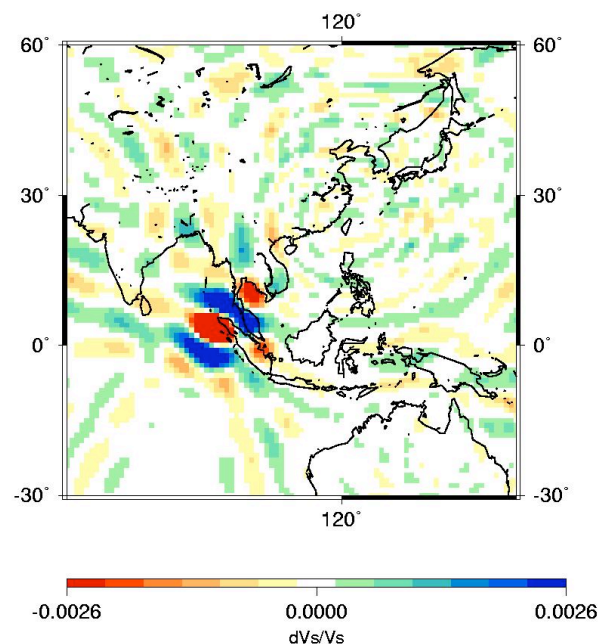
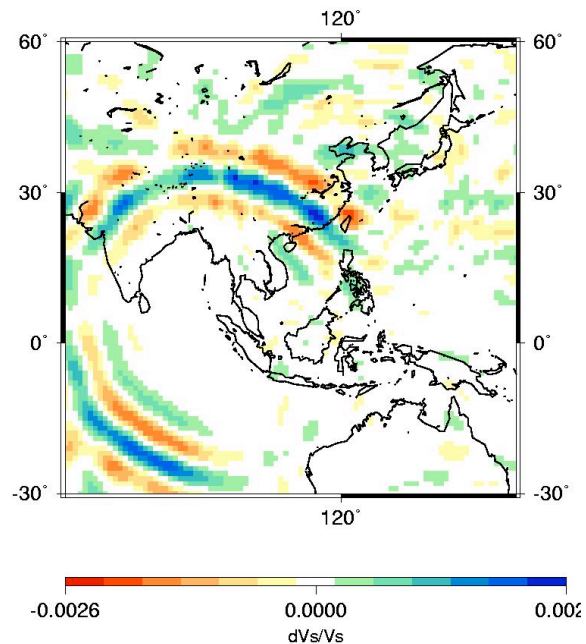
First conclusions

- Time reversal focus at the right time ($t_0 \approx -7000s$)
- and at the right place

$t = -6000s$

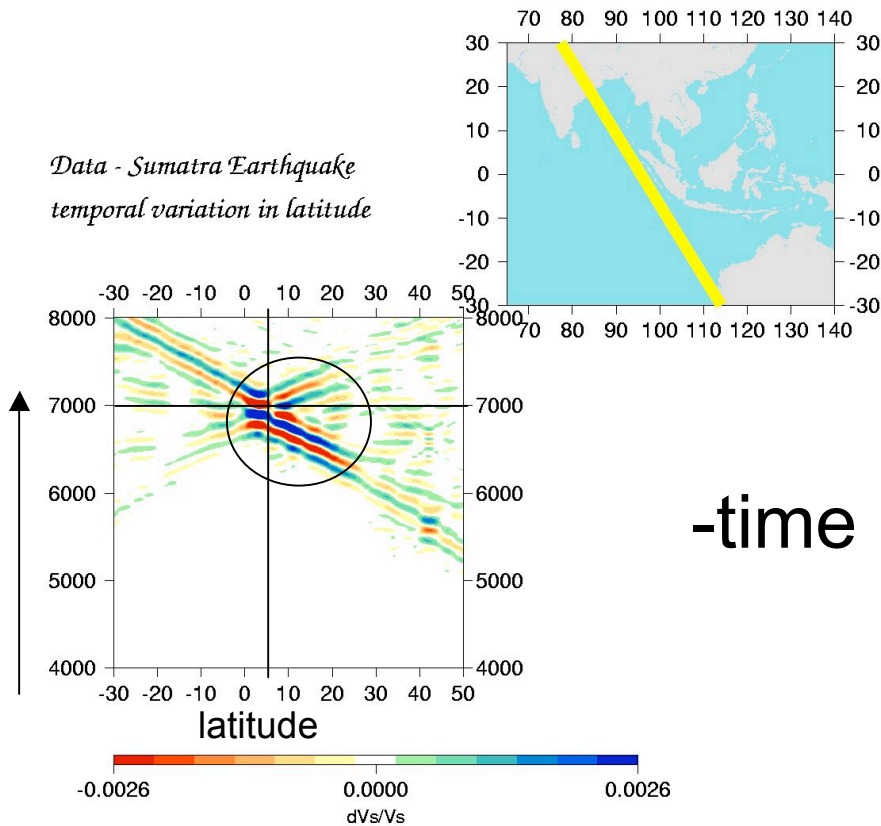
$t = -7000s$

$t = -7500s$

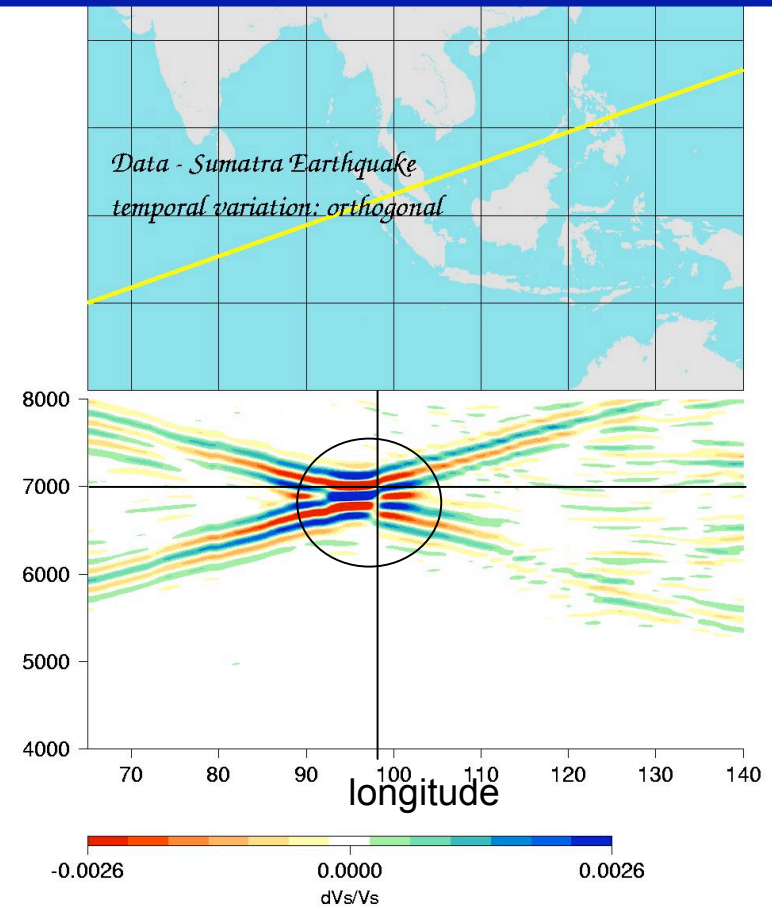


Can we get information about the history Of the seismic rupture?

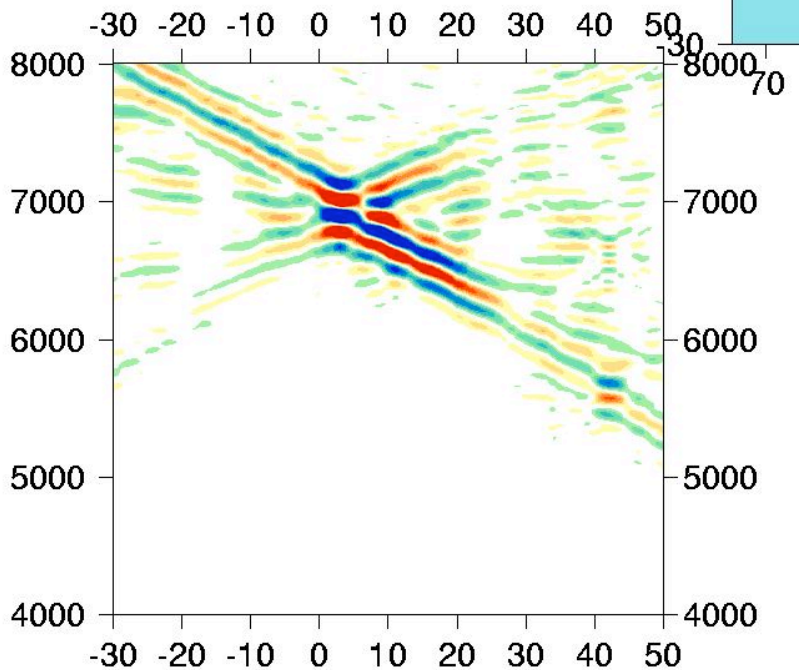
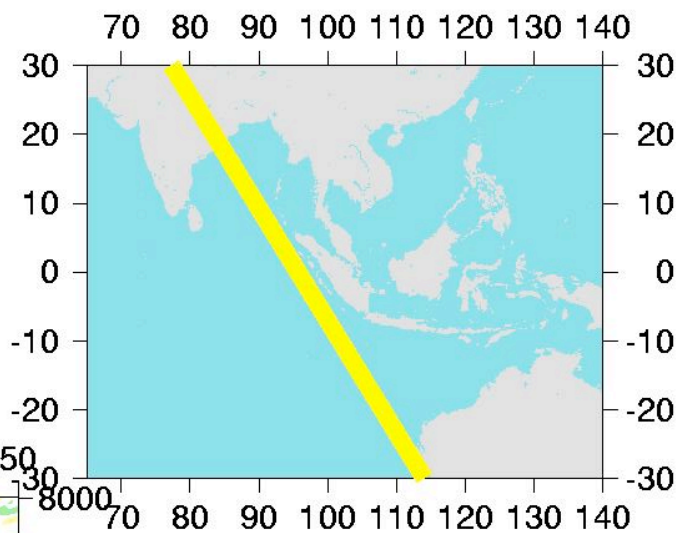
*Data - Sumatra Earthquake
temporal variation in latitude*



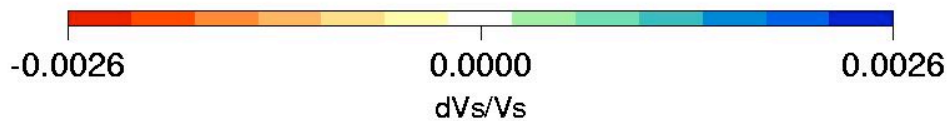
*Data - Sumatra Earthquake
temporal variation: orthogonal*



*Data - Sumatra Earthquake
temporal variation in latitude*



↑
-time





Source Rupture Imaging

$$\mathbf{u}_z(\mathbf{r}, t) = \sum_k -u_k(\mathbf{r}) \cos \omega_k t / \omega_k^2 \exp(-\omega_k t / 2Q_k) (\mathbf{u}_k \cdot \mathbf{F})_s$$

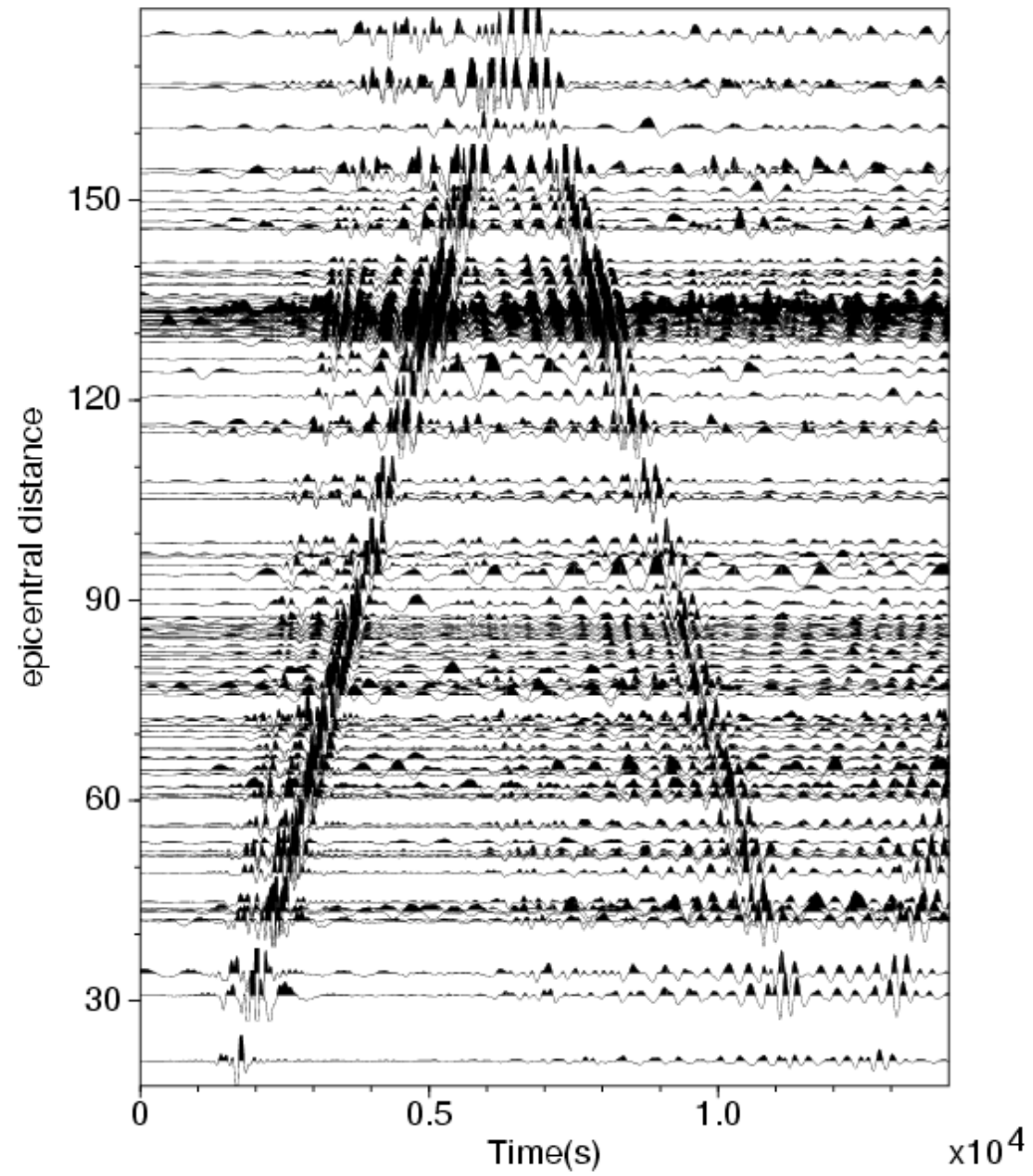
$$\mathbf{u}_z(\mathbf{r}, \omega) = \mathbf{G}(\mathbf{r}, \mathbf{r}_s, \omega) S(\mathbf{r}_s, \omega)$$

$\mathbf{G}(\mathbf{r}, \mathbf{r}_s, \omega)$ Green Function

$S(\mathbf{r}_s, \omega)$ Source Function

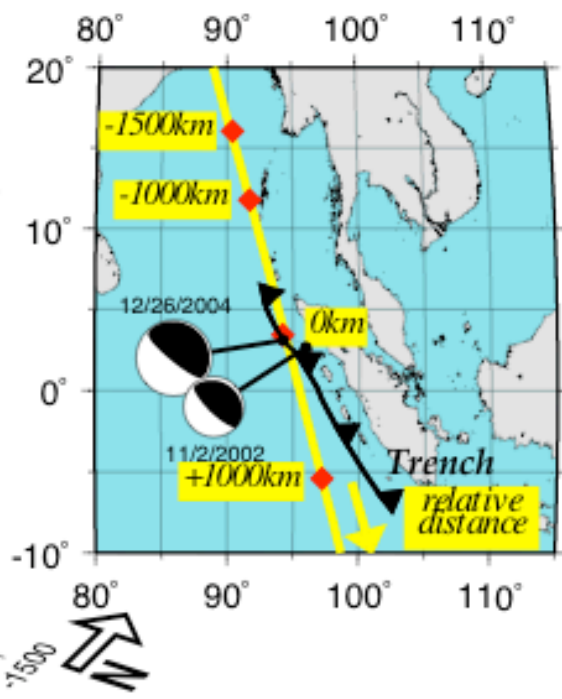
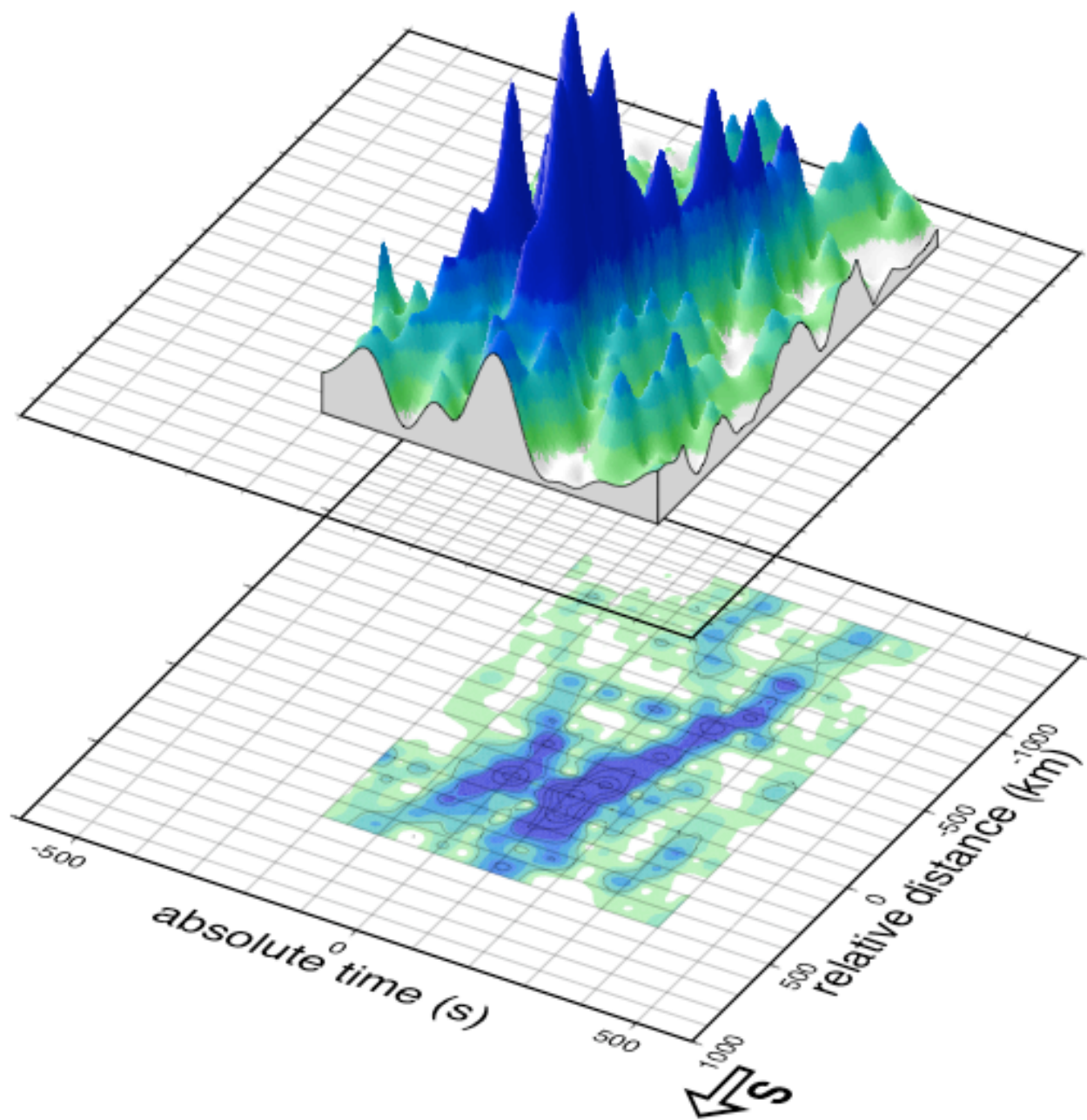
=> Reference source: delta function?

11/02/2002 data (1.2-9mHz)

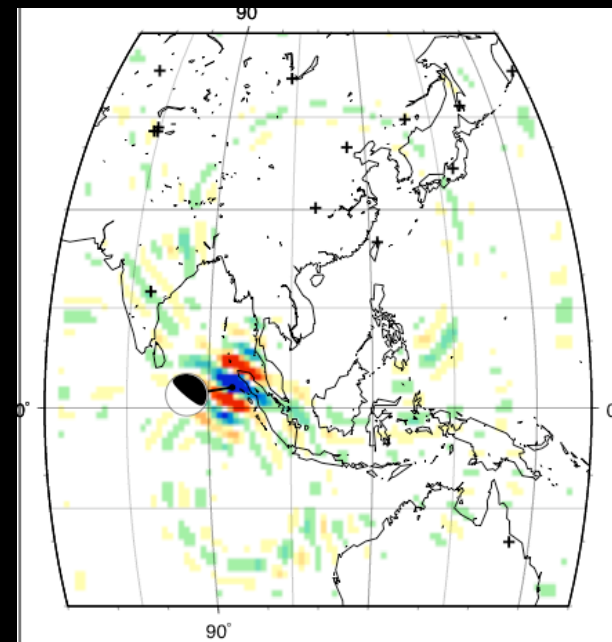
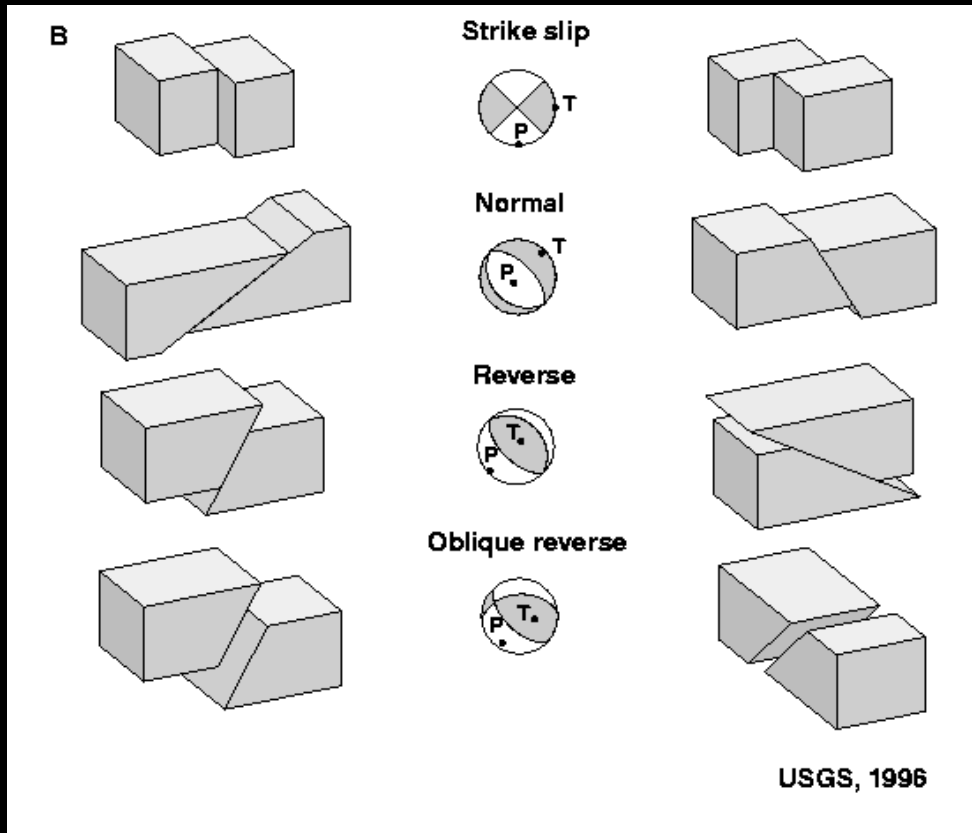


The 121 real records we work with in this experiment (#12).

absolute value of the displacement

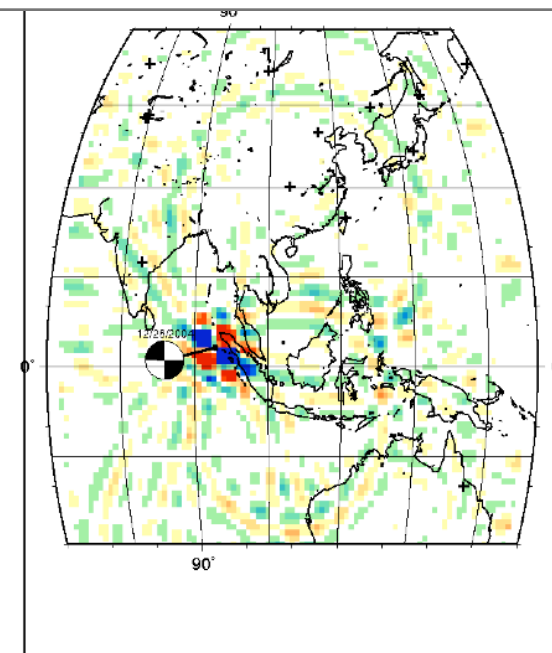


Different types of faults



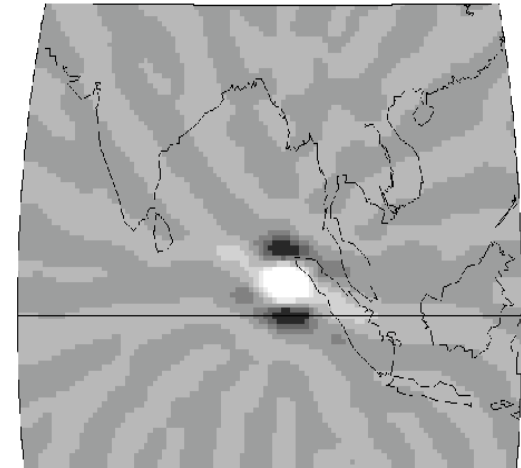
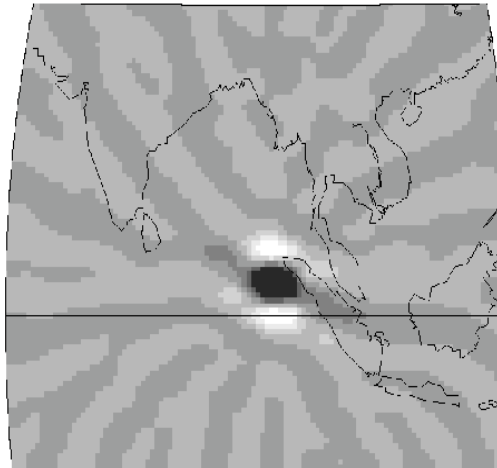
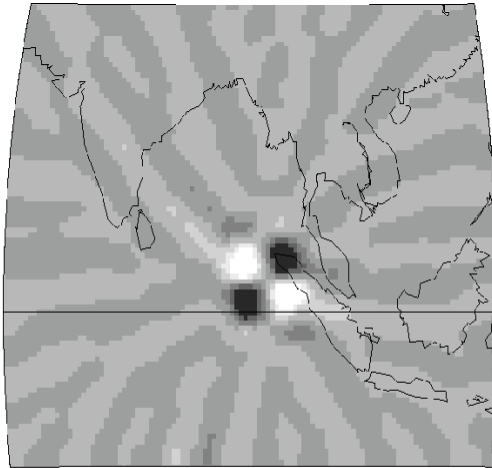
m

81



8

Different Radiation Patterns



Mozambique 22/02/2006 at 22h 19min 15.7s (TU).

latitude : -21.32°

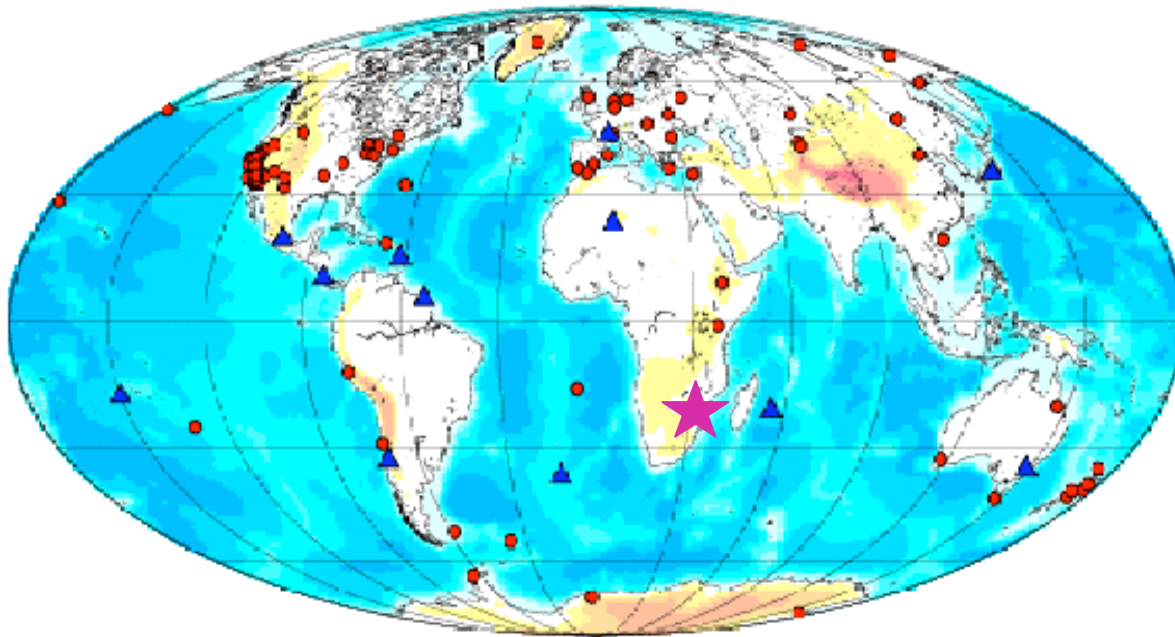
$M_w = 7.5$

longitude : 33.37°

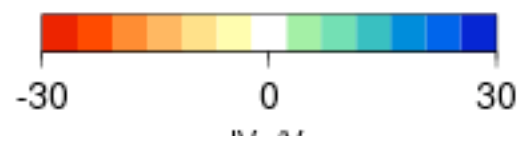
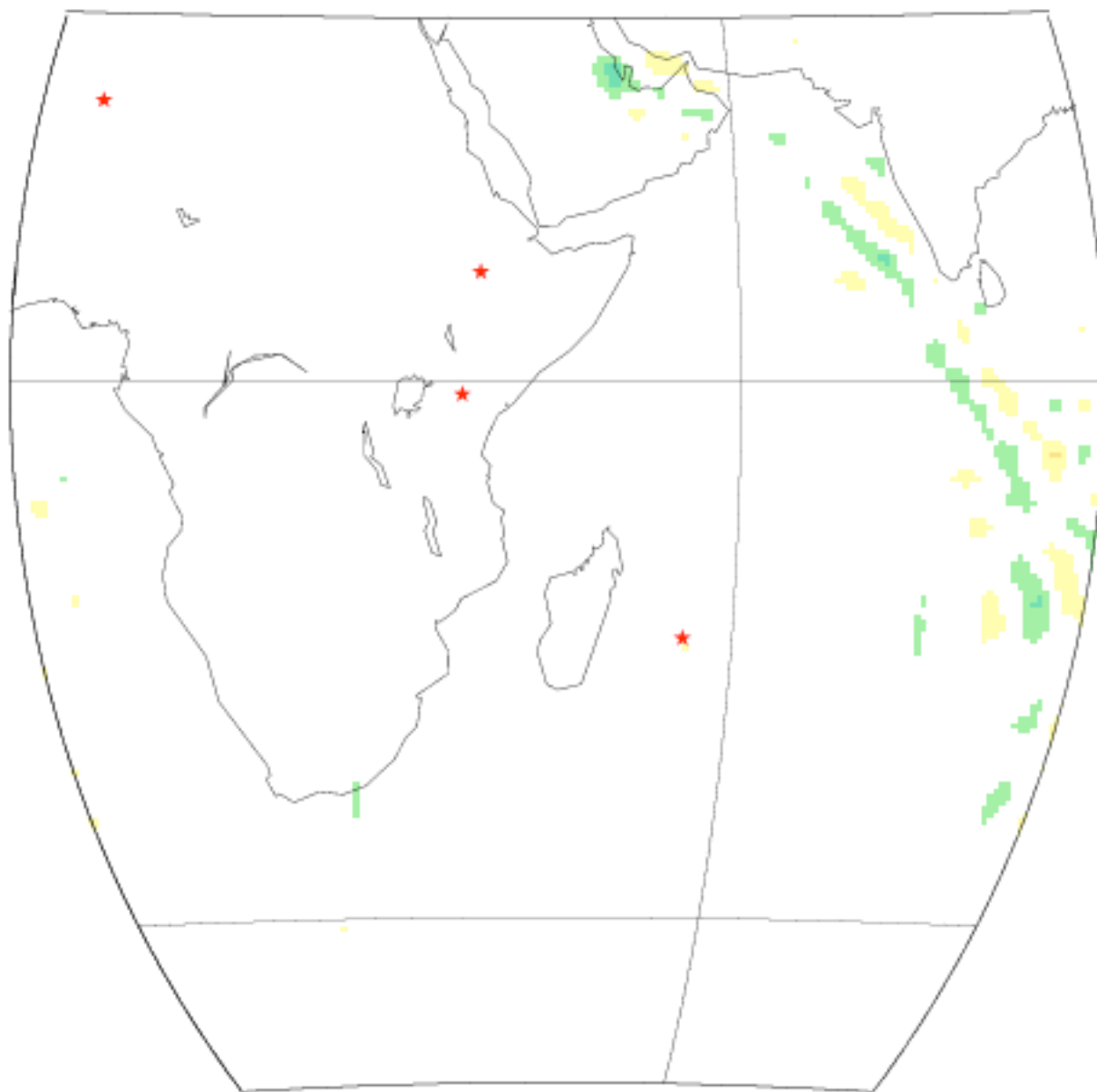
Normal fault N168°

depth : 12 km

Data of 112 stations filtered between 2 et 9 mHz.

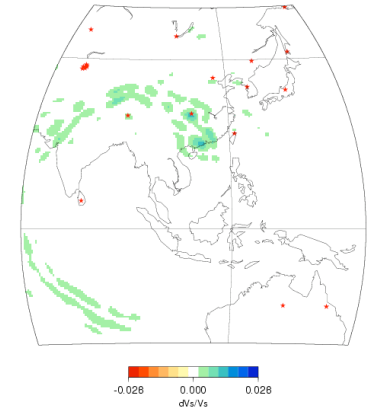


- Stations IRIS
- ▲ Stations GEOSCOPE
- ★ Location of earthquake

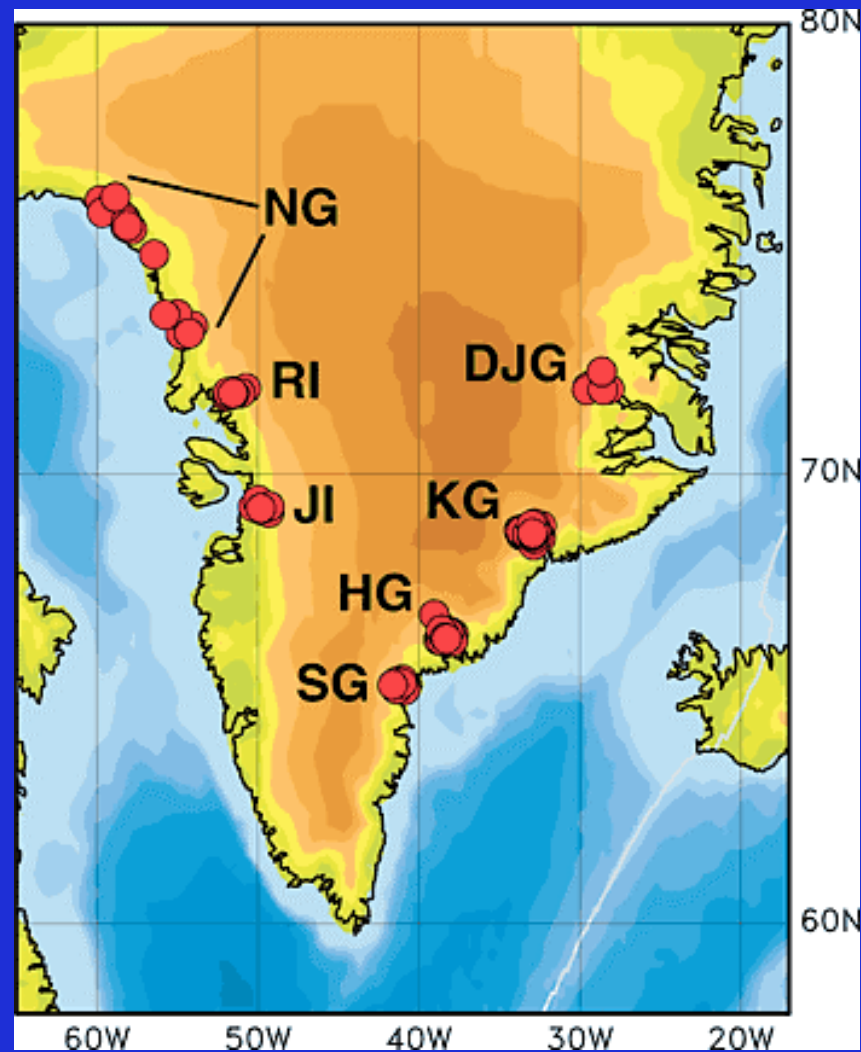
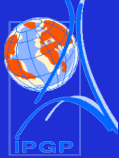


TIME REVERSAL

- Application to real seismograms with broadband FDSN stations
- Automated localization in time and in space of earthquakes
- Spatio-temporal Imaging of seismic source
<http://www.gps.caltech.edu/~carene>
<http://www.ipgp.jussieu.fr/~larmat>
- Detection of unknown seismic sources
(Glacial, “quiet” earthquakes, Seismic “Hum” of the Earth)
- Applications to seismic Tomography- Detection of mantle plumes...



Glacial Earthquakes



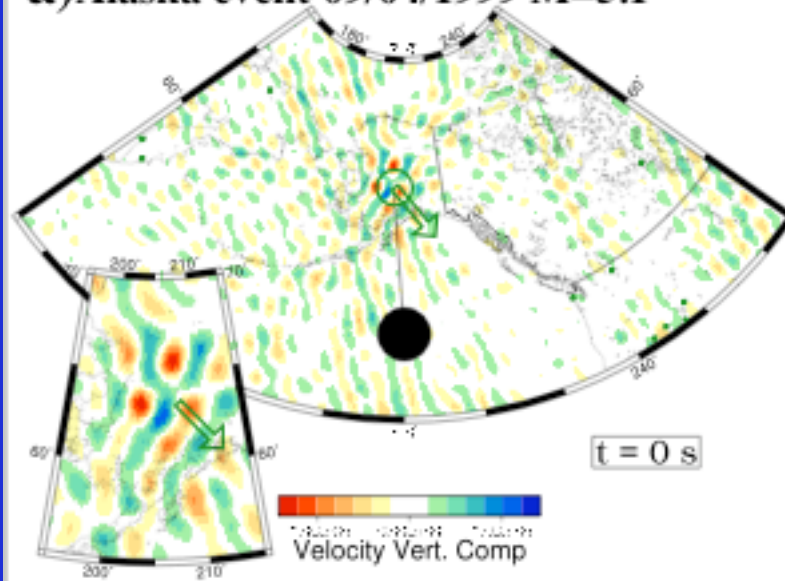
Ekstrom et al., 2003



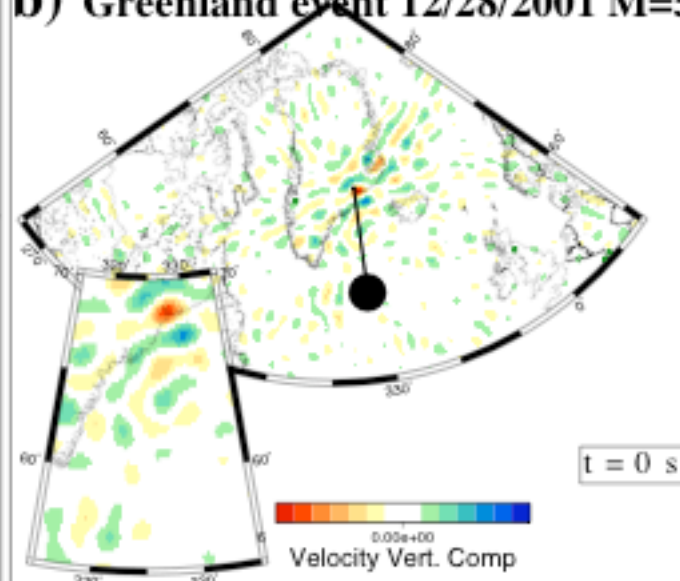
Glacial Earthquakes

- Carene Larmat, Jeroen Tromp (CalTech)
- Can we locate glacial earthquakes by time reversal?

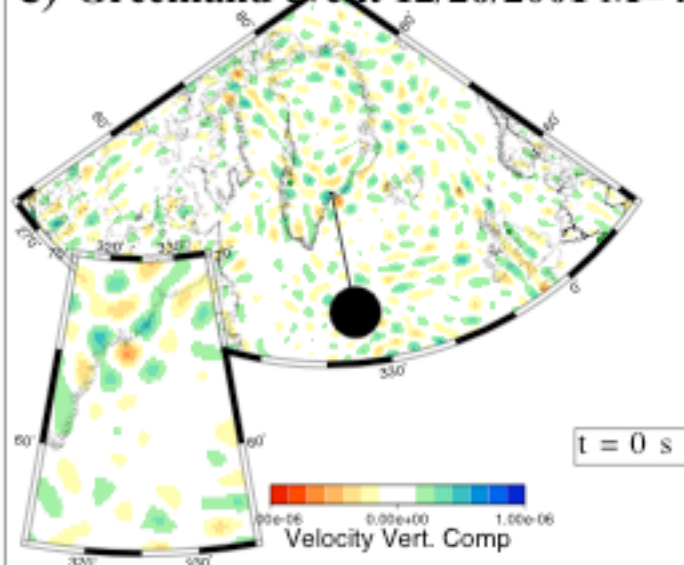
a) Alaska event 09/04/1999 M=5.1



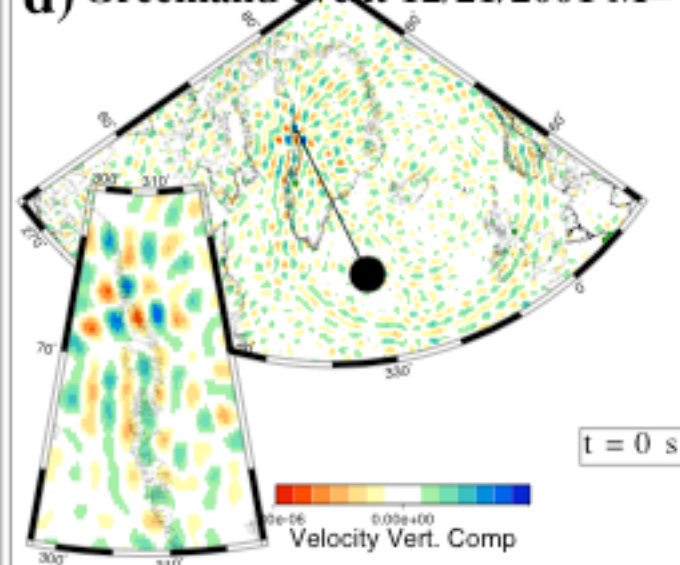
b) Greenland event 12/28/2001 M=5.0



c) Greenland event 12/26/2001 M=4.7



d) Greenland event 12/21/2001 M=4.8



- Time reversal of seismic waves

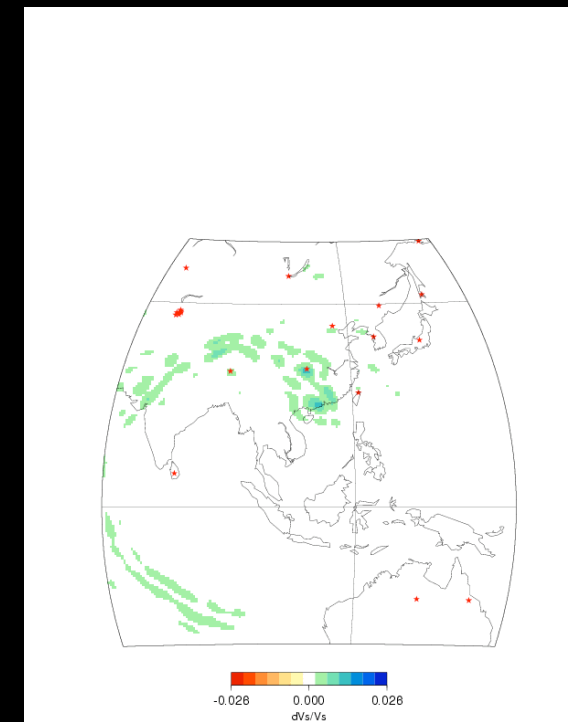
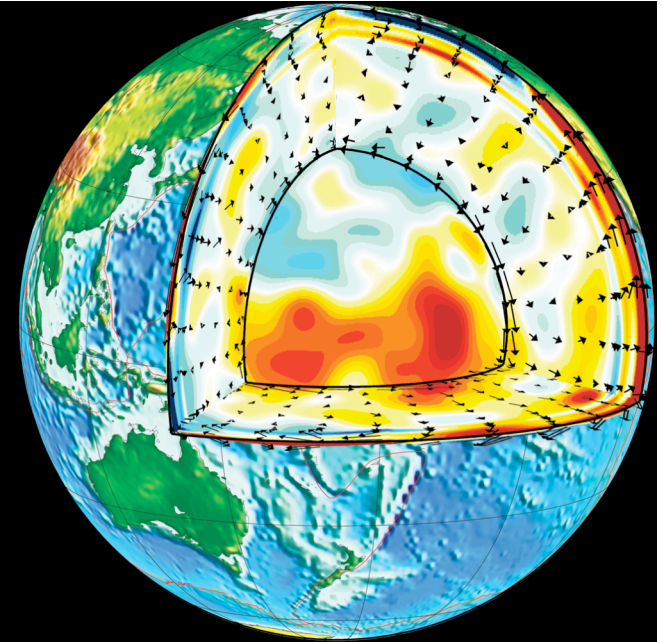
- ✓ Scientific challenges:

- Earthquakes in 3D models
(localisation, rupture, ...)

- Seismic noise....

- From global to regional scale
($T > 200s \rightarrow 30s$)

- Earth structure (Adjoint Tomo):
Image geological objects
(mantle plumes, slabs, ...)



Tsunami of 26/12/04

