The inverse problem of earthquake source mechanics with several applications

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> > Cargèse May 2007

The inverse problem of earthquake source mechanics :

Given a set of seismograms, find the details of the earthquake rupture process

$$u_k(\mathbf{x}_1, t_1) = \int_0^{t_1} dt \iint_{\Sigma} K_{ik}(\mathbf{x}_1, \mathbf{x}, t_1, t) \dot{a}_i(\mathbf{x}, t) dS$$

Seismogram = Green function kernel * slip rate

* means space-time convolution

The method of inversion is fully described in Das & Kostrov (JGR, 1990; PEPI, 1994) with all relevant equations fully written out

$\mathbf{A}\underline{x} = \underline{b}$

min ℓ_1 norm of $(\mathbf{A} \underline{x} - \underline{b}) = \underline{r}$

Constraints:

 $\underline{x} \ge 0$

 $x_i = 0$ for *i*'s prevented from slipping by causality Additional constraints such as preassigning seismic moment, looking for the smoothest solution that fits the data, etc. are easily incorporated into the method. It results in adding additional equations to the original system.

We shall use the method of Linear Programming to solve the problem.

(See Numerical Recipes by Press et al. 1989 for a simple version of the program and explanation of details of the method)

Synthetic tests:

Artificial data is constructed for some chosen rupture process and the data inverted using the *same* station distribution as would be used to invert the real data, giving us insight into the inversion – many tens of such tests are carried out.

Robustness tests

We test the preferred solution in the following way:

If we have a region of high (or low) slip, can we try to reduce (or increase) the slip in that region and still fit the data. This again leads to additional equations in our original system of equations. We even find the fault shape and area in this way.

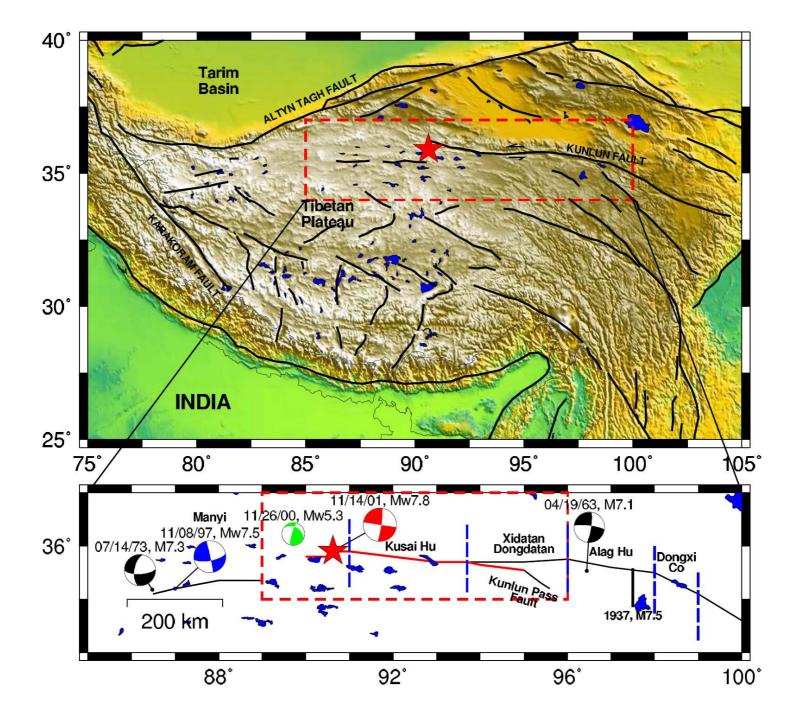
Robustness tests (contd)

If we have a very high rupture speed, we test if we can decrease it *and still fit the data*, and so on, testing all interesting aspects of the solution

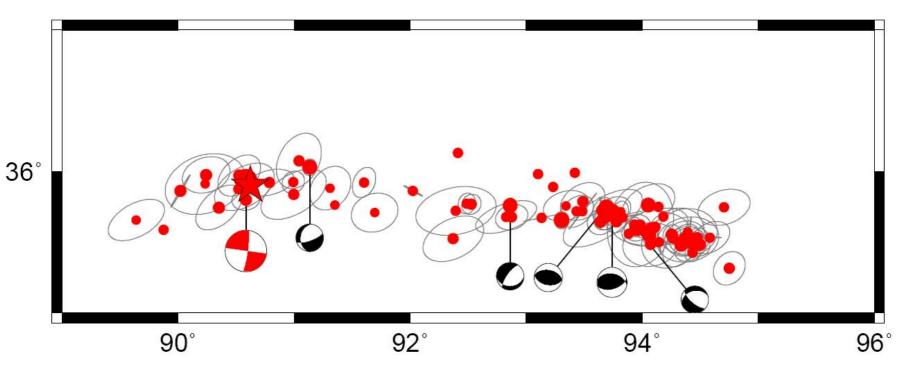
Unlike other methods of inversion, we allow the *entire* region behind the main rupture front to slip, if the data requires it. This is important to find if there is slip well behind the main rupture front. Theoretical models showed that such slip is possible.

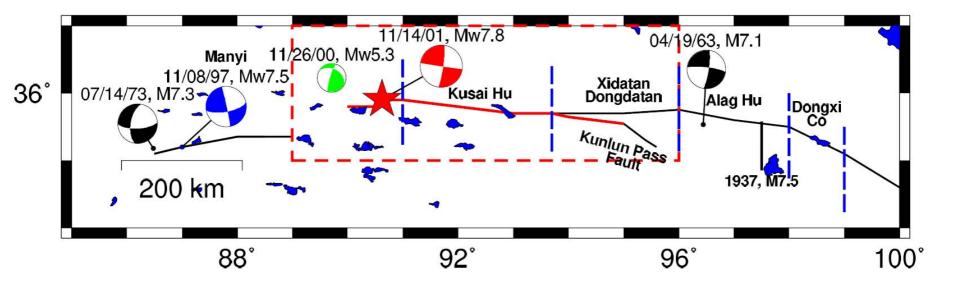
Also, unlike some other studies, we use only pure waves, e.g. when we invert SH waves, say, we terminate the seismogram when either ScSor SS arrive.

The Mw 7.8 2001 Kunlun, Tibet earthquake

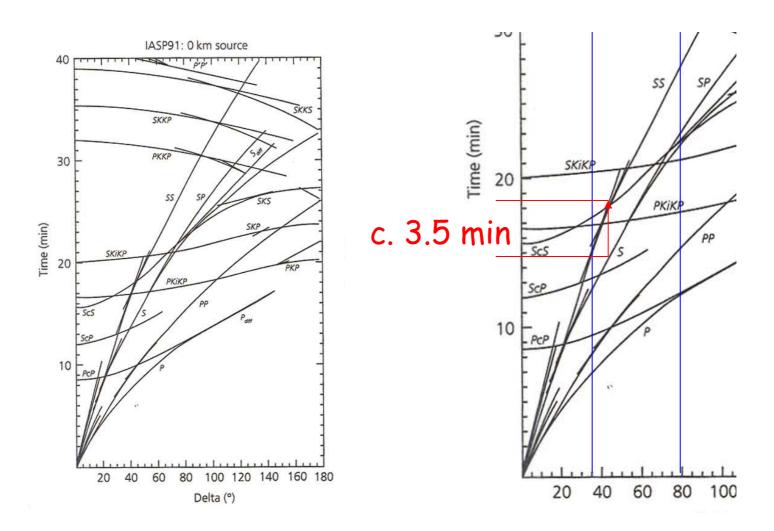


6-month

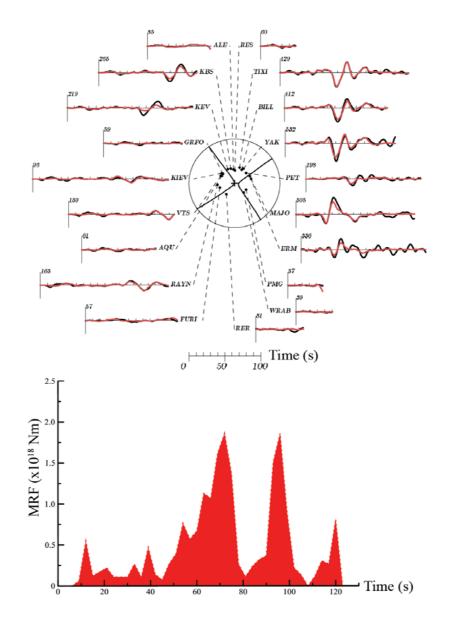




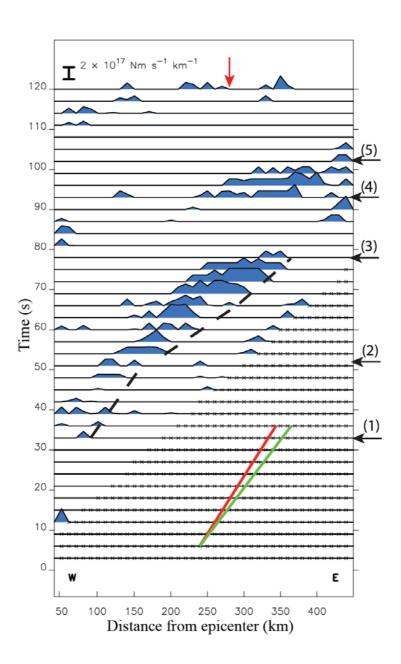
Source Duration Limitations

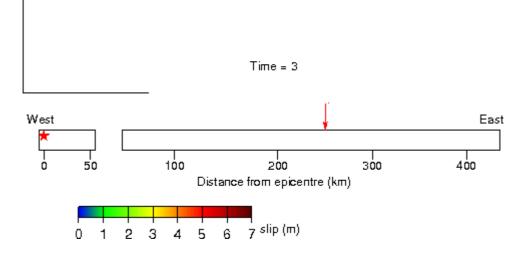


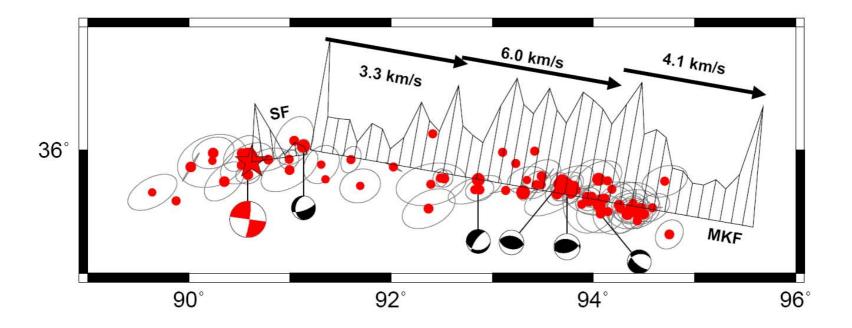
SH waves

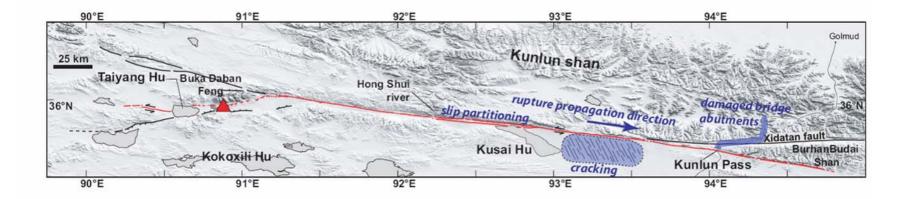


Preferred solution

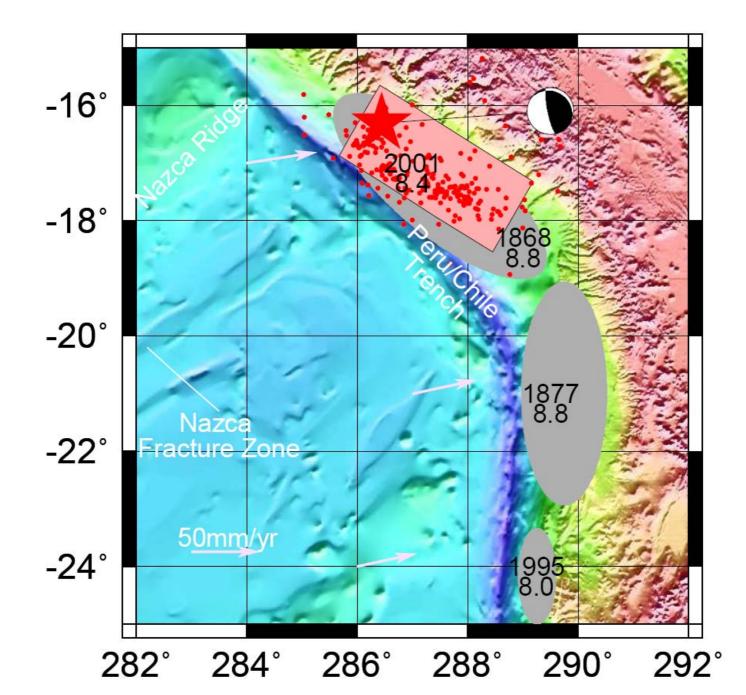




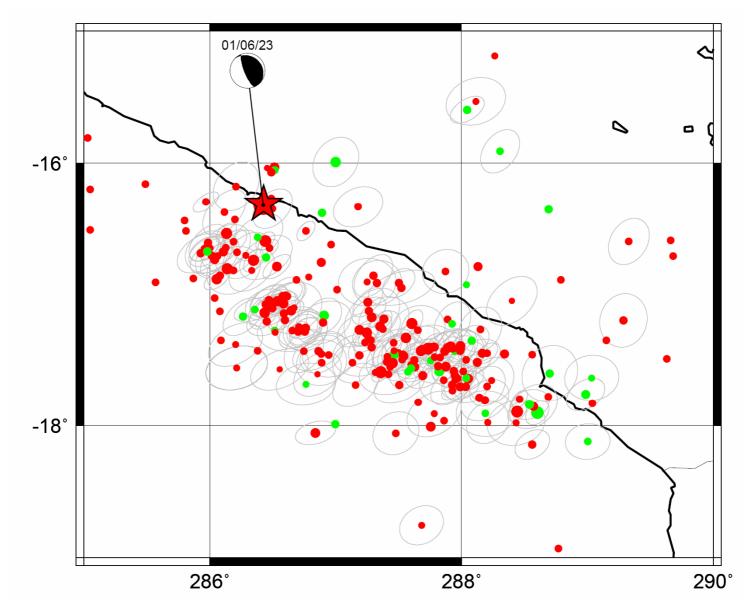




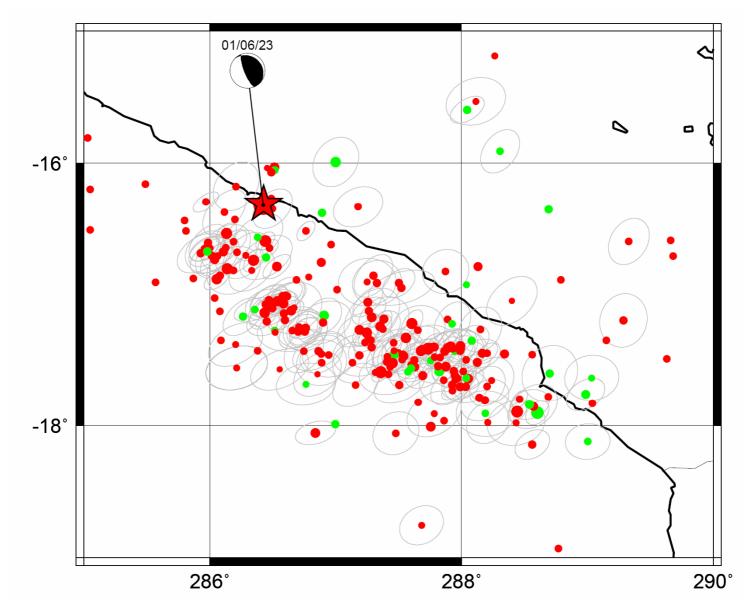




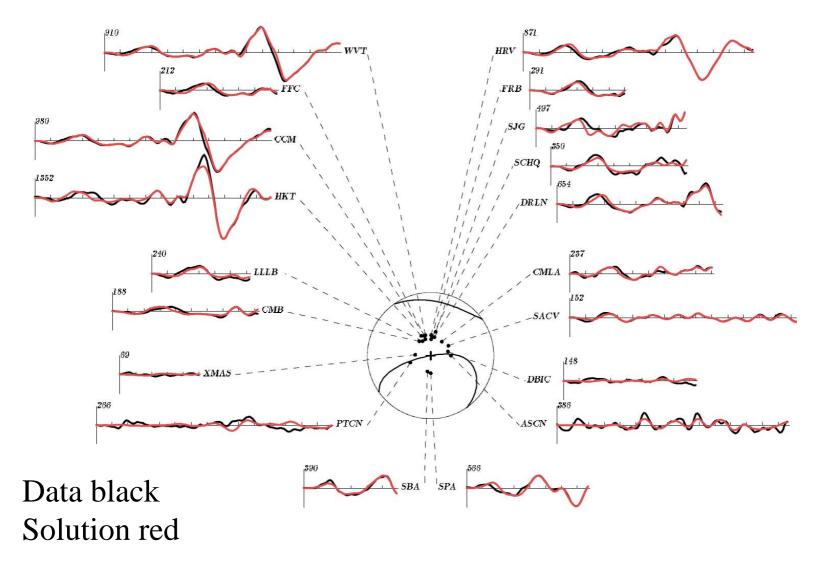
24 hour aftershocks

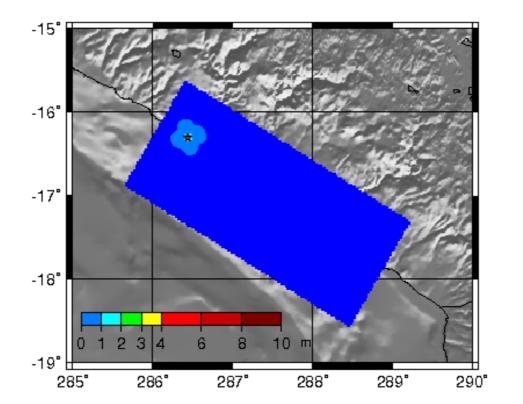


24 hour aftershocks



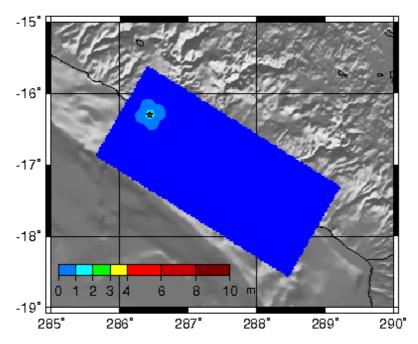
SH waves

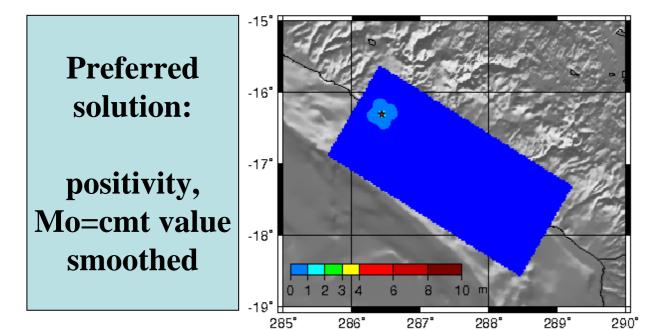


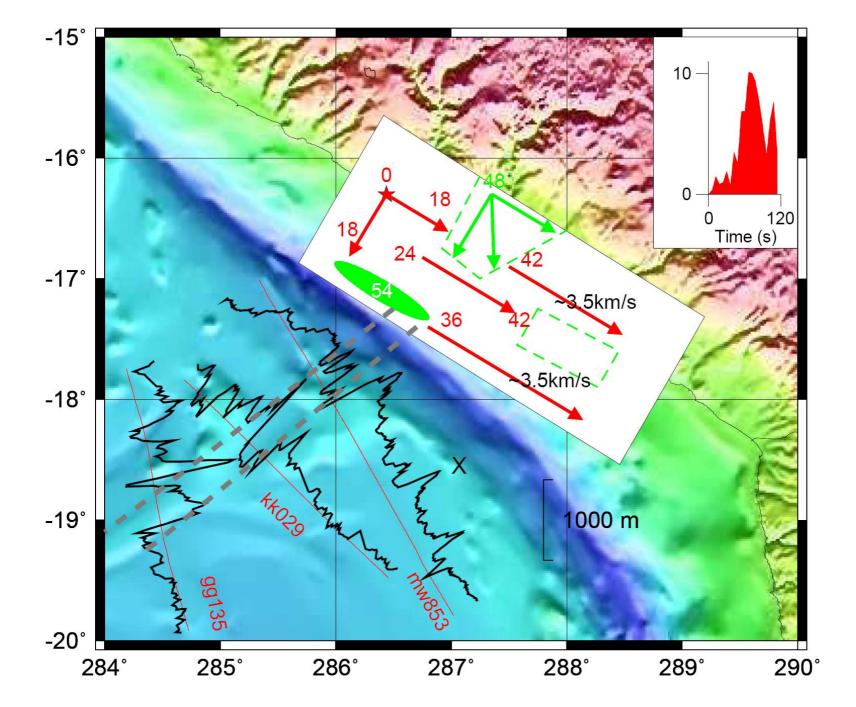


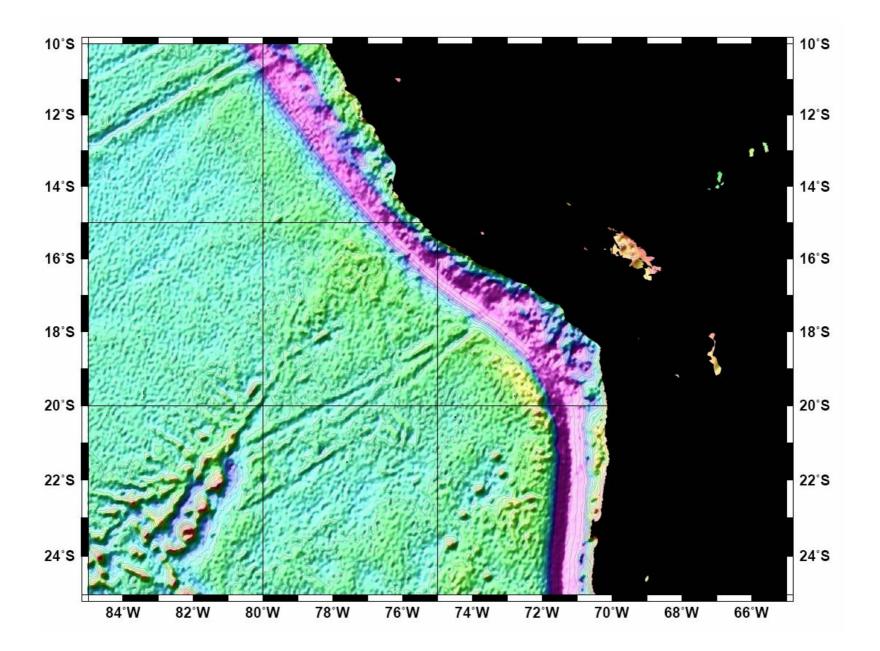
Time = 6

Solution in which we try to maximize slip on barrier during 0-60s

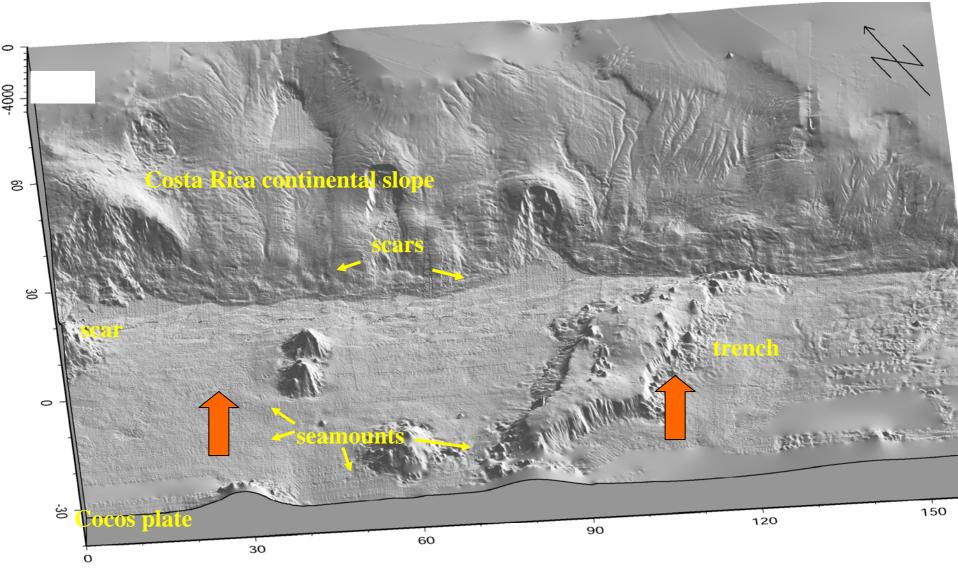


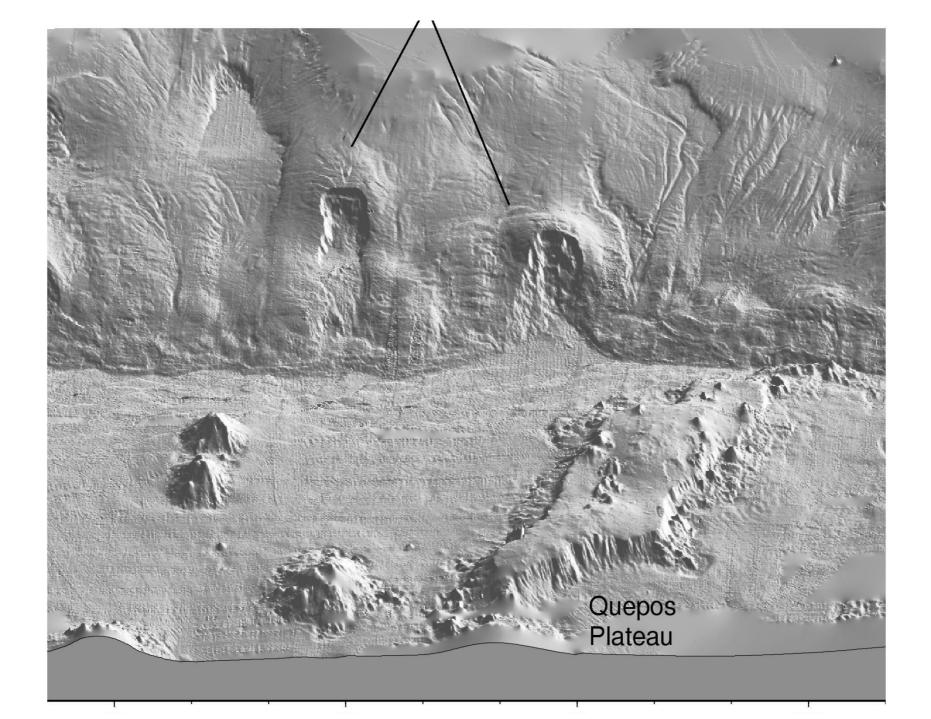


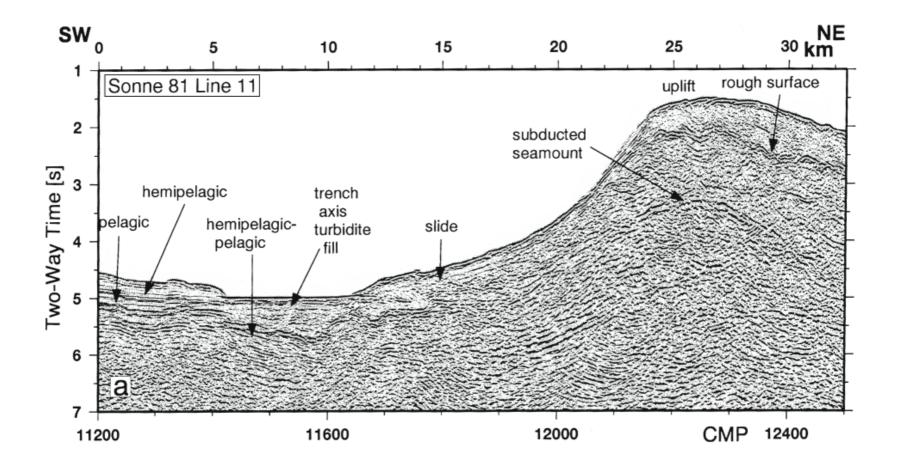




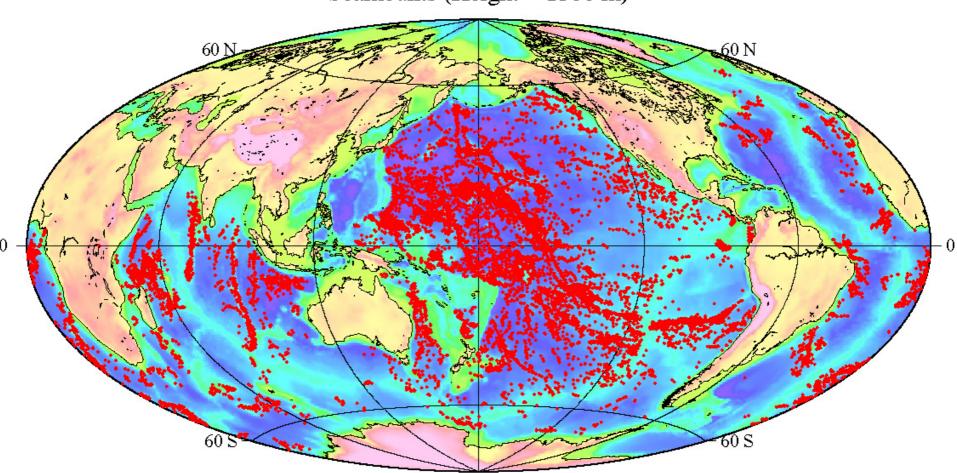
Seamount subduction along Costa Rica (von Huene et al.)



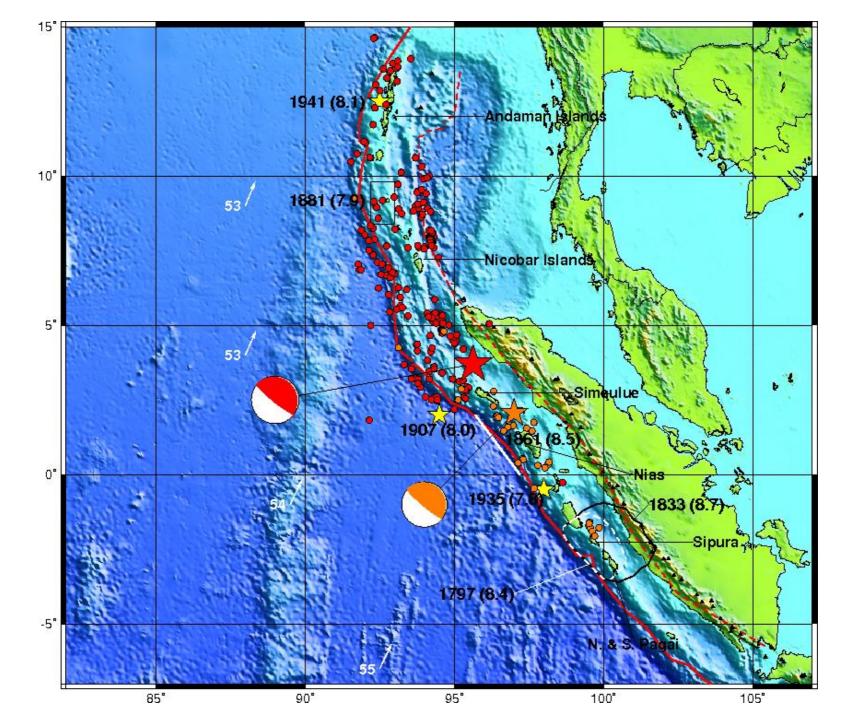


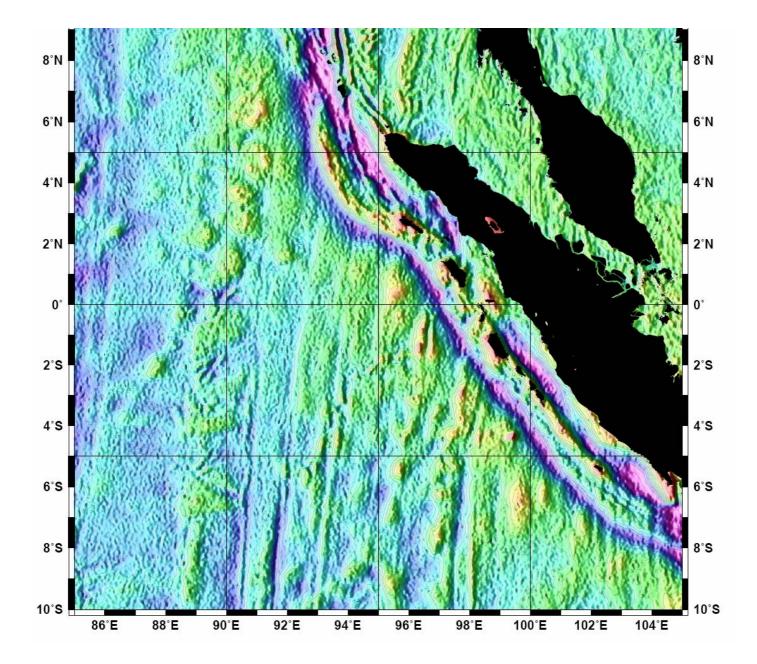




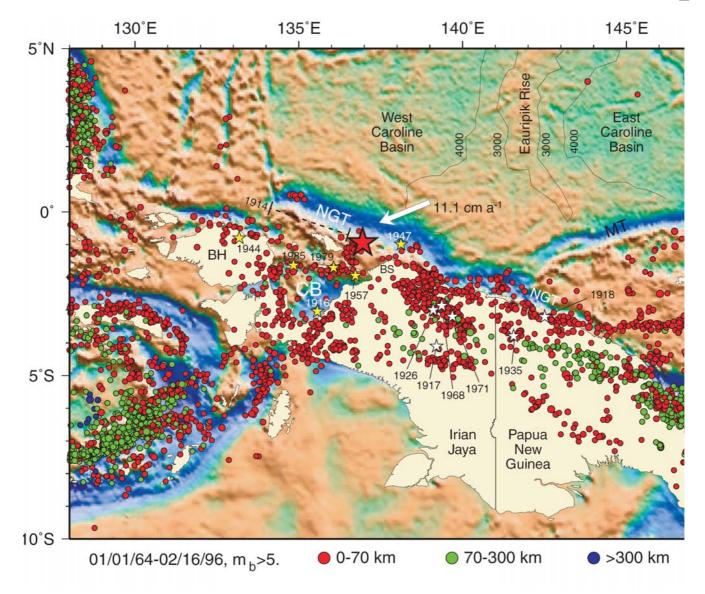


Source: Satellite altimetry (Wessel, 2001)

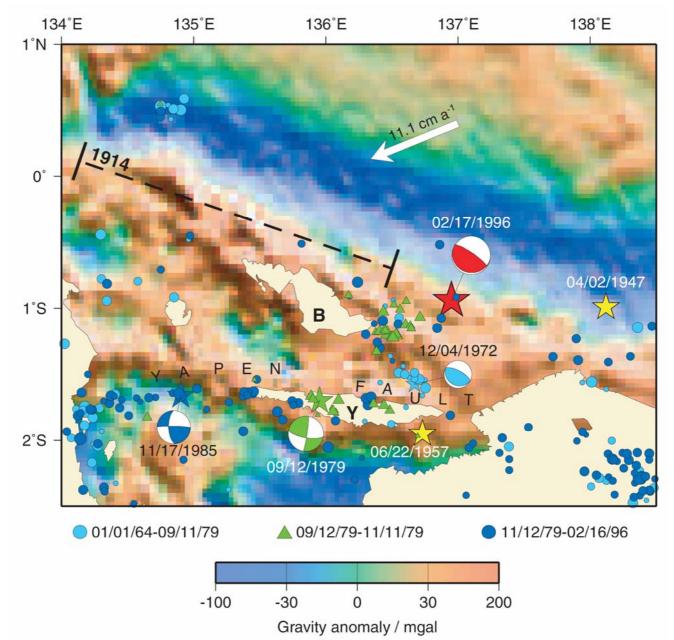




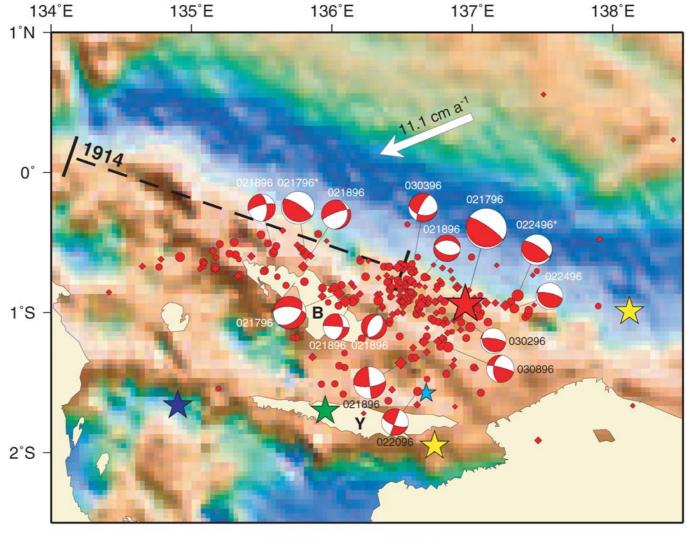
The Mw 8.2 1996 Biak, Indonesia earthquake



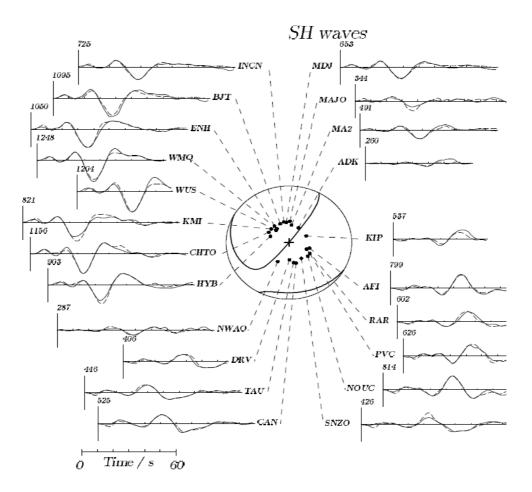
Reliably relocated seismicity prior to earthquake

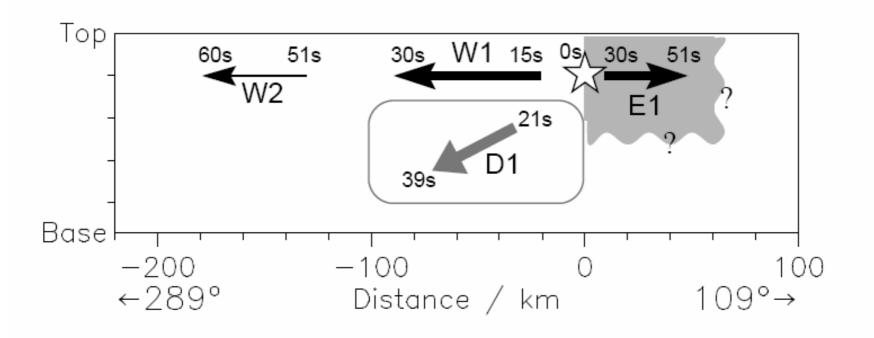


Reliably relocated 1-month aftershocks

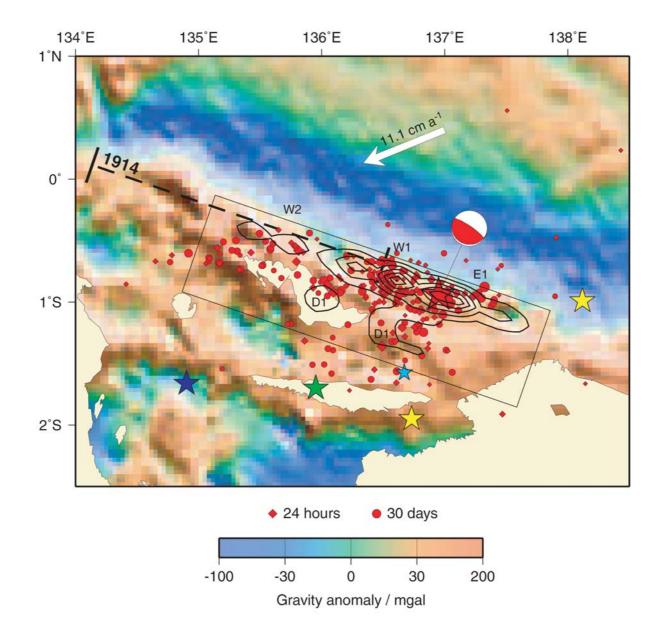


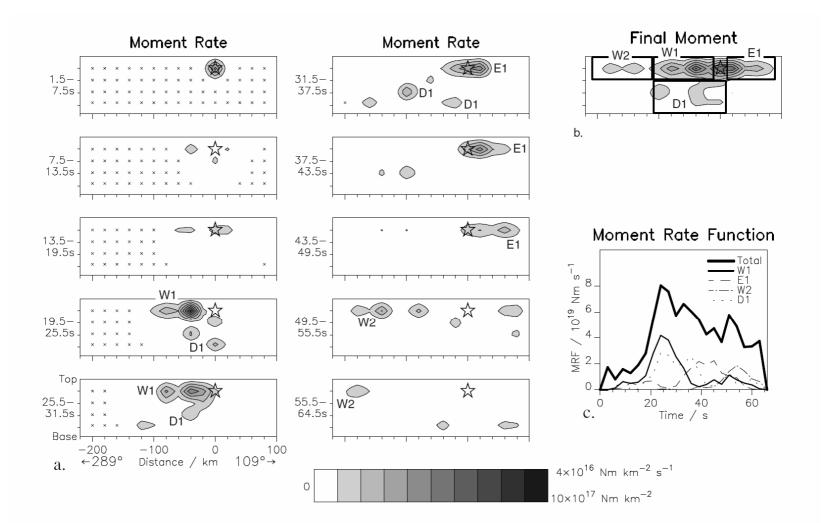
24 hours30 days

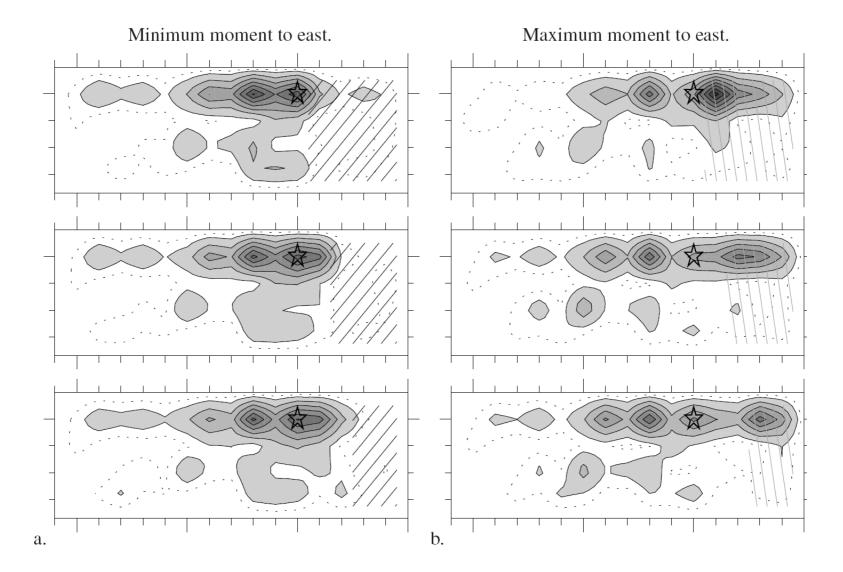




Final moment distribution (contours at 10²⁰ N mkm⁻²⁾

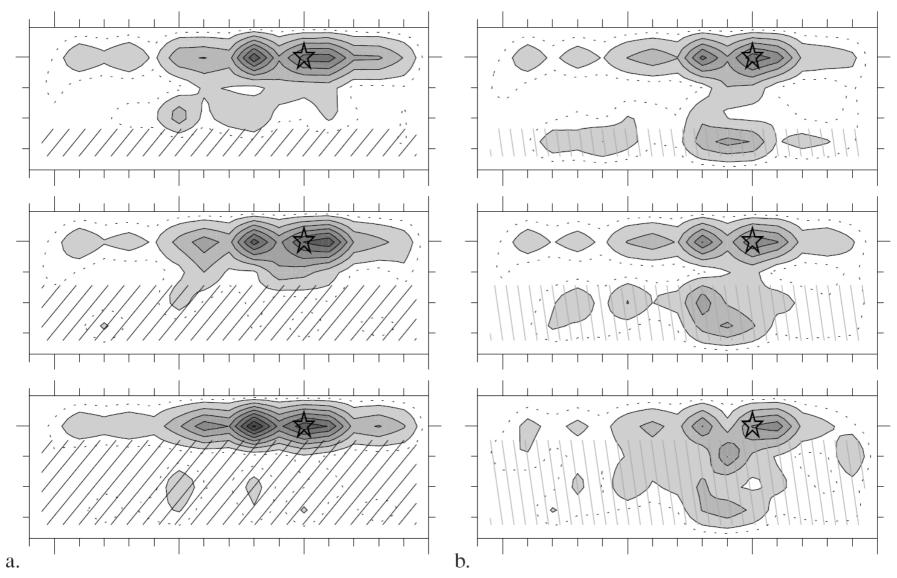


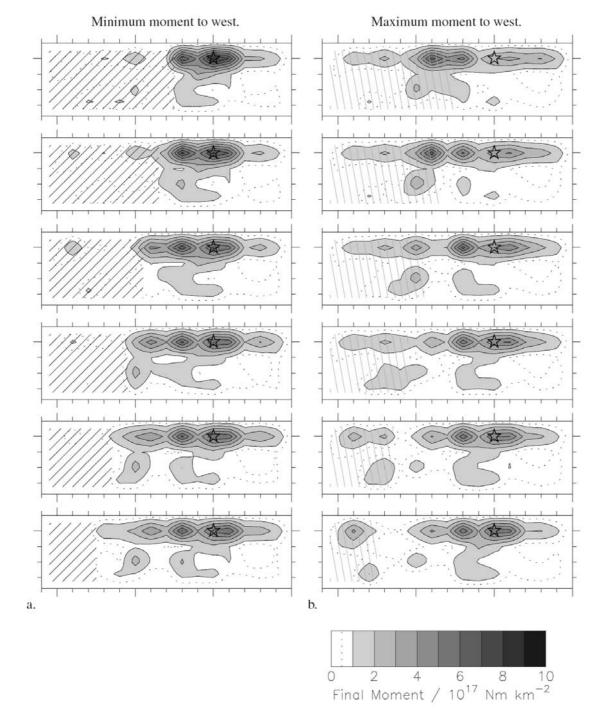




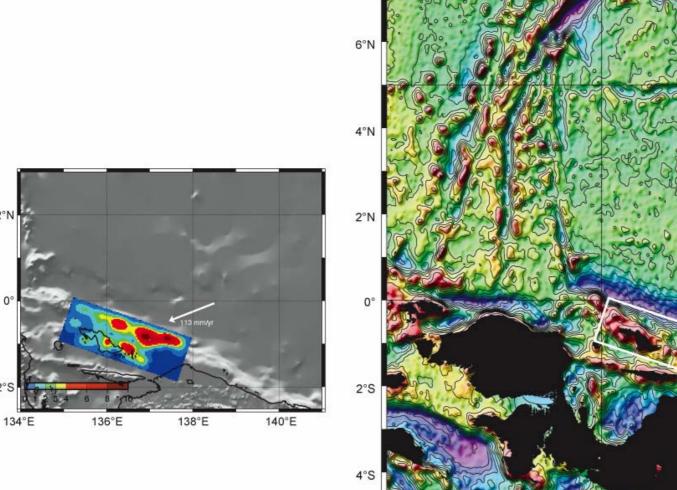
Minimum moment near base.

Maximum moment near base.





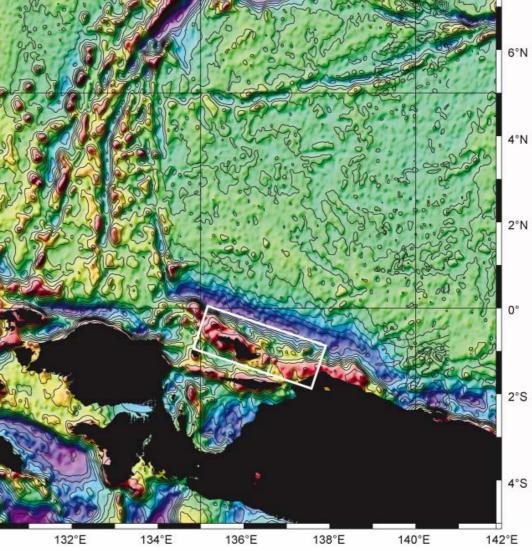
Mw 8.2 1996 Biak, Indonesia earthquake



130°E

2°N

0



Earthquake rupture stalled by a subducting fracture zone

D. P. Robinson, S. Das & A.B. Watts

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